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2000~2021年黄土高原生态分区 NEP 时空变化及其 驱动因子

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摘要:净生态系统生产力(NEP)是陆地生态系统碳源/汇定量评价的重要指标.以黄土高原地区及6个生态分区(黄土高塬沟壑 区A1、A2副区,黄土丘陵沟壑区B1、B2副区、沙地和农灌区(C区)和土石山区及河谷平原区(D区))为研究对象,结合遥感、 气象、地形和人类活动数据,采用相关性分析、多元回归残差分析和地理探测器等方法,估算区域NEP并分析其时空变化特 征及气候、地形、人为因子对NEP时空变化的影响.结果表明,在时间尺度上,2000~2021年,黄土高原NEP多年平均值(以 C计)为104.62g·(m²·a)⁻¹.黄土高原及各生态分区NEP均呈增长趋势,其中,黄土高塬沟壑区A2副区NEP年均增长率最大, 为9.04g·(m²·a)⁻¹;沙地和农灌区NEP年均增长率最小,为2.74g·(m²·a)⁻¹.除沙地和农灌区为弱碳源外,其余各生态分区均表 现为碳汇.在空间尺度上,黄土高原年均NEP呈现东南高西北低的分布格局,碳汇高值主要分布在黄土高塬沟壑区南部,碳 源区主要分布在黄土高塬沟壑区北部、沙地和农灌区的大部;NEP的空间变化有显著差异,高增幅主要分布在A2副区中南部 以及B2副区的西南部.黄土高原及各生态分区NEP时间变化受人为因素影响最大,人类活动数据与NEP的相关系数均大于 0.80, 且人为因素对NEP的贡献率均在50%以上;NEP的空间变化受气象因子的影响较大,降水、太阳辐射是影响空间变化 的主导因子.总之,黄土高原NEP的时空变化受自然因素和人类社会因素共同影响.研究结果可为黄土高原陆地生态系统减排 增汇及实现双碳目标提供参考.

关键词:黄土高原; 生态分区; 净生态系统生产力; 相关分析; 残差分析; 地理探测器 中图分类号: X171.1 文献标识码: A 文章编号: 0250-3301(2024)05-2806-11 DOI: 10.13227/j. hjkx. 202306059

Spatio-temporal Variation in NEP in Ecological Zoning on the Loess Plateau and Its Driving Factors from 2000 to 2021

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Abstract: Net ecosystem productivity (NEP) is an important index for the quantitative evaluation of carbon sources and sinks in terrestrial ecosystems. Based on MOD17A3 and meteorological data, the vegetation NEP was estimated from 2000 to 2021 in the Loess Plateau (LP) and its six ecological subregions of the LP (loess sorghum gully subregions ; A1, A2; loess hilly and gully subregions: B1, B2; sandy land and agricultural irrigation subregion: C; and earth-rock mountain and river valley plain subregion: D). Combined with the terrain, remote sensing, and human activity data, Theil-Sen Median trend analysis, correlation analysis, multiple regression residual analysis, and geographic detector were used, respectively, to explore the spatio-temporal characteristics of NEP and its response mechanism to climate, terrain, and human activity. The results showed that: ① On the temporal scale, from 2000 to 2021 the annual mean NEP of the LP region (in terms of C) was 104.62 g· (m²·a)⁻¹. The annual mean NEP for both the whole LP and each of the ecological subregions showed a significant increase trend, and the NEP of the LP increased by 6. 10 g • (m² • a)⁻¹ during the study period. The highest growth rate of the NEP was 9.04 g • (m² • a) -1, occurring in the A2 subregion of the loess sorghum gully subregions. The subregion C had the lowest growth rate of 2.74 g · (m²·a)⁻¹. Except for the C subregion, all other ecological subregions (A1, A2, B1, B2, and D) were carbon sinks. 2 On the spatial scale, the spatial distribution of annual NEP on the LP was significantly different, with the higher NEP distribution in the southeast of the LP and the lower in the northwest of the LP. The high carbon sink area was mainly distributed in the southern part of the loess sorghum gully subregions, and the carbon source area was mainly distributed in the northern part of the loess sorghum gully subregions and most of the C subregion. The high growth rate was mainly distributed in the central and the southern part of the A2 subregion and the southwest part of the B2 subregion. ③ Human activities had the greatest influence on the temporal variation in NEP in the LP and all the ecological subregions, with the correlation coefficient between human activity data and NEP being above 0.80, and the relative contribution rates of human factors was greater than 50%. The spatial distribution was greatly affected by meteorological factors, among which the precipitation and solar radiation were the main factors affecting the spatial changes in the NEP of the LP. The temporal and spatial variations in the NEP in the LP were influenced by natural and human social factors. To some extent, these results can provide a reference for the terrestrial ecosystem in the LP to reduce emissions and increase sinks and to achieve the goal of double carbon.

Key words: Loess Plateau(LP); ecological subregion; net ecosystem productivity; correlation analysis; residual analysis; geographic detector

工业革命以来, 化石燃料的燃烧和土地利用的 变化, 使大气中温室气体特别是二氧化碳(CO₂)浓 度激增^[1], 从而导致全球平均温度持续上升. 陆地生 态系统是地球上重要的碳汇, 每年抵消超过 30% 的 人为碳排放量^[2]. 净生态系统生产力(NEP)是陆地生 态系统与大气之间的净碳交换, 是定量评价陆地生 态系统碳源/汇的重要指标[3]. 在气候变化和人为干

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扰的背景下,准确掌握陆地碳源/汇的时空变化并揭 示其驱动机制对区域可持续发展和全球碳减排具有 重要意义.

中国以其约占全球6.5%的陆地面积,贡献了 相当于全球陆地生态系统净 CO2 吸收量的 10%~ 31%, 表明中国陆地生态系统在全球生态系统碳汇 中具有重要作用^[4]. 自 2000 年以来,中国植被 NEP 整体呈增加趋势^[5],然而不同区域NEP的年际变化 趋势并不相同. 2000~2020年西南地区^[6]和 2000~ 2015年青海高原^[7]的NEP呈显著地上升趋势,而 2000~2020年海南岛森林 NEP 呈现不显著地下降趋 势^[8]. 区域 NEP 的时空变化是多因素共同作用的结 果[9~11]. 气候变化可以驱动植被动态变化和调节土 壤呼吸,从而影响陆地生态系统的碳循环[6-8].一 般而言,降雨充沛和太阳辐射充足,加之较高的气 温,有利于促进植物的光合作用,进而增强植被的 碳汇能力[5].人类活动(植树造林、生态修复和森林 经营等)是驱动区域尺度生态系统碳源/汇时空变化 的重要因素[12,13].据报道,退耕还林工程实施后黄 土高原净初级生产力 (net primary productivity, NPP)、NEP及碳储量均呈增长趋势^[14]; 生态保护 工程实施后三江源地区 NEP 呈显著上升趋势^[15], 碳汇功能显著增强.不同区域气候变化和人类活动 对 NEP 的贡献率有很大差异^[6]. 综上所述, 我国不 同区域植被 NEP 变化趋势及对不同驱动因子的响 应,均存在显著的时空差异,因此,有必要针对区 域尺度,结合时空动态演变分析来揭示区域 NEP 变化的驱动机制.

黄土高原地域广阔, 气候类型多样, 自然地理 条件复杂和空间组合变化明显,降水和气温空间分 布很不平衡,植被覆盖存在很大的空间异质性[16], 是世界上水土流失最严重和生态环境最脆弱的地区 之一[14,17].近20年来,随着退耕还林/草、天然林保 护和山水林田湖草等工程的实施,该区的植被覆盖 度得到极大改善, 植被的固碳能力显著提高[18]. 对 生态系统碳源/汇的研究多集中于黄土高原整个区域 陆地生态系统的NPP、碳储量的研究^[19~21],而对不 同生态分区陆地生态系统碳源/汇的时空变化规律及 环境因子驱动机制仍然相对缺乏.为此,本文依据 杨艳芬等[16]的黄土高原生态分区结果,结合遥感、 气象、地形和人类活动数据, 探析 2000~2021 年黄 土高原及其6个生态分区 NEP 的时空分布特征、变 化趋势及其对不同驱动因子的响应,量化自然和人 类活动因子对 NEP 变化的贡献, 以期为黄土高原减 排增汇、实现双碳目标提供理论参考.

1 材料与方法

1.1 研究区概况

黄土高原(33°41′~41°16′N, 100°52′~ 114°33′E)位于中国中部偏北,总面积达6.4×10° km². 属大陆性季风气候, 全年≥10℃的积温 2 300~ 4 500 ℃, 无霜期 120~250 d, 日照时数 1 900~3 200 h^[22], 年平均气温 3.6~14.3 ℃, 年降水量 150~800 mm^[23],降水自东南向西北、由山地向平地递减,年 际变异较大,年内分布也不均匀[16].地势西南高、 东南低,海拔85~3700m.地貌类型多样,由丘陵、 高塬、平原、沙漠、草原和土石山地等组成,其中 山区、丘陵区和高塬区占2/3以上.土壤类型主要有 黄绵土、褐土、黑垆土等. 植被具有明显的分区差 异,从东南向西北依次为森林草原、典型草原和半 荒漠草原.杨艳芬等^[16]依据国家发改委的分区原则 和方法,综合考虑了黄土高原水土流失治理技术、 模式和生态恢复建设工程的区域性差异,将黄土高 原划分为4个生态分区:黄土高塬沟壑区(A)、黄土 丘陵沟壑区(B)、沙地和农灌区(C)和土石山区及河 谷平原区(D).其中黄土高塬沟壑区以六盘山为界, 划分为 A1 和 A2 两个副区; 黄土丘陵沟壑区以毛 乌素沙漠南缘为界,划分为 B1 和 B2 两个副区 (图1).



A:黄土高塬沟壑区,A1:黄土高塬沟壑区A1副区,A2:黄土高塬沟 壑区A2副区;B:黄土丘陵沟壑区,B1:黄土丘陵沟壑区B1副区,B2: 黄土丘陵沟壑区B2副区;C:沙地和农灌区;D:土石山区及河谷平 原区

图1 研究区位置及高程示意

Fig. 1 Location and elevation of the study area

1.2 数据来源

本研究使用了由美国国家航空航天局数据中心 (http://ladsweb.modaps.nasa.gov/)提供的 MODIS Terra NPP数据(MOD17A3HGF)用于 NEP的估算,空 间分辨率为 500 m,时间分辨率为 1 a,时间跨度是 2000~2021年.

数字高程模型数据来源于由美国国家航空航天

局(NASA)和美国国家测绘局(NIMA)联合测量的 SRTM产品,空间分辨率90m.坡度(slope)和坡向 (aspect)数据基于高程(ele)数据计算得到.

气温和降水量数据来源于国家地球系统科学数据中心——黄土高原分中心(http://loess.geodata.cn),时间分辨率为月度,空间分辨率为1km^[24],月气温数据求平均值获得年平均气温(T),月降水量数据求和得到年降水量(P),时间跨度均为2000~2021年;年太阳总辐射(SR)数据来自中国气象数据网(http://data.cma.cn/),通过对逐日日照时数数据进行计算得到年太阳总辐射数据,并利用ArcGIS10.2软件采用反距离权重插值法(IDW)将站点数据插值成栅格数据^[25],时间跨度为2000~2021年.

人类活动强度(HA)数据来自 Mu 等^[26]的研究, 该数据从多个来源收集 8类人类压力变量(即建筑环 境、人口密度、夜间灯光、农田、牧场、公路、铁 路和通航航道),生成年度全球人类足迹数据集来 综合表示人类活动强度,空间分辨率为1 km,时间 跨度为2000~2018年.

将所有数据统一投影至WGS1984/UTM 49N坐标 系,空间分辨率重采样至1km,便于进行后续数据 的处理与分析,

1.3 研究方法

1.3.1 NEP估算

NEP 定义为 NPP 与土壤异养呼吸(RH)的差值, 其计算公式如下:

NEP = NPP - RH

如果 NEP 大于零,表示植被固定的碳比土壤呼 吸释放的碳大,即碳汇;否则就是碳源.

根据前人的研究^[27-29],笔者对黄土高原地区土 壤异养呼吸的估算模型进行了对比,最终采用 Pei 等^[27]建立的经验公式,该公式已被多个研究成功运 用^[30,31].公式如下:

 $RH(x,t) = 0.22 \times [\exp\{0.091 \ 3T(x,t)\} +$

 $\ln\{0.3145P(x,t)+1\}$]×30×46.5%(2) 式中, *T*和*P*分别为像元*x*在*t*月的平均温度(°C)和 总降水量(mm).

将每年的月度RH数据相加得到年RH,后用年 NPP减去RH得到年NEP.

1.3.2 NEP变化趋势分析

采用 Theil-Sen 中值趋势分析方法来探讨黄土高 原 NEP 的变化趋势,其计算公式如下:

$$\beta = \text{Median}\left(\frac{\text{NEP}_j - \text{NEP}_i}{j - i}\right)$$
(3)

式中, β 为NEP变化趋势;NEP_i和NEP_i分别表示第*i* 年和第*i*年NEP的时间序列值(本研究中2000 $\leq i <$ *j*≤2021).*β*>0表示研究期内NEP变化呈增加趋势, β<0表示研究期内NEP变化呈减少趋势.

采用 Mann-Kendall 方法对 NEP 变化趋势进行 显著性检验.对于标准值 Z 大于 0,则序列呈上升 趋势;若小于 0,则序列呈下降趋势.当 Z 的绝对 值大于等于 1.64、1.96 和 2.58 时则说明该时间序列 分别通过了置信水平 90%、95% 和 99% 的显著性 检验.

1.3.3 皮尔逊相关分析

利用相关性系数和显著性检验来揭示各驱动因 子与NEP时空变化相关性的强弱.

1.3.4 多元回归残差分析

利用多元回归残差分析分离气候变化和人类活动对 NEP 变化的贡献. 通过建立年均气温、年降水量和年太阳总辐射与 NEP 的回归方程预测气候变化对 NEP 的影响(NEP_{cc}); NEP 观测值(NEP_{obs})与 NEP_{cc}之间的差值,即 NEP_{HA},表示人类活动对 NEP 的影响.

$$NEP_{cc} = a \times T + b \times P + c \times SR + d \qquad (4)$$
$$NEP_{HA} = NEP_{obs} - NEP_{cc} \qquad (5)$$

式中,NEP_{cc}和NEP_{abs}分别表示基于回归模型的NEP 预测值和NEP 观测值; $a, b, c \pi d$ 为模型参数; $T, P \pi SR$ 分别为年均气温(°C)、年降水量(mm)和年太 阳总辐射(MJ·m⁻²);NEP_{HA}为残差.

利用一元线性回归法计算 2000~2021 年的 NEP_{cc} 和 NEP_{HA}的变化趋势率,分别表示在气候变化和人 类活动影响下的 NEP 变化趋势.趋势率为正表示气 候变化或人类活动可促进 NEP 的增加;反之,表示 会导致 NEP 下降.根据表 1 对 NEP 变化的主要驱动 因素进行区分,计算气候变化和人类活动对 NEP 变 化的贡献率.

1.3.5 地理探测器模型

地理探测器是一种揭示空间异质性、基于4个 模块量化驱动因素对响应变量影响^[32]的空间统计方 法.本研究中主要使用因子探测和交互探测来分析 黄土高原 NEP空间差异的驱动机制.因子探测主要 是衡量驱动因素对响应变量空间变异的解释能力, 用q值来表示.

$$q = 1 - \frac{\text{SSW}}{\text{SST}} = 1 - \frac{\sum_{h=1}^{L} N_h \sigma_h^2}{N \sigma^2}$$
(6)

式中,SSW和SST分别表示层内方差之和和全区总 方差; $h = 1, 2, 3, \dots, L$ 为依赖或独立因素的分层; N_h 和N分别为层h和全区的单元数; σ_h^2 和 σ^2 分别为 层h和全区NEP值的方差; $0 \le q \le 1$.

用交互探测来判断两个驱动因素的解释能力是

表1 NEP变化的驱动因素判定标准及贡献率计算方法¹⁾

	Table 1 Criteria for	r determining the driv	ving factors of NEP c	hanges and calculation methods for	contribution rate	
$\mathrm{slope}(\mathrm{NEP}_{_{\mathrm{obs}}})$	驱动因素	驱动因素的划分标准		驱动因素的贡献率/%		
		$slope(NEP_{cc})$	$\mathrm{slope}(\mathrm{NEP}_{\mathrm{HA}})$	气候变化	人类活动	
	СС 和 НА	>0	>0	$slope(NEP_{CC})/slope(NEP_{obs})$	$\mathrm{slope}(\mathrm{NEP}_{\mathrm{HA}})\!/\mathrm{slope}(\mathrm{NEP}_{\mathrm{obs}})$	
>0	CC	>0	<0	100	0	
	HA	<0	>0	0	100	
	СС 和 НА	<0	<0	$slope(NEP_{cc})/slope(NEP_{obs})$	$\mathrm{slope}(\mathrm{NEP}_{\mathrm{HA}})\!\big/\mathrm{slope}(\mathrm{NEP}_{\mathrm{obs}})$	
<0	CC	<0	>0	100	0	
	HA	>0	<0	0	100	

1)NEP_{abs}为NEP观测值;NEP_{cc}为气候变化对NEP变化影响量;NEP_{HA}为人为因素对NEP变化影响量;CC为气候变化;HA为人为因素

增强、减弱或相互独立.方法是对两个独立因素 X1、 X2 的 q 值,以及 X1 和 X2 交互探测的 q 值进行比 较^[32].

基于之前的研究^[14,33,34],并考虑到具体的数据 可用性,分别选取了气象因子(*T、P、SR*)、地形因 子(ele、slope、aspect)和人为因子(HA)这7个潜在 因子进行相关性和地理探测器分析.其中地理探测 器在要求离散自变量的前提下,根据q统计值的最 大值原则,采用最优离散方法对连续自变量进行重 新分类,本研究基于R Studio软件中的"GD"包进 行分类.

2 结果与分析

2.1 黄土高原NEP年际变化

2000~2021年黄土高原年均NEP值(以C计,下同)2001年最低为7.42g·(m²·a)⁻¹,2018年最高为 171.81g·(m²·a)⁻¹,平均值为104.62g·(m²·a)⁻¹[图2 (a)];整体呈显著增长趋势(P<0.05),年均增长率



为 6.10 g·(m²·a)⁻¹; 22 年间, 黄土高原碳源面积 (NEP < 0)占比呈波动下降趋势, 其中 2001 年占比 最大. 而随着 NEP 的增加, 2021 年碳汇区域(NEP > 0)的面积百分比达到 77%, 与 2000 年相比增加了 27%[图 2(b)].

6个生态分区的年均 NEP 均呈显著增长趋势, 但增长速度不同(图3). A2 副区 NEP 年均值和年变化 率均居首位,分别为 186.84 g·(m²·a)⁻¹和 9.04 g·(m²·a)⁻¹,C区 NEP 年均值和年变化率均最小,分 别为-30.41 g·(m²·a)⁻¹和 2.74 g·(m²·a)⁻¹.就 NEP 年均 值而言,B区和C区均小于黄土高原整个区域的平 均水平;而除 A1副区和C区外,其余各生态分区变 化率均大于黄土高原的变化率.各生态分区最大值 主要集中在 2018 年和 2019 年,最小值除 A1副区外 均在 2001 年.其中,2018 年 A2 副区 NEP 值最大为 288.18 g·(m²·a)⁻¹,2001 年 C 区 NEP 值最小为-80.09 g·(m²·a)⁻¹.整体来看除 C 区为弱碳源,其余各分区 均为显著的碳汇.





2.2 黄土高原 NEP 空间格局及其变化

黄土高原 NEP 整体呈现东南高西北低的空间分 布格局(图4). NEP < 0的面积占总面积的 26.56%, 主要分布在 C 区大部分地区, A 区北部和 B 区西北 部也有分布; NEP > 0的面积占 73.44%, 其中 NEP 高值区[> 300 g·(m²·a)⁻¹]占 7.45%, 主要分布在 A2 副区西南部以及A1副区南部边缘地区.

NEP空间变化格局与其空间分布格局相似,呈现由东南向西北递减的趋势[图5(a)].2000~2021年,黄土高原固碳能力以增长为主,面积占比达98.55%,显著(P < 0.05)或极显著(P < 0.01)增加的面积占全区域的89.41%,其中,NEP的微增幅[2.5~



图 4 2000~2021 年黄土高原及其生态分区平均 NEP 空间分布 Fig. 4 Spatial distribution of annual mean NEP in the Loess Plateau and its ecological subregions from 2000 to 2021

5 g·(m²·a)⁻¹]、中增幅[5~10 g·(m²·a)⁻¹]和高增幅 [>10 g·(m²·a)⁻¹]面积占比分别为22.67%、47.28% 和13.12%,高增幅主要分布在A2副区中南部以及 B2副区的西南部(如陕西延安、宝鸡,甘肃庆阳和 平凉等地).负增长区域面积仅占1.45%,零星分布 在C区,D区以及A1副区等部分地区(图5).

2.3 驱动因子分析

2.3.1 驱动因子对黄土高原 NEP 年际变化的影响

黄土高原全区及各生态分区 NEP 的年际变化与 HA 均呈极显著正相关(*P* < 0.01),相关系数均在 0.80以上;除A2副区和D区外,全区及其余生态分 区的 NEP 与年降水量的年际变化相关性均达到了显 著(*P* < 0.05)或者极显著水平(*P* < 0.01);全区及各 生态分区 NEP 的年际变化与年均气温和年太阳总辐 射相关性均未达到显著水平(P>0.05)(图6). 残差分 析表明,气候因素和人类活动对NEP年际变化的贡 献率分别为37.33%和62.67%,表明人类活动是引起 NEP年际变化的主导因子(图7). 黄土高原约19.37% 的区域受气候因素的主导(贡献率>50%),主要分布 在C区北部、A1副区北部以及D区东北部[图7 (a)]. 人类活动对NEP变化的贡献率在79.81%的区 域占主导,其中贡献率大于75%的地区占区域面积 的1/4,集中分布在研究区中部和南部[图7(b)].在 6个分区中,人类活动对NEP增长贡献率最高,贡 献率均在50%以上,而气候因子贡献率最高,贡 献率均在50%以上,而气候因子贡献率最大 (76.35%),对C区贡献率最小(53.88%),相应C区 气候因素引起NEP增长的贡献率最大(46.12%),A2









2.3.2 驱动因子对黄土高原 NEP 空间分异性的影响

因子的独立效应方面,相关分析和因子探测均 表明 P(r = 0.72, q = 0.54)和 SR(r = -0.56, q = 0.37)是影响黄土高原 NEP 空间格局的主要驱动因子(图 8),其中 SR 和 T与 NEP 呈负相关,而 HA 和 aspect 对 NEP 的解释力相对较弱.各生态分区的主要影响 因子不同,A区、BI副区、B2副区、C区、D区分别是P、T、P和SR、SR、ele(图8).整体来看,气象因素是黄土高原及各分区NEP空间变化的主要影响因素,地形和人类活动的解释力相对较弱.

为了量化两因子的交互作用,进一步做了交互 探测分析,表2中列出了影响黄土高原全区和各生 态分区 NEP 空间变化的主要影响因素(P、T、SR、 ele)之间及其与其他影响因子(HA、slope、aspect)交 互作用的q值.结果表明,成对因子的q值大于单个 因子的q值或它们的总和,说明两因子交互作用后 对 NEP 空间变化的解释力更强.黄土高原 $P \cap T$ 和 $P \cap$ ele 交互作用的q值最高,均为0.72.A1副区、B1 副区和D 区 $P \cap T$ 的交互解释力最强,q值分别为 0.77、0.86和0.56,A2副区 SR $\cap T$ 交互解释力最强, q值为 0.89,B2 副 区 $P \cap$ ele 解释力最强,q值为 0.82,而C 区 $T \cap$ HA 交互作用q值最高,为0.60 (表2).

3 讨论

3.1 时间尺度 NEP 变化分析

以往研究表明,自1999年开始实施退耕还林/ 草工程以来,黄土高原地区植被覆盖度明显提 高^[35-37].刘国彬等^[35]报道了黄土高原2017年的林草



Fig. 7 Contribution rates of climate change and human activities to the increase in NEP in the Loess Plateau and its ecological subregions

覆盖率比1999年提高了33.6%,水土保持林草及封 禁治理面积达到24万 km²以上;王逸男等^[36]报道了 2000年至2020年黄土高原植被覆盖度由0.39增加到 0.61, 且 2017年后提升速度加快. 国家政策和相关 生态恢复措施等人为因素是引起黄土高原植被覆盖 度增加的主要原因, 2000~2020年人类活动对黄土 高原植被覆盖度变化的贡献率为65.22%,气候变化 为34.78%[37]. 植被恢复、退耕还林/草和森林经营等 措施能够通过植物光合作用将大气中的CO2固定在 植物和土壤中,从而提升了黄土高原生态系统的固 碳能力[38].张佑铭等[39]报道了1990~2015年黄土高原 植被的 NPP 与植被固碳总体呈增加趋势,年均 NPP 增速为 2.74 g·(m²·a)⁻¹;杨丹等^[19]发现 2000~2015年 黄土高原植被 NPP 的平均增速为 3.62 g·(m²·a)⁻¹.本 研究结果表明, 2000~2021年黄土高原全区年均 NEP 呈显著增长趋势(P<0.05), 增长速率为 6.10 g·(m²·a)⁻¹,人类活动是引起黄土高原NEP年际变化 的主要因素,贡献率为62.67%,这与Feng等^[14]的研 究结果一致.

6个生态分区 NEP 均呈增长趋势,但增长速率 有所差异.A 区整体表现为显著的碳汇区,碳汇增速 也较快,这与该区高植被覆盖率有关^[40];而A2 副区 NEP 增长速率最快,这是因为A2 副区温度、降水等 自然条件优越^[16],且是固沟保塬、退耕还草的主要 区域^[14],加之坡耕地整治和小流域综合治理等一系 列生态修复工程的实施,土壤侵蚀明显减弱,植被 以及区域生态环境得到了极大改善^[41]. B区碳汇量相 对较低,这可能是由于早期该区过度放牧等不合理 的土地利用, 使生态环境遭到严重破坏, 但在退耕 还林/草等生态工程的支持下,碳储量共计增加 5.27×10⁴ t,还林还草的碳汇贡献率超过 50%, NEP 增长速率加快,植被覆盖度明显提高,取得了显著 的经济和生态效益^[42].C区是黄土高原主要的弱碳源 区域,气候干旱、降水稀少,不合理的水资源利用 导致地表植被减少, 植被的光合能力下降, 植被生 产力受到限制.该地区相较于其他分区,更容易受 到气候变化的干扰,整体碳汇增长率较低[40],且该 区受人类活动影响较大,长期的过牧滥牧造成严重 的草原退化和沙化,通过轮牧、禁牧和休牧等措施 的实施[43],以及生态工程的建设,使得该地生态治 理目标基本实现^[44].D区气候条件优越,林地广布, 植被覆盖率高^[16],属于显著的碳汇区.该区也是重 要的农业区,且聚集了西安、郑州和太原等大型城 市,区域经济活动频繁,受人类活动影响强度较 大,近年来实施的山水林田湖草沙保护和修复工 程, 增强了生态系统的固碳能力, 使得 NEP 年际变 化率较高[45].

3.2 空间尺度 NEP 变化分析

空间尺度上,相关分析和因子探测结果均表明,年降水量和年太阳总辐射是影响黄土高原NEP

8



Correlation coefficient and factor detection q value between NEP spatial distribution and driving factors in the Loess Plateau

表2 黄土高原及其生态分区交互探测结果

Table 2 Results of interactive detection of Loess Plateau and its ecological subregions

					8 8		
交互因子	全区	A1副区	A2副区	B1副区	B2副区	C区	D区
$P \cap aspect$	0.67	0.72	0.71	0.66	0.53	0.31	0.32
$P \cap ele$	0.72	0.75	0.83	0.78	0.82	0.54	0.55
$P \cap \text{slope}$	0.63	0.74	0.78	0.49	0.52	0.33	0.32
$P\cap \mathrm{HA}$	0.65	0.74	0.84	0.53	0.63	0.34	0.44
$P \cap SR$	0.64	0.71	0.76	0.60	0.66	0.47	0.39
$P \cap T$	0.72	0.77	0.87	0.86	0.74	0.50	0.56
$\mathrm{SR}\cap\mathrm{aspect}$	0.48	0.34	0.69	0.44	0.62	0.35	0.32
$\mathrm{SR}\cap \mathrm{ele}$	0.66	0.65	0.87	0.74	0.79	0.46	0.50
$\mathrm{SR}\cap\mathrm{slope}$	0.52	0.34	0.80	0.41	0.56	0.38	0.32
$\mathrm{SR}\cap\mathrm{HA}$	0.52	0.38	0.80	0.50	0.73	0.51	0.41
$SR \cap T$	0.63	0.70	0.89	0.74	0.68	0.44	0.48
$T \cap aspect$	0.26	0.60	0.27	0.66	0.34	0.46	0.26
$T \cap ele$	0.26	0.51	0.27	0.68	0.50	0.56	0.36
$T \cap \text{slope}$	0.37	0.60	0.38	0.58	0.50	0.45	0.39
$T \cap HA$	0.25	0.68	0.37	0.69	0.51	0.60	0.36
$ele \cap aspect$	0.25	0.44	0.22	0.64	0.50	0.34	0.32
$\mathrm{ele} \cap \mathrm{slope}$	0.31	0.44	0.43	0.63	0.54	0.37	0.38
$\mathrm{ele} \cap \mathrm{HA}$	0.17	0.59	0.40	0.68	0.54	0.39	0.37

空间分异的主要驱动因子,与之前大多数学者的研 究结果相一致^[30,33]. 黄土高原年降水量均呈东南高西 北低的空间格局^[46],与NEP空间分布高度一致.水 分是 NEP 变化的主要限制因子,降水的增加促进浅 根植物的发育,进而促进NEP的增加^[47];而黄土高 原太阳辐射空间上自东南向西北递增,与NEP空间 分布相反.热量的提高加大了区域的水分蒸散,使 土壤水分降低、植物受到干旱胁迫,光合能力下 降, 生态系统生产力受到制约, 进而抑制了植被的 碳吸收和干物质积累[48]. 交互探测的结果也证明了 这一点,黄土高原全区及6个生态分区 $P \cap SR$ 的q值均大于单个因子的q值或大于两个因子q值的和 (表2),表明年降水量和年太阳总辐射共同作用时 会增加对 NEP 的解释能力. 此外, B 区和 D 区与人类 活动呈负相关关系,这可能是该区域人口集聚效应 强,城乡建设用地急剧增长占用了大量的耕地,土 地利用格局发生变化,且煤炭开采等人类干扰对 NEP的增长产生了负面效应^[49].除B2副区和C区外, 黄土高原全区和其他分区的P∩T或SR∩T的q统计 值最高,表明气候因素是引起黄土高原 NEP 空间变 化的主要原因.

4 结论

(1)2000~2021年黄土高原全区及各分区年均 NEP均呈显著增长趋势,年均增长率为2.74~9.04 g•(m²•a)⁻¹,A2副区最大,C区最小.2021年77%的 区域为碳汇,与2000年相比碳汇面积增加了27%, 除C区为弱碳源外,其余各分区均为显著的碳汇.

(2)黄土高原年均 NEP呈现东南高西北低的空间分布格局.2000~2021年,98.55% 的区域固碳能力增长,A2副区中南部以及 B2副区西南部增幅较大.

(3)人类活动是影响黄土高原NEP时间变化的 主要因素,对NEP年际变化的贡献率为62.67%,气 候变化的贡献率为37.33%;气候因素是影响黄土高 原NEP空间变化的主要驱动因子.

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