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典型温冰川区湖泊的稳定同位素空间分布特征

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摘要:2014年8月对拉市海表层及不同深度湖水进行采样,分析拉市海湖水的氢氧稳定同位素的空间变化及其影响因素,探讨典型温冰川区域湖泊的水文补给特征.结果表明,拉市海表层湖水的8¹⁸0、8D值分别在-12.98%。~8.16%和-99.42%。~73.78%。之间波动,平均值分别为-9.75%。和-82.23%。;表层湖水的8¹⁸0及过量氘表现出相反的空间变化特征,有河水注入的区域8¹⁸0值较低而过量氘值较高;垂直方向上过量氘随深度变化较小,表明湖水在垂直方向上混合较充分,不同深度层上过量氘表现出自东向西先增大后减小的变化趋势,这可能与人湖河流的分布、湖泊所处的地理位置及自然条件等密切相关;同位素对比研究发现,拉市海的主要补给源为大气降水及河水,冰雪融水可能间接补给拉市海;对拉市海与青藏高原地区典型湖泊和非冰川区湖泊的氧同位素组成对比发现,冰川区湖泊中稳定同位素表现出明显的高程效应(拉市海除外),8¹⁸0随海拔升高而降低.非冰川区湖泊蒸发效应较为明显,同位素值明显偏正.

关键词:氢氧稳定同位素;过量氘;空间分布;拉市海;玉龙雪山

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Spatial Distribution of Stable Isotope from the Lakes in Typical Temperate Glacier Region

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Abstract: We focused mainly on the spatial variation and influencing factors of hydrogen and oxygen stable isotopes between water samples collected at the surface and different depths in the Lashi Lake in August, 2014. Hydrological supply characteristics of the lake in typical temperate glacier region were discussed. The results showed that the values of $\delta^{18}O$ and δD in the Lashi Lake ranged from -12.98% to -8.16% with the mean of -9.75% and from -99.42% to -73.78% with the mean of -82.23%, respectively. There was a reversed spatial variation between $\delta^{18}O$ and d. Relatively low values of $\delta^{18}O$ with high values of d were found at the edge of the lake where the rivers drained into. Meanwhile, the values of d in the vertical profile varied little with depth, suggesting that the waters mixed sufficiently in the vertical direction. The d values increased at first and then decreased from east to west at different layers, but both increase and decrease exhibited different velocities, which were related to the river distribution, the locality of the lake and environmental conditions etc. River water and atmospheric precipitation were the main recharge sources of the Lashi Lake, and the melt-water of snow and ice might also be the supply resource. The $\delta^{18}O$ values of lake water in glacier region decreased along the elevation (except for Lashi Lake), generally, this phenomenon was called "altitude effect". Moreover, high isotopic values of the lake water from non-glacier region were due to the evaporation effect.

Key words: hydrogen and oxygen stable isotope; deuterium excess; spatial variation; Lashi Lake; Mt. Yulong

氢氧同位素作为自然界水体的主要组成部分,对于研究区域水体的水文平衡和水文循环具有重要意义[1~3].湖泊水体8¹⁸O主要反映了流域的降水同位素特征、湖泊蒸发强度、流域水文状况及湖水的滞留时间等.因此,利用稳定同位素示踪湖泊水体的演化过程,为解决湖泊水量、水质和水文循环过程提供一种重要的方法.目前国外关于同位素技术在湖泊水文学中的应用研究较多:Gibson等^[4]利用稳定同位素模型模拟了高纬度地区湖泊的蒸发过程;Biggs等^[5]运用同位素质量平衡方法模拟

高海拔区湖水的蒸发和流入比率,研究了湖泊的现代水文过程.就国内研究而言,氢氧稳定同位素质量平衡方法已经被广泛运用于评估湖泊的水量平衡、长期和短期蒸发过程等方面:吴敬禄等[6]通

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作者简介: 史晓宜(1990~),女,硕士研究生,主要研究方向为水文 与水资源,E-mail:xyshi14@lzu.edu.cn 过对长江中下游 45 个湖泊不同季节水体δ¹⁸O分析,研究了湖泊水体的演化过程; Tian 等^[7]、巩同梁等^[8]对羊卓雍错流域稳定同位素水文循环进行了相关研究; 高晶等^[9]、臧娅琳等^[10]通过深入研究青藏高原典型湖泊区多种水体的稳定同位素时空分布特征,揭示了高原湖泊的水文过程; Bao等^[11]利用稳定同位素物质平衡方法估算了青海湖的蒸发量,发现计算结果与同时期实测蒸发量数据相当.

云贵高原是我国淡水湖泊分布较多的地区之 一,主要分布在滇中和滇西北地区,该区域湖泊一般 海拔较高,湖盆较深,且均在石灰岩、砂岩地区,故 其湖泊特色与长江中下游浅水湖泊迥异,与分布我 国西北、东北者更不相同. 丽江-玉龙雪山地区是我 国典型的温冰川分布区,区内分布有数十个高原湖 泊,如拉市海、程海、沪沽湖和九子海溶蚀高原洼 地雨季形成的小湖泊等,不仅是重要的区域水资源 和旅游资源,而且在区域水循环和环境研究中具有 特殊的意义. 玉龙雪山冰川是该区域重要的水资源 之一,近年来,丽江地区气温上升,玉龙雪山冰川发 生显著变化,冰川面积不断缩减的同时,冰川末端海 拔明显上升. 气候变化和玉龙雪山地区冰川的快速 消融和退缩,势必对湖泊蓄水量和水文平衡等造成 影响,因此,需开展湖泊在区域水资源和水循环过程 中的重要性研究.

玉龙雪山地区水体同位素的研究较多, He 等^[12~14]对玉龙雪山冰川区积雪及其融水径流中稳定同位素进行了一系列研究,揭示了冰川区水体δ¹⁸O的时空变化及其气候效应: 庞洪喜等^[15,16]就冬

夏季玉龙雪山冰川不同水体稳定同位素的分馏开展了对比研究,并对丽江市大气降水的影响因子进行了深入探究; Pu 等[17,18]利用同位素示踪了丽江-玉龙雪山地区的水文过程. 然而,该地区尚无湖水稳定同位素时空变化方面的研究和详细报道,因此,有必要展开对该区域湖泊的同位素组成及水文补给、蒸发等水文过程的研究.

本文选择丽江-玉龙雪山地区典型湖泊拉市海为研究对象,分析湖泊水体中8¹⁸0以及过量氘的空间变化特征及影响因素,探讨典型温冰川区湖泊的水文补给状况,通过揭示典型温冰川区湖泊的现代水文循环特征,以期为该区域湖相沉积中稳定同位素记录研究提供参考.

1 材料与方法

1.1 研究区域概况

玉龙雪山(27°10′~27°40′N,100°9′~100°20′E)位于青藏高原东南缘、横断山脉的南端,是我国典型的温冰川分布区. 其冰川分属于四个流域,即大具沟流域、白水河流域、漾弓江流域和仁河流域. 雪山东麓从北向南依次为大具盆地、甘海子盆地和丽江盆地. 该地区受到季风及多种环流系统的影响,表现出干湿季分明、降水主要集中在季风期的南亚季风气候特征.

拉市海位于丽江盆地西部 8 km 处,一个封闭型的季节性宽浅湖泊(如图 1),湖面海拔2 437 m,水深 2.5~4.5 m,蓄水量达4 150万 m^{3[19]}. 湖泊受到位于北侧的美泉河及南侧的清水河等的补给,湖水由指云寺暗河人口排泄,与金沙江支流相连^[20].

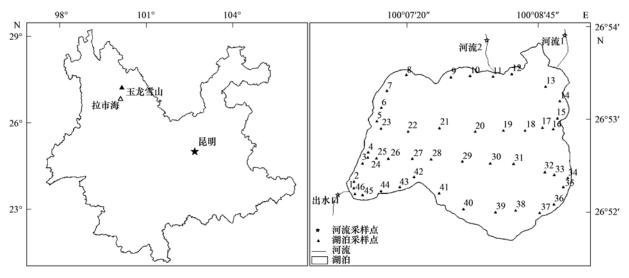


图 1 拉市海湖水采样点分布示意

Fig. 1 Location of research area and sampling sites

1.2 样品采集与分析

2014年8月在拉市海进行水样采集,采集路线如图1所示.首先沿湖周边进行表层水样采集,样品数共计46个;其次沿两条平行断面(偏北断面:标号16~23;偏南断面:标号25~33)在湖中进行采样,在每个采样点自表层向下以1m为间隔采集水样,样品数共计56个.同期,采集两条入湖河流(河水1和河水2)水样,具体位置如图1所示.

取样过程中,用采集的水样清洗聚乙烯样品瓶3次后再装样,并在每个样品瓶上编号,用 GPS 记录采样点地理位置,同时记录详细信息. 样品密封后放入玉龙雪山冰川与环境观测研究站内 - 15℃的冰箱低温保存. 之后以冷冻方式运回中国科学院寒区旱区环境与工程研究所冰冻圈国家重点实验室. 利用液态水同位素分析仪 Picarro L2130-i 进行测定,测量结果以相对于维也纳标准平均海水(Vienna Standard Mean Oceanic Water, V-SMOW)千分差的形式表示,δ¹⁸O和 δD 的测量误差分别为±0.025‰和±0.1‰.

2 结果与分析

2.1 拉市海表层湖水氢氧同位素组成及关系

同位素测试结果表明,表层湖水的 $\delta^{18}O$ 值在 -12.98‰~-8.16‰~之间 波 动,平均 值 为 <math>-9.75‰; δD 值在 -99.42‰~-73.78‰之间波 动,平均值为 <math>-82.23‰. 根据标准偏差,湖水的 $\delta^{18}O$ 值变化幅度较小,标准差为 1.36, δD 值变化明显,波动幅度为 25.64‰. 入湖河水的 $\delta^{18}O$ 值分别为 -13.90‰(河水 1)和 <math>-13.59‰(河水 2). 对比发现,湖水的 $\delta^{18}O$ 值普遍高于入湖河水及降水(平均值为 $-14.57‰)^{[21]}$ 的 $\delta^{18}O$ 值,这主要与湖水的蒸发有关.

研究表明,全球降水中氢和氧同位素存在一种线性关系. Craig^[22]把这种关系定义为全球大气降水线(global meteoric water line, GMWL),表达式为: $\delta D = 8 \, \delta^{18}O + 10$. 对拉市海表层 46 个湖水样品分析表明, δD 和 $\delta^{18}O$ 的关系为: $\delta D = 5.53 \, \delta^{18}O - 28.36 (R^2 = 0.99)$,其斜率高于全球大气降水线 $8(\delta D = 8 \, \delta^{18}O + 10)$. Zeng 等^[23]根据在丽江采集的降水资料,得出当地降水线为: $\delta D = 7.59 \, \delta^{18}O + 1.35 \, (R^2 = 0.99)$. 入湖河水的 $\delta^{18}O$ 值分别为 -13.90%0和 -13.59%0,明显低于降水及湖水的同位素值. 由图 2 可知,拉市海湖水采样点均位于我国大气降水线及当地降水线的下方,说明湖水经历了明显的蒸发.

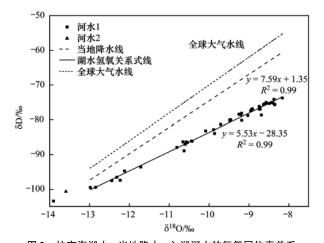


图 2 拉市海湖水、当地降水、入湖河水的氢氧同位素关系
 Fig. 2 The δ¹⁸O and δD values of lake water, rainwater,
 and river water in the study area

2.2 拉市海表层湖水δ¹⁸O和过量氘的空间变化

利用 Arcgis 对拉市海表层湖水δ¹⁸O进行插值,得出δ¹⁸O的空间分布(如图 3). 由图 3 可知,表层湖水δ¹⁸O存在一定的空间变化,有河水注入的区域为低值区,这是由于拉市海的补给河流主要分布在东北岸,入湖河水的δ¹⁸O值分别为 - 13.90‰和-13.59‰,明显低于湖水δ¹⁸O的平均值,故在河水汇入处受到稀释作用的影响,导致该区域δ¹⁸O值较低. 而湖泊东西岸的δ¹⁸O值相对较高,主要是由于河水缓慢汇入湖泊的过程中,与湖泊边缘(北岸)的水体相互混合,使其同位素值显著降低,同等蒸发的条件下,无河水注入的区域δ¹⁸O值相对较高.

水在蒸发过程中的动力分馏作用使氢和氧稳定同位素的平衡分馏被破坏,在降水中 δD 和 $\delta^{18}O$ 之间的关系出现一个差值,Dansgaard ^[24]把它定义为过量 $fi.d = \delta D - 8 \delta^{18}O$. 另外,d 值不仅仅反映水体非平衡程度,而且能反映蒸发速率的快慢. 由表 1 所示,过量氘值在 $-9.14\%c \sim 4.43\%c$ 之间波动,平均值为 -4.23%c. 而全球降水中过量氘值为 10%c,说明拉市海湖水的蒸发效应比较显著. 以克里金插值法得到过量氘的空间变化图(如图 4). 由图 3、4 可知, $\delta^{18}O$ 与过量氘表现出相反的空间分布特征.

2.3 拉市海深水剖面过量氘的空间变化

图 5 为 8 个深水剖面中过量氘随深度变化.本研究选取 8 个深水剖面(采样点序号分别为偏北断面中 19、20、21、22;偏南断面中 28、29、30、31) 共 40 个水样作出过量氘随水深变化的曲线.从中可知,8 个剖面的过量氘值随深度变化较小,偏北断面中变化幅度最大的是 22 号采样点,变化范围为-5.69‰~-4.23‰,相差 1.45‰;偏南断面中过

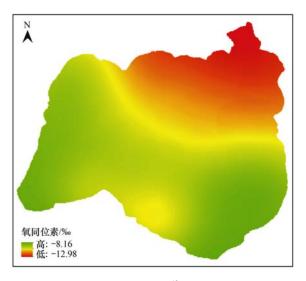


图 3 拉市海表层水样 8¹⁸O的空间变化

Fig. 3 Spatial variation of surface water δ¹⁸O values in Lashi Lake

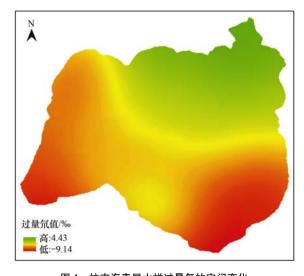
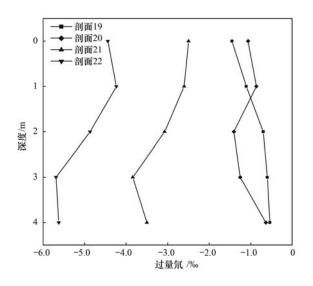


图 4 拉市海表层水样过量氘的空间变化

Fig. 4 Spatial variation of surface water deuterium excess in Lashi Lake



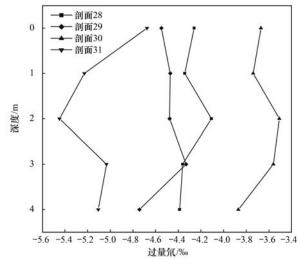


图 5 拉市海湖水过量氘随深度变化

Fig. 5 Variations of deuterium excess values with depth for the lake water

量氘值随深度变化相差最大的是 31 号采样点,最大值为 -4.68‰,最小值为 -5.45‰,相差 0.77‰,表明湖水在垂直方向上混合比较充分.由图 6 可知,偏北断面和偏南断面的采样点在不同深度层上自东向西表现出先增大后减小的变化趋势,但增减速率表现出相反特性,这可能与入湖河流的分布有关^[9].

3 讨论

3.1 拉市海水文补给特征

通常情况下,湖水的补给主要来自于大气降水、河水、地下水等,位于冰川区的湖泊,还可能受到冰雪融水的补给.本研究区水资源包括地表水和地下

水,主要赋存于冰川、积雪、河流、湖泊和地下岩层中. 该区域夏季受西南季风的影响,降水较为充沛,降水量约为800 mm [25,26],大量的降水为拉市海提供了重要的水源补给. 根据实地考察和资料记载 [20],拉市海受到入湖河流美泉河及清水河等的补给. 由表 1 可知,湖水的 δ^{18} O值 -9.75%,人湖河水(河水 1、河水 2)的 δ^{18} O值 分别为 -13.90%。-13.59‰,湖水的 δ^{18} O值明显高于河水的 δ^{18} O值,表明湖水的稳定同位素组成受到蒸发效应的影响. 此外,本研究选取白水河流域的泉水双胞泉和丽江最大的泉眼群黑龙潭作为地下水代表,其 δ^{18} O值分别为 -14.50‰和 -14.72‰(如表 1),与人湖河水的 δ^{18} O值相比,差值不大,在 -0.60‰ $\sim -1.13‰$ 之间

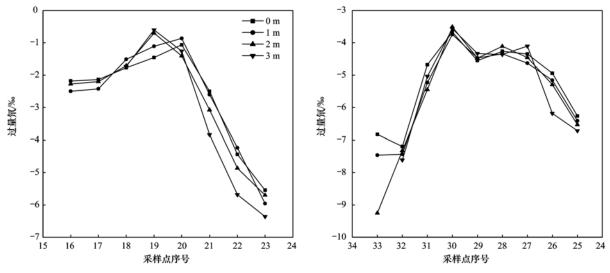


图 6 拉市海不同深度过量氘自东向西变化曲线

Fig. 6 Variations of deuterium excess in different depth from east to west in Lashi Lake

表 1 玉龙雪山周边不同水体 δ^{18} O 的平均值¹⁾

Table 1 Average δ¹⁸O values of different water bodies in Mt. Yulong region

采样类型	采样时间 —	$\delta^{18} O/\% o$			 文献
木件矢型		最小值	最大值	平均值	文
河水 1	2004年9月	_	_	- 13. 90	本研究
河水 2	2004年9月	_	_	- 13. 59	本研究
湖水	2014年8月	- 12. 98	-8.16	-9.75	本研究
双胞泉	2009年5~10月	- 15. 06	- 14. 17	- 14. 50	[17]
黑龙潭泉水	2009年5~10月	- 15. 13	-14.21	- 14. 72	[17]
冰雪融水	2009年5~9月	-13.41	-11.65	- 12. 59	[17]
丽江市降水	2008年6~9月	_	_	- 14. 57	[21]

1)"一"表示未测

波动. 已有研究表明^[27],由于该区域特殊的地貌条件,地表水与地下水相互补给的速度较快,转换过程中8¹⁸0含量变化较小. 因此,地下水也可能补给拉市海.

位于丽江盆地北部的玉龙雪山是盆地内水系的主要发源地,由于冰雪融水和丰沛的降水补给,玉龙雪山周边发育有大量的小河流,总体上由北向南注入丽江盆地^[28,29].因此,冰雪融水可能通过补给地下水及河水,间接补给拉市海. 范弢^[20]通过遥感资料指出玉龙雪山南侧的融水径流,首先进入文海湖,随后于湖南部附近岩溶地带的落水洞泄通过地下暗河车门举大泉补给拉市海,说明拉市海可能受到冰雪融水的补给.

在全球气候变暖的背景下,冰川变化对以冰雪融水为主要补给的高原湖泊变化有重要的影响^[30].研究表明,玉龙雪山地区冰川持续退缩,1957~2009年,消失冰川共计6条,现仅存13条,其面积为4.42 km²,相比1957年总面积减少了7.18 km²;冰

川表面形态破碎严重,冰川积累区已出现冰面河和冰面湖等^[31,32].同时,玉龙雪山作为丽江盆地内重要的水源发源地,冰川面积的明显减小,势必会引起本区水资源的动态变化,这可能对拉市海湖泊的水量及水文补给状况产生一定影响.

3.2 拉市海湖水与其他地区湖水的δ¹⁸O值比较

通过对湖泊进行稳定同位素研究,一方面可以揭示湖泊的现代水文循环过程,另一方面也为湖相沉积中稳定同位素记录提供参考. 表 2 综合了本研究及夏季期间其他地区不同类型湖泊水体的8¹⁸O平均值. 臧娅琳等^[10]对羊卓雍错湖泊表层8¹⁸O值及过量氘的空间变化进行研究,对比发现,羊卓雍错湖和拉市海表现出一致的同位素分布特征:有河水注入的边缘区8¹⁸O值较低,表层湖水的8¹⁸O值空间变化幅度不大,但羊卓雍错湖中心区8¹⁸O值较高,拉市海湖水的高值区主要分布在无河水注入的边缘区. 这种空间分布的差异,可能受到湖泊形状、地理位置及入湖河流的分布等的影

[9,10]

对青藏高原区典型高原湖泊(如羊卓雍错、纳木错和普莫雍错)的氧同位素组成进行对比研究,发现青藏高原区湖泊的δ¹⁸O 值随着海拔高度的增加而减小,呈现明显的"高程效应". 但拉市海的δ¹⁸O值并不符合高程效应,可能是因为该区域在夏季受

季风活动影响较大,降水中重同位素值较贫化,导致 拉市海水体的δ¹⁸O值相对较低.同时,对于非冰川区 湖泊(巴丹吉林沙漠湖泊和青海湖),水体的δ¹⁸O值 明显较高,蒸发作用影响极为显著,这些差异表明湖 水的同位素组成主要与其补给类型和地理位置、自 然条件等因素密切相关^[35].

表 2 拉市海湖水与其他地区湖水的 $\delta^{18}O$ 值比较

Table 2 Comparison of δ¹⁸O values for different lakes

湖泊	经度	纬度	采样日期	样品数/个	$\delta^{18}\!O/\%\!\sigma$	海拔/m	文献
拉市海	100°06′ ~100°09′	26°51′ ~ 26°53′	2014年8月	46	-9.75	2 437	本研究
羊卓雍错	90°08′ ~91°45′	28°27′ ~29°12′	2006年8月	30	-5.4	4 440	[9]
纳木错	90°16′ ~91°03′	30°30′ ~ 30°55′	2005~2008年的6~9月	_	-6.7	4 718	[33]
普莫雍错	90°13′ ~90°33′	28°29′ ~28°38′	2006年8~9月	21	-7.2	5 030	[9]
巴丹吉林沙漠湖泊	99°23′ ~104°34′	39°04′ ~42°12′	2012年8~9月	38	5. 9	_	[34]
青海湖	97°50′ ~ 101°20	36°15′ ~38°20′	_	_	2.0	4 500	[11]

4 结论

- (1)拉市海表层湖水的 $\delta^{18}O$ 、 δD 值分别在 $-12.98‰ \sim -8.16‰和 -99.42‰ \sim -73.78‰之间波动,平均值分别为 <math>-9.75‰和 -82.23‰,\delta D$ 和 $\delta^{18}O$ 的关系为: $\delta D = 5.53$ $\delta^{18}O 28.36$,其斜率高于地降水线斜率 7.59,表明湖水经历了明显的蒸发效应.
- (2)拉市海表层湖水的δ¹δO及过量氘值的空间变化相反,有河水注入的区域,湖水的δ¹δ0值较低而过量氘值较高,无河水补给的东西岸δ¹δO值较高而过量氘较低.
- (3)垂直剖面上湖水的过量氘值随深度变化较小,表明拉市海湖水在垂直方向上充分混合;过量氘值在不同深度层上表现出自东向西先增大后减小的变化趋势,但增减速率呈现出相反特性. 蒸发效应和不同补给源的同位素组成是影响湖水同位素空间分布的主要因素.
- (4)青藏高原冰川区湖泊水体中8¹⁸0值随海拔升高而减小,表现出明显的高程效应,因拉市海受季风影响较大,湖水同位素值并不符合高程效应.非冰川区湖泊蒸发效应较为显著,同位素值明显偏正. 参考文献:
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