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季节非对称升温对喀斯特土壤 CO, 释放的影响

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摘要:全球变暖呈现季节非对称升温特征,若在研究全球变化对生态系统的影响时未充分考虑该特征,很可能导致研究结果失真。基于红外线辐射增温法,野外模拟不同升温情景下喀斯特土壤 CO_2 释放的短期(4 a) 特征。升温情景包括不升温(对照)、对称升温(全年同步升温 2.0%)和非对称升温(冬春/夏秋季升温幅度为 2.5%/1.5%、3.0%/1.0%、3.5%/0.5%和 4.0%/0%)。结果表明,与对照相比,升温样地土壤 CO_2 通量显著提高,增加了 $0.26~\mu$ mol·(m^2 ·s) $^{-1}$,增幅为 17.41%,其中冬春季通量增加了 $0.23~\mu$ mol·(m^2 ·s) $^{-1}$.在平均升温 2.0%情景下,土壤 CO_2 释放的温度系数(Q_{10})变幅为 $1.53\sim3.24$ 之间,平均值为 2.23。对称升温处理中夏秋季土壤 CO_2 通量升温贡献率(80%)远高于冬春季(20%);非对称升温处理夏秋季和冬春季平均升温贡献率相当(46% 和 54%).5 个升温情景下 CO_2 通量和 Q_{10} 呈现随升温的非对称性增加而降低的趋势,其中对称升温处理 CO_2 通量显著高于中度、高度和极端非对称升温处理。各处理中,夏秋季 Q_{10} 均大于冬春季,这可能与土壤含水量、土壤微生物、可溶性无机碳和植被生长等有关。研究揭示,基于对称升温情景可能会高估全球变暖对喀斯特土壤 CO_2 释放的影响。

关键词: 非对称升温; 对称升温; 温度系数(Q_{10}); 土壤 CO_2 通量; 喀斯特

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Effects of Seasonal Asymmetric Warming on Soil CO₂ Release in Karst Region

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Abstract: Seasonal asymmetric warming is one of the distinguishing features of global warming. However, if this feature is not considered in studying the effects of global changes on terrestrial ecosystems, it might probably cause misunderstanding of these studies. The releasing features of soil CO2 in Karst regions under various warming scenarios were simulated following a four-year continuous warming period using infrared radiators. A total of six treatments was arranged: no warming (ambient temp, CK); symmetric warming (ambient plus 2.0% full year, SW); and, lowly, moderately, highly, and extremely asymmetric warming (ambient plus 2.5%/1.5%, 3.0%/1.0%, 3.5%/0.5%, and 4.0%/0% in the winter-spring/summer-autumn seasons, respectively, LAW, MAW, HAW, and EAW). The results showed that compared to CK, soil CO_2 efflux in all the warming plots significantly increased by 0. 26 μmol·(m²·s)⁻¹, or 17. 41%. In the winter-spring seasons, soil CO₂ efflux in the warming treatments increased by 0.23 µmol·(m²·s)⁻¹. The Q₁₀ values ranged from 1.53 to 3.24 with an average of 2.23 under the scenario of warming up by 2.0°C. The warming-induced contribution of CO₂ efflux in the summer-autumn seasons (80%) was obviously higher than that in the winter-spring seasons (20%) in the SW treatment, whereas the mean contribution in the summer-autumn seasons (46%) was closer to that in the winter-spring seasons (54%) in the asymmetric warming treatments. Both soil CO₂ efflux and Q_{10} showed a tendency towards decrease with the increase in the asymmetry of warming under the five warming scenarios. The soil CO2 efflux in the SW treatment was significantly (P < 0.05) higher than those in the MAW, HAW, and EAW treatments. The Q_{10} values in the summer-autumn seasons was larger than those in the winter-spring seasons under each warming treatment or across all warming treatments, which was probably related to soil water content, soil microbe, dissolved inorganic carbon, and vegetation growth. The results revealed that it may potentially overestimate the effects of global warming on soil CO₂ releasing subject to symmetric warming. **Key words**: asymmetric warming; symmetric warming; temperature coefficient (Q_{10}); soil CO₂ efflux; Karst

受自然因素和人为活动的影响, 1880~2012 年间全球表面气温升高了 0.85℃, 尤其是地表温度 [0.086~0.095℃·(10 a) ⁻¹], 而且未来气候还将持续变暖^[1]. 近 100 年来我国地表气温显著升高,升温幅度略高于全球或北半球平均值,尤其是近 50年以来^[2]. 关于升温过程, Karl 等^[3]分析 1951~1990 年间全球气候资料指出, 北半球大部分地表最低温上升幅度是最高温的 3 倍. 研究证实了气候变

暖存在明显的季节差异,即季节非对称升温,表现为冬春季升温幅度大于夏秋季^[1,2,4].

土壤呼吸(包括自养和异养呼吸)释放 CO₂,是 陆地生态系统碳循环的重要环节,影响全球气候.

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干旱、喀斯特地区钙质土中存在的大量无机碳(如 碳酸钙) 也可能是土壤潜在 CO, 释放源[5~7]. Ramnarine 等[6]通过培养试验发现钙质耕作土释放 的 CO, 中 62%~80%来源于无机碳. 作为陆地生 态系统中生物和非生物过程的调控因子, 温度变化 往往会引起土壤过程和功能变化, 包括土壤呼 吸[8~13]. 研究表明全球变暖或模拟升温能促进土壤 呼吸,释放更多的 CO2,进而可能加剧全球变 暖[10,11],但促进幅度因生态系统类型、增温方式和 幅度及增温季节与持续时间的不同而存在很大差 异, 甚至敏感性随温度的升高而降低[14], 使气候变 暖对土壤呼吸或 CO, 通量的影响机制异常复杂[15], 再者笼统地基于平均温变化(而不是非对称升温) 的估算可能在一定程度上误估生态系统对气候变暖 的反馈[10,16]. 王学娟等[13]利用开顶箱增温法提高 苔原生态系统生长季(6~9月)温度,发现短期(3 a)增温改变了土壤微生物群落结构,但并未明显改 变与碳循环相关的酶活性、土壤微生物生物量和异 养呼吸.

受青藏高原和东亚季风、西南季风的共同影响,地处低纬高原的滇中喀斯特地区气候变化更明显、更复杂,属于气候变化敏感区^[17,18],如滇中昆明地区1961~2007 年平均温升高了 2.08℃^[18].然而并不清楚气候变暖,尤其是非对称升温情景下,该地区土壤 CO₂ 释放的变化情况,也不清楚这种变化对土壤碳源/汇的影响。而正确评估其影响将为生态模型的参数估算与验证提供参考和依据^{15]}.基于此,笔者采用红外线辐射增温法,对喀斯特小样地进行人工控制增温,对比研究对称升温和非对称升温情景下土壤 CO₂ 释放特征,以期为评估未来气候变暖情景下该地区土壤碳源/汇演变趋势提供科学依据.

1 材料与方法

1.1 研究区概况

研究区位于云南省石林县境内(24°40′N,103°22′E),海拔为1750~1800 m,属亚热带季风气候区,是高原喀斯特的典型代表.该区域年均温为15.8℃,最高月均温为20.9℃(6月),最低月均温为8.6℃(1月).年均降水量为948 mm,雨季(5~10月)降水量占全年的80%左右.1961~2016年(数据来源于石林县气象局,其气象站距离本研究区0.8 km),该区域年平均升温幅度为0.29℃·(10 a) $^{-1}$,其中冬春季为0.44℃·(10 a) $^{-1}$,夏秋季为

0. 14° C • (10 a) $^{-1}$.

2012 年,在研究区内选取面积为 1 ha 的石漠 化生态系统(非农地),该区域立地条件相对一致,地势平坦,最大相对高差不足 5 m,地表岩石出露率约为 17%,土壤类型为钙质红色石灰土.区内乔灌木已被破坏,多低矮植物如绣线菊(Spiraea salicifolia)、鬼针草(Bidens pilosa)、苦刺花(Sophora viciifolia),植被高度低于样地大部分出露岩石.2012 年 8 月,该区域植被盖度约为 35%.样地土壤容重为 1.21 g·cm⁻³,土壤总孔隙度为 54.34%,最大持水量为 374.6 g·kg⁻¹,有机碳含量为 14.75 g·kg⁻¹,土壤无机碳(以碳酸钙计)含量为 6.81%(重量比),土壤 pH 值为 7.5.

1.2 试验设计

于2012年8月在石漠化生态系统的固定样地 内设置 18 个 2 m×2 m 小区(小区内相对高差不超 过40 cm, 且岩石出露率不超过10%, 小区间距大 于2 m),安装升温装置和温度监测探头(Decagon, DC, 美国), 开展升温预备试验. 采用红外线辐射 增温法,该方法是目前国内外生态系统小样地控制 试验中模拟增温的主流方法[9,14,19,20],即在每个小 区内悬挂2根功率可调的红外线辐射增温管(增温 管最大功率为 2.0 kW, 长度为 2 m, 悬挂高度为 1.2~1.5 m),对试验小区地表进行人工增温. 由 于地表土壤性质、植被覆盖、微地形等差异,通常 增温幅度以地表上方 3~10 cm 空气温度为基 准[14,20], 本研究以地表上方 5 cm 处气温为基准 (若无特殊说明,下同). 在每个试验小区内随机安 设7个温度探头 (T_1) ,小区外(距离小区至少0.5m)安设5个温度探头(T_0),裸露岩石处不安设,安 装高度离地表 5 cm. 温度探头与高精度自动控温器 (SMUL4, 精度为±0.01℃, 国产)电连, 利用控温 器实时自动调节增温管输出功率, 使实际增温幅度 (*T*₁ - *T*₀)与试验设定的增温幅度相同(±0.01℃误 差范围内). 运行3个月的预备试验显示:试验增温 幅度不超过 4.0℃时,实际增温幅度($T_1 - T_0$)与试 验设定增温幅度的相对误差小于5%,能满足本研 究增温需要.

于 2012 年 12 月在 18 个小区内开展升温试验,设置不升温(对照, CK)、对称升温(全年同步升温 2.0°C, SW)和非对称升温处理,其中非对称升温包括低度非对称升温($\Delta 2.5$ °C/ $\Delta 1.5$ °C, 12 月 ~ 翌年 5 月升温 2.5°C, 6 ~ 11 月升温 1.5°C, LAW)、中度非对称升温($\Delta 3.0$ °C/ $\Delta 1.0$ °C, 12 月 ~ 翌年 5 月升

温 3.0℃, 6~11 月升温 1.0℃, MAW)、高度非对 称升温(Δ3.5℃/Δ0.5℃,12 月~翌年5月升温 3.5℃,6~11 月升温0.5℃, HAW)和极端非对称 升温($\Delta 4.0$ ℃/ $\Delta 0$ ℃, 12月~翌年5月升温4.0℃, 6~11 月不升温, EAW). 本研究中所指升温幅度 均以地表上方 5 cm 处空气温度为基准, 升温期为 2012年12月至2016年11月. 每处理3个小区,即 3次重复,采取完全随机区组设计.本研究设定的 各处理升温幅度均为2.0℃(对照除外),主要是由 于过去 50 年滇中地区升温幅度为 0.2℃・(10 a) -1[18], 据此预测未来 100 年该地区升温幅度将达 2.0℃,这相当于 IPCC[1] 所预测的较低的全球升温 情景:IPCC 第五次评估报告指出,基于由浓度驱动 的 CMIP5(耦合模式比较计划第五阶段)模式模拟 得出,与1986~2005年相比,预估4个气候变化情 景下 2081~2100 年全球平均表面升温范围将分别 为 0. 3 ~ 1. 7、1. 1 ~ 2. 6、1. 4 ~ 3. 1 和 2. 6 ~ 4. 8℃ (RCP 2.6、RCP 4.5、RCP 6.0 和 RCP 8.5). 需要 说明的是气象学上春季是指3~5月、夏季6~8 月、秋季9~11月、冬季12月~次年2月.//

本研究属于生态系统小样地试验, 试验小区较 多(18个), 为持续研究较长时间段内土壤 CO₂ 通 量变化情况,采用成熟稳定可靠、适合长期研究的 碱液吸收法测定 CO, 通量, 即 2015 年 12 月(持续 升温3 a 后), 在每个小区随机安设3个具塞 PVC 管(高度 15 cm, 内径 15 cm), PVC 管另一端插入 土体 3 cm 左右. 在 PVC 管内放一盛有 50 mL 1 mol·L NaOH 溶液的吸收瓶(吸收土壤释放的 CO₂), 密闭. 每隔 10 d 左右更换吸收瓶, 测定其溶 液中的 CO₂-C, 换算成期间的土壤 CO₂ 通量, 即每 个月测定 3 次土壤通量. 试验期为 2015 年 12 月至 2016年11月(共12个月). 同时设置3个空白试 验,即 PVC 管底部封闭,仅放置盛有 NaOH 溶液的 吸收瓶. 在每个小区内安装 5~6 个土壤温湿度监 测探头(ZK-ZD10A, 国产), 自动监测土壤 15 cm 处温湿度,每2h输出一组土壤温度和含水量 数据.

1.3 采样和分析方法

分别于2016年4月和10月,在每个小区内按 "S"形多点(不小于5个)采集表层土样(0~15 cm),采样时除去土表凋落物,小区内土样混合形成一个混合样.每处理采集土壤混合样3个,共18 个.土样采回后,用手选法除去活体根系和可见植物残体,碾碎过10目筛,混匀,调节含水量至田间 最大持水量的 40%,用于测定土壤微生物生物量碳、氮和磷(MBC、MBN 和 MBP),以及可溶性有机碳(DOC)和无机碳(DIC),分析测试前在室温条件下预培养 7 d^[21]. 取部分样品在室内风干、磨细过100 目筛,用于测定其他性质.

NaOH 溶液吸收的 CO_2 -C 以及土壤 DOC 和 DIC 采用碳-自动分析仪(Phoenix-8000)测定,其中 CO_2 -C 分析方法详见文献[21], DOC 和 DIC 提取方法详见文献[22]; MBC、MBN 和 MBP 用熏蒸提取法测定^[21];土壤 pH 值(土水比为 1:2)用复合电极测定.

1.4 统计分析

碳-自动分析仪测定的 CO_2 -C 换算成同期土壤 CO_2 通量,再计算冬春季、夏秋季和全年土壤平均 通量.用温度系数(Q_{10})描述土壤 CO_2 通量对温度 的敏感性,公式为:

$$Q_{10} = \left(\frac{E_{\rm W}}{E_{\rm C}}\right)^{\frac{10}{\Delta T}} \tag{1}$$

式中, $E_{\rm w}$ 和 $E_{\rm e}$ 分别为升温和对照处理中土壤 ${\rm CO}_{\rm e}$ 通量, ΔT 为升温幅度.

用升温贡献率(CP)描述不同季节土壤 CO₂ 通量增量占总增量的比例,公式为:

$$CP = \frac{E'_{W} - E'_{C}}{2 \times (E_{W} - E_{C})} \times 100\%$$
 (2)

式中, E'_{w} 和 E'_{c} 为冬春季或夏秋季升温和对照处理中土壤 CO_{2} 通量,由于 E_{w} - E_{c} 为全年通量增量的平均值,故需统一时间(不同季节的通量增量除以2).

用单因素方差分析(One-way ANOVA, LSD 法) 检验不同处理之间土壤温湿度、土壤可溶性碳、微 生物生物量、土壤通量等参数差异的显著性,用线 性回归分析土壤温度、土壤含水量与 CO₂ 通量的相 关性,以上分析在 SPSS 17.0 软件上完成.

2 结果与分析

2.1 升温对土壤性质的影响

试验升温第 4 年,研究区未升温(对照)样地上方 5 cm 处气温年平均值为 15.4℃,其中夏秋季和冬春季平均温分别为 18.8℃和 12.0℃(表 1). 各升温处理气温升高幅度与试验设定升温幅度完全一致,即全年平均升温 2.0℃,夏秋季和冬春季分别升温 0~2.0℃和 2.0~4.0℃. 土壤 15 cm 处温度升高幅度(2.3~2.4℃)稍大于地表气温,其中夏秋季和冬春季土壤升温幅度分别为 0~2.4℃和 2.5~5.0℃.

表 1	试验区地表气温和土壤温湿度1)

Table 1	Air temperature	soil temperature	and soil water	content of the	experimental plots
rabie i	Air temperature.	son temperature,	and son water	coment of the	experimental blots

AL TH	1× -k- *k-	ţ	也表气温/℃	C	=	上壤温度/℃	CJ	土	壤含水量/	%	рŀ	I
处理	样本数	全年	夏秋季	冬春季	全年	夏秋季	冬春季	全年	夏秋季	冬春季	夏秋季	冬春季
对照	3	15.4 b	18.8 e	12. 0 f	18.5 b	21.8 е	15. 1 f	30. 2 a	44. 6 a	15. 8 a	7.4 a	7.6 a
对称升温	3	17.4 a	20.8 a	14.0 a	20.9 a	24. 2 a	17.6 e	28. 3 b	44. 1 a	12.5 b	7.1 b	7.5 a
低度非对称升温	3	17.4 a	20. 3 b	14.5 b	20.8 a	23.5 b	18.2 d	28.4 b	44.7 a	12. 1 bc	7.1 b	7.5 a
中度非对称升温	3	17.4 a	19.8 c	15.0 с	20.9 a	23.0 с	18.8 c	28.0 b	44. 2 a	11.8 bc	7. 2 ab	7.6 a
高度非对称升温	3	17.4 a	19.3 d	15.5 d	20.9 a	22.4 d	19.5 b	27.6 bc	44. 3 a	$10.\ 8\ \mathrm{cd}$	7. 3 ab	7.6 a
极端非对称升温	3	17.4 a	18.8 e	16.0 e	20.9 a	21.8 e	20. 1 a	27.3 с	44. 5 a	10. 1 d	7.4 a	7.6 a

1) 地表气温为样地上方 5 cm 处空气温度, 土壤温湿度为土壤 15 cm 处的温度和含水量; 各列不同字母表示处理间均值差异达 95% 显著水平, 下同

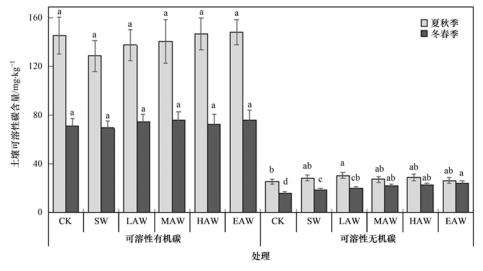
持续升温第 4 年,研究区内降水量为1 024.5 mm,其中夏秋季降水量为 821.4 mm,占全年降水量的 80%以上.升温处理中土壤年平均含水量在27.3%~28.4%之间,显著低于对照,为对照的90%~94%(表1).夏秋季各处理土壤平均含水量接近,在 44.5%左右.冬春季土壤平均含水量随同期升温幅度提高而降低,其中升温样地含水量比对照低 21%~36%.冬春季增温(地表气温增加 2.0~4.0℃、土壤增温 2.5~5.0℃)加剧了土壤水分蒸发^[23];夏秋季频繁降水有效补充了土壤水分,在一定程度上抵销了增温对土壤水分蒸发的影响,而且夏秋季增温幅度低于冬春季.

夏秋季升温能显著降低对称升温和低度非对称升温处理中土壤 pH 值,对冬春季土壤 pH 的影响不显著(表1). 夏秋季土壤 pH 略低于冬春季,约低0.2~0.4个单位.

无论采样季节,不同处理间 DOC 含量差异均

不显著(图 1). 同一处理中夏秋季 DOC 含量约为 冬春季的 1.9~2.1 倍. 除 LAW 处理中 DIC 含量 显著高于 CK 外,夏秋季其他处理间 DIC 含量差 异均不显著. 而冬春季 DIC 含量随增温幅度的提高而增加,其中对照处理 DIC 含量显著低于各升温处理.

持续升温对土壤微生物生物量的影响因季节而异(表2). 在夏秋季, SW 处理中土壤 MBC、MBN和 MBP含量均显著高于 CK, 是后者的 1.1~1.2倍. 此外, 在夏秋季 LAW和 MAW 处理土壤 MBN和 MBP含量均显著高于 CK. 而在冬春季, 升温对土壤 MBC、MBN和 MBP的影响均不显著. 总体上, SW 处理土壤 MBC、MBN和 MBP含量显著低于非对称升温处理. 无论是冬春季或夏秋季, 升温对微生物生物量 C/P比(MBC/MBP)的影响较小,各处理间差异均不显著,比值在 37.5 左右. 但升温可以显著降低 SW和 LAW处理中夏秋季微生物生物量



CK、SW、LAW、MAW、HAW、EAW:对照、对称升温、低度非对称、中度非对称、高度非对称、 极端非对称升温处理;不同字母表示处理间均值差异达95%显著水平

图 1 不同升温情景土壤可溶性有机碳和无机碳含量

Fig. 1 Soil dissolved organic C and inorganic C under different warming scenarios

表 2 不同升温情景下土壤微生物生物量碳、氮、磷及碳氮比和碳磷比

Table 2 Microbial biomass C, N, P, C: N and C: P under different warming scenarios

处理	微生物生物量	碳/mg•kg -1	微生物生物量	፟፟፟፟፟፟፟፟፟፟፟፟፟፟፟፟፟፟፟፟፟፟፟፟፟፟፟፟፟፟	微生物生物量	遺磷/mg⋅kg - 1	微生物生物	勿量碳氮比	微生物生物	可量碳磷比
处理	夏秋季	冬春季	夏秋季	冬春季	夏秋季	冬春季	夏秋季	冬春季	夏秋季	冬春季
对照	285. 2 ± 27. 6 b	75. 2 ± 6. 8 a	19.4 ± 1.9 c	5. 1 ± 0. 5 a	7.6 ± 0.5 c	2.0 ± 0.2 a	14.7 ± 0.7 a	14. 8 ± 0. 5 a	37. 5 ± 1. 4 a	37. 6 ± 1. 3 a
对称升温	357. 1 ± 36.3 a	80. 5 ± 6.8 a	28.4 ± 2.0 a	$5.5 \pm 0.5 \text{ a}$	9. 5 ± 0.4 a	2.1 ± 0.2 a	12. 6 \pm 0. 9 b	14.6 ± 0.5 a	37. 6 ± 2.8 a	$38.4 \pm 1.5 \text{ a}$
低度非对称升温	327. 6 ± 30 . 4 ab	82. 4 ± 3.9 a	26. 1 $\pm2.$ 2 ab	$5.8 \pm 0.4 \text{ a}$	$8.7\pm0.4~\mathrm{ab}$	2.2 ± 0.2 a	12. 6 \pm 0. 4 b	14.2 ± 0.7 a	37.6 ± 1.7 a	$37.6 \pm 1.8 \text{ a}$
中度非对称升温	320.4 ± 29.6 ab	77. 8 ± 4.3 a	23. 0 \pm 2. 0 b	5.1 ± 0.4 a	$8.9\pm0.4~\mathrm{ab}$	2.1 ± 0.2 a	13. 9 ± 0.1 ab	15.3 ± 0.6 a	$36.0 \pm 1.4 \text{ a}$	37. 1 ± 2.0 a
高度非对称升温	300. 2 \pm 26. 1 b	80. 7 ± 4.5 a	$20.~8\pm1.~9~\mathrm{c}$	$5.8 \pm 0.5 \text{ a}$	$8.~1\pm0.~5~\mathrm{bc}$	2.2 ± 0.2 a	14.4 ± 0.9 a	13.9 ± 0.6 a	37. 1 ± 1. 1 a	$36.7 \pm 1.6 \text{ a}$
极端非对称升温	288. $8 \pm 23. 8 \text{ b}$	86.7 ± 7.5 a	20. $0 \pm 1.8 \text{ c}$	$6.0 \pm 0.5 \text{ a}$	$7.~8\pm0.~5~\mathrm{bc}$	$2.3 \pm 0.2 \text{ a}$	14. 4 ± 0. 4 a	14. 4 ± 0. 3 a	37.0 ± 1.7 a	37.7 ± 1.2 a

C/N 比(MBC/MBN = 12.6), 而其他处理或季节 MBC/MBN(14.5 左右)对升温的响应不敏感.

2.2 土壤 CO, 释放特征

持续升温对喀斯特石漠化土壤 CO_2 通量有显著影响(表 3). 总体上,升温样地 CO_2 通量[1.73 $\mu mol \cdot (m^2 \cdot s)^{-1}$] 显著高于 CK [1.47 $\mu mol \cdot (m^2 \cdot s)^{-1}$],SW 处理中 CO_2 通量[1.86 $\mu mol \cdot (m^2 \cdot s)^{-1}$]显著高于非对称升温处理[1.69 $\mu mol \cdot (m^2 \cdot s)^{-1}$]。在不同升温情景模式下,5个升

温处理中土壤 CO_2 通量比 CK 高 $8.8\% \sim 26.5\%$. 升温处理之间,土壤通量呈现随升温的非对称性增加而降低的趋势,即 SW > LAW > MAW > HAW > EAW. 统计结果显示 SW 处理中 CO_2 通量显著 (P < 0.05)高于 MAW、HAW 和 EAW 处理. 土壤 CO_2 通量对不同升温情景的响应和敏感性差异明显(表3). 在平均升温 2.0% 情景下, Q_{10} 表现随升温非对称性增加而降低的趋势,变幅在 $1.53\sim 3.24$ 之间,平均值为 2.23.

表 3 不同升温情景下土壤 CO, 释放特征

Table 3 Releasing features of soil CO₂ under different warming scenarios

-		·孟昌./ 1/2	\ -1//		0	~ 11 \	41 MD	エナトラス・ペ
处理	CO_2	通量/μmol·(m²	·s) ·		Q_{10}	- IL	<u> </u>	贡献率/%
处 连	全年	夏秋季	冬春季	全年	夏秋季	冬春季	夏秋季	冬春季
对照	1.47 ± 0.06 d	2. 08 ±0. 10 d	$0.86 \pm 0.03 \text{ c}$	_	#/ e	1 11	_	$-\langle \cdot \rangle$
对称升温	1. 86 ± 0.06 a	$2.70 \pm 0.12 \text{ a}$	$1.02 \pm 0.02 \text{ b}$	3. 24	3. 69	2. 35	79.49	20.51
低度非对称升温	且 1.79 ±0.08 ab	2. 52 ± 0 . 11 ab	1. 06 ± 0.04 ab	2.68	3. 59	2. 31	68.75	31. 25
中度非对称升温	且 1.72 ± 0.06 bc	2.35 ± 0.10 bc	1.09 ± 0.04 ab	2. 19	3. 39	2. 20	54.00	46. 00
高度非对称升温	且 1.66 ±0.06 c	$2.~20~\pm0.~09~\mathrm{cd}$	1. 12 ± 0.04 a	1. 84	3. 07	2. 13	31.58	68. 42
极端非对称升温	且 1.60 ± 0.06 c	$2.06 \pm 0.09 d$	1. 14 ± 0. 04 a	1.53	30	2. 02	_	107. 69

不同季节,各处理土壤 CO,通量及其对升温情 景的响应差异明显(表3). 在5个升温处理中, 夏 秋季和冬春季土壤 CO₂ 平均通量分别为 2.37 和 1.09 μmol·(m²·s)⁻¹. 在夏秋季, SW 处理中土壤 CO₂ 通量最高[2.70 μmol·(m²·s)⁻¹], 其次为 LAW 处理, MAW 和 HAW 处理次之, EAW 处理的 最低[2.06 μmol·(m²·s)⁻¹]. 在冬春季, 土壤 CO₂ 通量大小顺序则正好与夏秋季相反,SW 处理土壤 通量不足 EAW 处理的 90%. 在同一处理中, 夏秋 季土壤 CO, 通量均显著高于冬春季, 是冬春季 1.8 ~2.6 倍. HAW 和 EAW 处理中土壤 CO2 夏秋季通 量与 CK 差异不显著, 尤其是 EAW 处理, 甚至略低 于 CK. 除此之外, 升温均能显著提高不同季节土 壤 CO₂ 通量, 夏秋季和冬春季提高幅度分别为 14%~30%和19%~33%.升温处理中夏秋季和冬 春季平均升温幅度分别为 1.0℃和 3.0℃,对应的 Q10平均值分别为 3.63 和 2.18. 总体上或在同一处

理中,夏秋季 Q_{10} 均大于冬春季. 在不同季节, Q_{10} 值大体上均随升温的非对称性增强而降低,如在冬春季 Q_{10} 值由 SW 处理的 2. 35 降低至 EAW 处理的 2. 02.

升温处理中夏秋季土壤 CO₂ 通量升温贡献率 (CP,56%)高于冬春季(44%,表3). 同等升温幅度下,夏秋季 CP(80%)远大于冬春季(20%). 而非对称升温处理中夏秋季和冬春季平均 CP 为 46%和 54%. 在非对称升温处理之间,冬春季 CP 随同期升温幅度提高而增加,即由 LAW 处理的 31%提高到 EAW 处理的 107%. 当冬春/夏秋季升温比为3.0 时,夏秋季 CP(54%)稍高于冬春季(46%).

3 讨论

3.1 升温对喀斯特土壤 CO, 释放的影响

温度是陆地生态系统中生物和非生物过程的调控因子,在一定程度上控制土壤碳收支和循

环^[8~13].通常升温可提高土壤微生物和酶活性,促进 SOC 矿化,也驱动和促进根及根际呼吸^[10,11,13].升温可能提高喀斯特地区碳酸钙等碳酸盐岩的风化和溶解^[24,25].因此,普遍认为土壤呼吸随温度的升高而增加,与气候变暖存在正反馈过程^[10,11].如Rustad等^[26]对云杉-冷杉林持续增温(3a),发现增温幅度为5℃左右时土壤 CO₂ 释放量增加了 25%~40%.本研究用红外线辐射器对喀斯特小样地进行持续增温(4a),结果显示升温显著提高了土壤 CO₂通量,比未升温处理高 17%,表明未来全球变暖将加速我国西南高原喀斯特地区土壤 CO₂的释放.

通常用 Q_{10} 来反映土壤呼吸的温度敏感性. Q_{10} 值越大,则土壤呼吸对温度的敏感性越高,同等温 度变幅对 CO, 通量的影响也越大. 本研究中, 在全 年平均升温 2.0℃ 情景下, 5 个升温处理中土壤 CO, 通量 Q₁₀值变幅在 1.53~3.24 之间(平均值为 2.23),这与我国西南喀斯特地区类似研究结 果[27~29]以及全球尺度上温带、热带/亚热带森林和 草原土壤呼吸 Q_{10} 值接近 $(2.2 \sim 2.7)^{[30]}$,稍高于全 球平均水平(1.7~2.0)[30,31]和纬度略低的我国亚 热带阔叶林(1.86)[32],但明显高于全球沙漠/荒漠 地区(1.43)[31]以及热带稀树草原(1.39)[33]. 据此 推断我国西南高原喀斯特地区对气候变化敏感. 郑 建萌等[17]和 Li 等[18]从该地区气象演变趋势的角度 也得到类似结论. 唐夫凯等[29] 对黔西南喀斯特土 壤呼吸的温度敏感性进行研究, 发现草地、灌草 地、灌丛和次生林土壤呼吸 Q_{10} 值范围为 1.92 ~ 3.67. 而该地区轮作旱地(2.02)[28]和马尾松人工 林(1.92~2.10)^[27]土壤呼吸 Q₁₀值也在 2 左右. 本 研究中喀斯特石漠化生态系统(非农地)土壤 CO, 通量对温度变化较敏感可能与本区较低的气温有关 (多年平均气温为 15.8℃). Kichlighter 等^[34] 总结 他人的研究发现,在低温背景下土壤呼吸 Q_{10} 较高, 在高温下 Q₁₀较为平稳, 温度超过 20℃ 时 Q₁₀一般 在 2 左右, 在 0℃时 *Q*10可高达 8.

通常认为冬春季土壤呼吸对温度变化的敏感性要大于夏秋季^[15,29,31,32,34]. 但本研究升温处理中冬春季土壤呼吸 Q_{10} 值反而小于夏秋季. 这可能与以下因素有关:①土壤含水量. 表 4 显示 6 个处理中夏秋季土壤含水量与 CO_2 通量的相关性(r 为 0.604 ~ 0.687) 明显小于冬春季(r 为 0.779 ~ 0.831); 而夏秋季土壤温度与 CO_2 通量的相关性(r 为 0.889 ~ 0.912) 则大于冬春季(r 为 0.791 ~ 0.825), 这表明夏秋季频繁的降水以及土壤干湿交

替可有效促进土壤呼吸和 CO₂ 释放^[5,35],在水分不 受限制的情况下温度可能是土壤呼吸的主控因子, 因此夏秋季提高温度则能显著促进土壤呼吸. 在冬 春季持续干旱情况下,土壤含水量极低(表1), CO₂ 通量的主控因子可能发生转移, 如转移为土壤 含水量[20,24,29,33,34],因此升温对冬春季土壤呼吸的 贡献有限. ②土壤微生物. MBC/MBN 可反映微生 物群落结构信息,而 MBC/MBP 反映土壤微生物对 磷素有效性调节作用的强弱[36]. 本研究中冬春季 升温对 MBC、MBN 和 MBP 含量, 以及 MBC/MBN 和 MBC/MBP 均无显著影响. 这表明冬春季升温对 土壤微生物呼吸(即异养呼吸)的作用可能有限. 而 夏秋季升温对微生物生物量有不同程度的提高作 用, 且当夏秋季升温幅度不低于 1.5℃ 时 MBC/ MBN 显著降低, 暗示夏秋季升温可以提高土壤微 生物数量和土壤氮素利用率,在一定程度上改变微 生物群落结构^[36],从而促进土壤有机碳矿化和CO, 释放. ③土壤 DIC 含量. 夏秋季土壤 DIC 含量约为 冬春季的1.1~1.6倍(图1),而土壤含水量为冬 春季的3.5~5.4倍(表1),表明冬春季土壤溶液 中 DIC 浓度高于夏秋季. 而土壤溶液中 DIC 的增加 主要与土壤 CO₂ 分压(pCO₂)的提高有关^[37], 尤其 在 pH 值较高(表 1)的背景下[5],进而促进土壤 CO₂ 的吸收和碳酸钙的沉淀^[24], 即发生土壤无机 固碳过程,这可能在一定程度上抑制了冬春季土壤 中 CO₂ 的释放. ④植被生长. 夏秋季是研究区植物 生长季,植物根系活跃,提高生长季的温度通常能 促进根及根际呼吸,即自养呼吸[14,15].

本研究中,尽管冬春季平均升温幅度(3.0℃)远大于夏秋季(1.0℃),但5个升温处理冬春季CO₂通量仅为全年的30%,而且平均升温贡献率小于夏秋季(表3).如 MAW 处理中夏秋季CO₂通量占全年的68.30%,其升温贡献率(54%)超过冬春季(46%).表明夏秋季不仅是喀斯特地区土壤CO₂最大的释放阶段,也将是全球变暖所导致的该地区土壤CO₂释放增加最迅速的阶段.这暗示了该地区夏秋季是人为调控土壤碳排放的关键期.

3.2 非对称升温对喀斯特土壤 CO₂ 释放的影响

气象资料显示全球变暖过程中冬春/夏秋季升温比通常在 3 左右[1.2.4], 研究区内 1961~2016 年冬春/夏秋季升温比为 3.1 [冬春季为 0.44℃ · (10 a) $^{-1}$, 夏秋季为 0.14℃ · (10 a) $^{-1}$, 夏秋季为 0.14℃ · (10 a) $^{-1}$]. 本研究中 SW、LAW、MAW、HAW 和 EAW 处理中冬春/夏秋季升温 比分别为 1.0、1.7、3.0、7.0 和夏秋季不升温

表 4	不同升温情景下土壤 CO。	,通量与土壤温度和土壤含水量的相关性 ¹⁾

7D 11 4	0 1.1 1. 100	CCI 1 11.	. 1 1 .	1 1°CC .	
Table 4	Correlations between soil CO	, ettlux and soil tem	peratures and soil water.	contents under different	warming scenarios

Al rm		土壤温度			土壤含水量	
处理 -	全年	夏秋季	冬春季	全年	夏秋季	冬春季
对照	0. 905 **	0. 912 **	0. 825 **	0. 814 **	0. 687 *	0. 831 **
对称升温	0. 901 **	0. 908 **	0. 791 *	0. 808 *	0. 604 *	0. 818 **
低度非对称升温	0. 889 **	0. 895 **	0. 802 *	0. 799 *	0. 626 *	0. 808 *
中度非对称升温	0. 899 **	0. 901 **	0.814*	0. 787 *	0. 647 *	0. 792 *
高度非对称升温	0. 865 **	0. 897 **	0. 820 **	0. 734 *	0. 664 *	0. 788 *
极端非对称升温	0. 866 **	0. 889 **	0. 818 **	0. 753 *	0. 660 *	0. 779 *

1) 土壤 CO_2 通量数据为单次数值(每 10d 左右 1 次,每个月 3 次);土壤温湿度为土壤 15 cm 处的温度和含水量,其数值为土壤 CO_2 通量单次数值测定期间土壤温湿度的平均值;*表示 P<0.05,**表示 P<0.01

(无尽大),即 MAW 升温情景与本区以及全球气候 变暖的非对称性接近. 表 3 显示 SW 处理中土壤 CO₂ 通量比非对称处理高 4%~16%, 且显著高于 MAW、HAW 和 EAW 处理, 其中比 MAW 处理高 8. 14%. SW 处理 Q_{10} 值(3. 24) 也明显高于非对称 升温处理(1.53~2.68, MAW 处理为2.19). 这表 明基于对称升温可能会高估全球变暖对喀斯特地区 土壤 CO₂ 释放的影响, 进而影响未来全球变化情况 下土壤碳源/汇的判断. 其原因如前文3.1 所述, 研 究区内夏秋季是土壤 CO₂ 释放(本底值大)和受气 候变暖影响(Q_{10} 值大、升温贡献率大)最大的时期, 在年均温增幅相同的情况下, 非对称升温处理中夏 秋季升温幅度小于对称升温,因而非对称性情景下 土壤 CO₂ 通量也小于对称升温. 苏宏新等[10] 运用 模型模拟, 发现当大气 CO, 浓度不变情况下, 基于 平均温度变化的估算可能在一定程度上高估了暖温 带蒙古栎林生态系统呼吸对气候变暖的正反馈.

需要指出的是,本研究为升温初期(4 a)土壤CO₂通量特征,也未考虑全球变暖其他因子的变化,如降水量、大气CO₂浓度等,因此本研究结论还需在更大范围、更长升温期、更复杂的气象环境、更多土壤/植被类型的喀斯特地区验证.

4 结论

(1)我国西南高原喀斯特地区对气候变暖敏感,未来全球变暖将加速该地区土壤 CO_2 的释放. 持续升温第4年,喀斯特土壤 CO_2 通量比不升温处理增加了17%,其中冬春季 CO_2 通量升温贡献率为44%. 在平均升温2.0℃情景下,5个升温处理土壤 CO_2 通量 Q_{10} 值变幅为1.53~3.24,平均值为2.23,其中夏秋季 Q_{10} 值大于冬春季,这可能与土壤含水量、土壤微生物、可溶性无机碳和植被生长等有关.

(2)5个升温情景模式下,CO₂ 通量和 Q₁₀随升温的非对称性增加而降低,其中对称升温处理中CO₂ 通量显著高于中度、高度和极端非对称升温处理.鉴于喀斯特地区冬春/夏秋季升温比约为 3 的情景(相当于本研究中度非对称升温情景),推测基于对称升温情景模式可能在一定程度上高估了全球变暖对喀斯特土壤 CO₂ 释放的影响.

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