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新疆车尔臣河流域绿洲带地下水咸化与污染主控因素

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摘要:车尔臣河流域绿洲带地下水是当地居民生产生活的重要水源,查明地下水咸化和污染的控制因素,对可持续开发利用地下水资源和保护沙漠绿洲带生态环境具有重要意义。在该区共采集36组单一结构潜水样品和54组多层结构潜水-承压水样品,在确定地下水主要离子分布特征和污染现状基础上,应用水文地球化学方法(Piper图、多元统计、Gibbs模型和离子间内联关系)综合揭示地下水咸化和污染的主控因素。研究区地下水从山前倾斜平原单一结构潜水至冲洪积平原多层结构潜水-承压水具有一定分带性,Cl-Na水(87.8%)是研究区地下水主要的水化学类型。单一结构潜水水质明显优于多层结构潜水-承压水,多层结构潜水-承压水成化主要由Na*(均值9969 mg·L⁻¹)、Cl⁻(均值13687 mg·L⁻¹)和SO₄²⁻(均值5840 mg·L⁻¹)导致,自然的水文地球化学过程是导致地下水水质恶化的主要原因。地下水化学主要受水岩作用和蒸发-浓缩作用控制,硅酸盐岩和蒸发盐岩矿物溶解是地下水化学组分的重要来源,而强烈的蒸发盐岩溶解过程对冲洪积平原区多层结构潜水-承压水的控制程度更高,加之在蒸发-浓缩和阳离子交换作用共同影响下,冲洪积平原区多层结构潜水-承压水咸化现象更为严重。此外,农田集中区农业施肥对地下水NO₅产生一定影响。

关键词:车尔臣河流域;地下水;咸化;污染;控制因素

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Controlling Factors of Groundwater Salinization and Pollution in the Oasis Zone of the Cherchen River Basin of Xinjiang

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Abstract: Gronndwater is a vital resource for local human fife and production in the oasis zone of the Cherchen River Basin of Xinjiang. Understanding the controlling factors of groundwater salinization and pollution is of great significance for the sustainable utilization of groundwater resources and protection of the ecological environment in desert oasis areas. In light of this, a total of 36 single structure unconfined groundwater samples and 54 multi-layered structure unconfined and confined groundwater samples were collected from the oasis zone of the Cherchen River Basin and evaluated for the distribution characteristics and pollution status of major ions. Hydrogeochemical methods (e. g., Piper diagram, multivariate statistics, Gibbs model, and relationships among ions) were used to determine the main controlling factors of groundwater salinization and pollution. Differences in hydrogeochemical zonation were found from the single structure unconfined aquifers in sloping plains of piedmont areas to the multi-layered structure unconfined and confined aquifers in alluvial-proluvial plain areas, and Cl-Na (87. 8%) was the main hydrochemical type in the groundwater of the study area. The quality of single structure unconfined groundwater was starkly better than that of the multi-layered unconfined and confined groundwater, which was mainly caused by Na⁺ (mean value of 9 969 mg·L⁺), Cl⁻ (13 687 mg·L⁺), and SO₄²⁻ (5 840 mg·L⁻). Moreover, the natural hydrogeochemical process was the main reason for the deterioration of groundwater quality. The hydrochemistry was mainly controlled by the water-rock interaction and evaporation processes. The mineral dissolution of silicates and evaporites was an important source of chemical ions in the groundwater. Furthermore, the chemical weathering of evaporites combined with the processes of evaporation and cation exchange had a significant influence on the salinization of multi-layered unconfined and confined groundwater in alluvial-proluvial plain areas. In add

Key words: Cherchen River Basin; groundwater; salinization; pollution; controlling factors

地下水是全球水文循环的一个重要组成部分^[1], 其水质的好坏直接影响着生态系统中物质循环、能量转换和信息传递. 然而,地下水环境在受原生地质背景影响或人类活动干扰时,水化学组分随之发生相应的响应变化,即被干扰的元素地球化学循环过程在一定程度会富集某些化学物质^[2~4]. 显然,地下水中毒害组分富集到一定程度会导致其失去原有的水资源功能,进而间接减少了可利用地下水资源 量[5]. 水质型缺水现象在国内外均有报道,特别是在地表水资源相对匮乏的干旱半干旱地区,地下水减化与污染问题已经对这些地区的人类健康、经济生产和生态环境产生了不利影响[6,7],进一步加剧了人

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向为地下水环境,E-mail:lipshydro@163.com * 通信作者,E-mail:zjzhoujl@163.com 类活动与水资源短缺之间的矛盾.深入揭示地下水 咸化与污染的主控机制,不仅为劣质地下水治理提 供有力的科学依据,同时对促进干旱半干旱地区人 类-社会-环境可持续发展具有深远的现实意义.

地下水化学组分主要受来源和迁移转化过程控制^[8],就水化学组分来源而言,总体可划分为自然背景输入和人类活动排放这2大类^[1],同时部分来源受自然因素和人类活动双重影响,如酸性污染物参与的水岩作用过程^[9];水化学组分迁移转化过程相对来源更为复杂,地下水在下渗途径和水平运移过程中,通常会发生一系列的水文地球化学作用,如水岩作用、蒸发-浓缩作用、氧化还原作用和阳离子交换作用等^[10],并伴随着外源物质的进一步入渗.面对地下水化学复杂的形成机制,国内外在揭示地下水化学特征基础上,逐渐探索出一系列较为可靠的研究体系,这为探索干旱半干旱地区地下水咸化与污染成因机制提供了方法学基础.

车尔臣河流域是我国塔里木盆地南缘较大的内 陆河流域,绿洲带多与荒漠相邻,并与其它绿洲带相 隔[11],强蒸发和弱降水的干旱炎热气候迫使当地居 民不得不开发利用地下水资源. 据已有研究, 车尔臣 河流域地下水存在一定咸化现象[7],同时地表水受生 活污水、农业化肥和畜禽粪便等多种污染排放影 响^[12]. 据 δ^2 H-H₂O和 δ^{18} O-H₂O证据,该流域地下水主 要接受地表河水、灌溉水和大气降水等补给[11],在补 给水下渗过程中,污染物通常随补给来源一同进入 含水层,这为地下水带来了潜在的污染风险[3]. 然而, 当前车尔臣河流域绿洲带地下水咸化和污染的控制 机制研究相对有限.基于此,本文利用车尔臣河流域 90组地下水化学样品(36组单一结构潜水样品和54 组多层结构浅层潜水-承压水样品),在调查地下水化 学分布特征基础上,利用 PIG (pollution index of groundwater)对该区地下水进行污染评价,利用水文 地球化学方法和多元统计方法综合确定地下水咸化 的主控因素,以期为车尔臣河流域绿洲带地下水咸 化与污染防治提供科学基础.

1 材料与方法

1.1 研究区概况

研究区主要位于新疆巴音郭楞蒙古自治州若羌县和且末县域内,地理位置为东经83.75°~89.67°,北纬36.19°~39.81°(图1).本区属温带大陆性干旱气候,年均气温约为10.1°C,日均温差达15.9°C.年均降水量仅为18.6 mm,其中最小年降水量低至1.9 mm,干旱指数为77,属极端干旱区.干旱炎热的气候条件,导致年均蒸发量高达2506.9 mm. 匮乏的降水

量和巨大的蒸发量,加之全年盛行的东北风,直接导致研究区呈现多风沙、浮尘和沙暴天气,这进一步加块了沙漠向西南地区推进的速度,一定程度恶化了若羌县和且末县绿洲带生态环境.车尔臣河是塔里木盆地东南最大的河流,多年平均径流量为8亿m³[13],其补给来源主要为冰雪融化水和地下水.

车尔臣河流域地势总体呈南高北低,由南至北分别为山区、冲洪积扇、冲洪积平原和沙漠地貌.区内除现代河床和沟谷外,多被第四系松散堆积物覆盖,岩性主要为上细下粗的亚砂土、卵砾石,底部多分布西域砾岩和泥岩(图1).据含水层类型地下水可划分为山前倾斜平原的单一结构潜水和冲洪积平原的多层结构潜水-承压水,地下水位埋深多介于0.8~165 m^[11],地下水位埋深≤5 m的区域占区内总面积的64.6%.潜水分布于山前倾斜平原,承压区分布于冲洪积平原,接近沙漠区.地下水主要补给来源为河道水下渗、渠系水和灌溉水入渗,径流方向与地形坡向和河流流向基本一致,受气候条件和低洼地形控制,浅层地下水蒸发和补给河道是主要的排泄方式.此外,冲洪积平原区承压水循环速率整体偏慢.

1.2 样品采集与测试

车尔臣河流域西北为塔克拉玛干沙漠,西南毗 邻荒漠和山区,仅车尔臣河流域呈长条形绿洲带,这 为地下水样品采集带来了极大困难,近年来整个流 域同期地下水化学资料极其有限. 在收集研究区水 文地质资料基础上,本着尽可能覆盖全流域原则,于 2014~2021年共采集地下水化学样品90组,其中单 一结构潜水样品36组(山前倾斜平原),这部分样品 可反映人类活动极弱区地下水化学特征;多层结构 潜水-承压水区样品54组(冲洪积平原),这部分样品 一定程度反映出生活和农业活动区地下水化学特征 (图1). 主要离子检测工作由新疆地质矿产局第二水 文地质工程地质大队化验室完成. 采用多参水质检 测仪(HANNA, HI9828)原位检测pH指标,采用火焰 原子吸收分光光度法检测 K+和 Na+,采用乙二胺四乙 酸二钠滴定法检测 Ca2+和 Mg2+,采用硝酸银容量法检 测 CI⁻,采用硫酸钡比浊法检测 SO₄²⁻,采用酸碱滴定法 检测 HCO3-,采用紫外分光光度法检测 NO3,采用离子 选择电极法检测下,采用干燥-比重法检测溶解性总 固体(TDS). 水化学样品设有空白样、空白加标样以 及平行样,确保主要离子检测结果的可靠度.水化学 数据采用电荷平衡误差[%CBE,式(1)]检验检测结 果的准确性:

$$\%CBE = \frac{\sum \text{阳离子} - \sum \text{阴离子}}{\sum \text{阳离子} + \sum \text{阴离子}} \times 100$$
 (1)

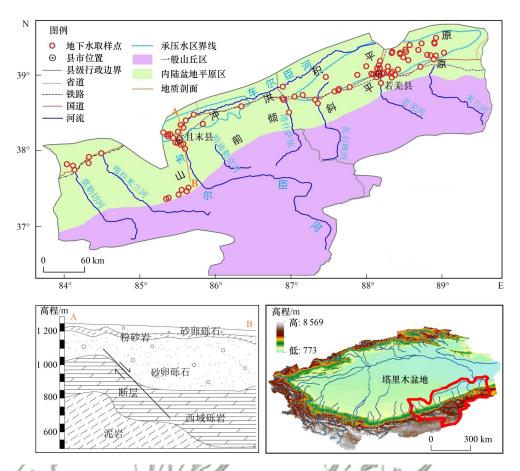


图 1 车尔臣河流域地下水采样点分布及水文地质示意

Fig. 1 Distribution map of groundwater sampling sites and hydrogeological diagram of Cherchen River Basin

式中, \sum 阳离子表示地下水中 K^+ 、 Na^+ 、 Ca^{2+} 和 Mg^{2+} 浓度 累加值, \sum 阴离子表示地下水中 Cl^- 、 SO_4^{2-} 、 HCO_3^- 、 NO_3^- 和 F^- 浓度累加值,地下水样品的%CBE $(-6.4\% \sim 4.2\%)$ 绝对值均低于 $10\%^{[14]}$.

1.3 地下水污染评价

PIG 评价法最早由 Subba Rao 于 2012 年提出^[15],该方法主要用于评价多种水文地球化学作用导致的地下水水质变化^[16],具体过程划分为 5 步. 首先确定各水化学指标的 RW(relative weight),据各水化学指标对人体的毒害特性,于 1 ~ 5 之间赋值,本研究的RW值参考文献 [15];其次,确定WP(weight parameter),各水化学指标的WP值计算如下:

$$WP = \frac{RW}{\sum RW}$$
 (2)

第3步确定 SC(status of concentration):

$$SC = \frac{C_i}{DS_i}$$
 (3)

式中, C_i表示各评价指标的实际检测浓度, DS_i表示饮用水标准限值,各评价指标的 DS_i值参考《地下水质量标准》(GB/T 14848-2017) Ⅲ类水限值和《生活饮用水卫生标准》(GB 5749-2022) 限值^[17,18], 我国水质标

准无对应参考值时,DS,值取自文献[15].

第4步,利用各水化学指标计算的WP和SC值,确定OW(overall groundwater quality)值:

$$OW = WP \times SC$$
 (4)

最后,将被评价的每种水化学指标的OW值进行 累加,获得地下水样的最终PIG值:

$$PIG = \sum OW \tag{5}$$

相应的计算参数取值见表 1. 据 Subba Rao 给出的地下水污染等级划分标准 [15],单个水化学指标 (OW) 和地下水总体评价结果 (PIG) 可划分为: OW/PIG < 1.0时为无明显污染; 1.0 \leq OW/PIG < 1.5时为低污染; 1.5 \leq OW/PIG < 2.0时为中污染; 2.0 \leq OW/PIG < 2.5 时为高污染; 2.5 \leq OW/PIG 时为重污染.

2 结果与分析

2.1 地下水化学特征

研究区地下水化学特征统计见表 2. 地下水 pH 总体介于 6. 30~8. 43,均值为 7. 85,多属中性至弱碱性水,单一结构潜水 pH(8. 05)略高于多层结构潜水-承压水 (7.72). ρ (TDS) 总体介于 423~ 359 376

表1 PIG计算涉及的相关参数取值1)

Table 1 Values of parameters related to PIG

指标	pН	TDS	K^{+}	Na^+	Ca ²⁺	Mg^{2+}	Cl-	SO_4^{2-}	HCO ₃	NO_{3}^{-}	F^{-}	合计
RW	5	5	1	4	2	2	4	5	3	5	5	41
WP	0.122	0.122	0.024	0.098	0.049	0.049	0.098	0.122	0.073	0.122	0.122	1.000
DS	7.5	1000	10	200	75	30	250	250	300	45	1.0	/

1)"/"表示没有相关数据

mg·L⁻¹,均值为 31 407 mg·L⁻¹,其中,单一结构潜水样品中淡水[ρ (TDS) < 1 000 mg·L⁻¹]占比为 33. 3%,多层结构潜水-承压水样品中淡水占比仅为 6. 0%,车尔臣河流域地下水多为微咸水及以上级别,且多层结构潜水-承压水的咸化程度高于单一结构潜水.由表2可知,研究区地下水较高的 TDS 主要由 Na⁺(均值9 969 mg·L⁻¹)、Cl⁻(均值13 687 mg·L⁻¹)和 SO₄²⁻(均值5 840 mg·L⁻¹)引起,这在一定程度反映出蒸发-浓缩作用对冲洪积平原区地下水化学的影响特征.

研究区地下水阳离子浓度平均值顺序为: $Na^* > Mg^{2*} > K^* > Ca^{2*}$,其中多层结构潜水-承压水阳离子浓度平均值顺序与研究区地下水保持一致,单一结构潜水阳离子浓度平均值顺序与多层结构潜水-承压水略有差异,即: $Na^* > Ca^{2*} > Mg^{2*} > K^*$. 地下水中阴离子浓度平均值顺序为: $Cl^* > SO_4^{2*} > HCO_3^{-*} > NO_3^{-*} > F^-$,单一结构潜水和多层结构潜水-承压水阴离子浓度顺序保持一致. 变异系数可反映地下水主要离子的空间分布差异[19],研究区地下水阳离子 $K^*(4.33, 变异系数)$ 和 $Mg^{2*}(3.22, 变异系数)$ 以及阴离子 $Cl^*(2.58, 变异系$

数)和 $SO_4^{2-}(2.15, 变异系数)$ 受空间尺度影响较大,这 些离子的空间变异性与研究区不同地貌类型分布 相关.

Piper图是水文地球化学分析常用的一种工 具[20],为进一步突出主导离子和主要水化学类型,部 分研究对 Piper 图进行分区和优化,即依据地下水中 主要离子的毫克当量百分比是否高于50%来确定主 导离子和主要水化学类型[4,21,22]. 由图 2 可知,车尔臣 河流域地下水主导阳离子主要为 K+Na+型(84.4%), 无主导型(14.4%)和 Mg2+型(1.1%)占比并不大,主导 阴离子主要为无主导型(55.6%),其次为CI型 (22.2%)和SO²型(22.2%). 单一结构潜水和多层结 构潜水-承压水主导阳离子类型相似,主导阴离子类 型存在一定差异,单一结构潜水(30.6%)中SO₄-型水 样比例较多层结构潜水-承压水(16.7%)高,多层结 构潜水-承压水(31.5%)CI型水样比例较单一结构潜 水(8.3%)高. 研究区地下水化学类型为:Cl-Na水 $(87.8\%) > \text{Cl-Ca} \cdot \text{Mg } \% (10.0\%) > \text{SO}_4 - \text{Ca } \%$ (2.2%),表明车尔臣河流域地下水中Na⁺和Cl⁻富集

表 2 车尔臣河流域地下水水质常规指标数理统计1)

Table 2 Statistics of regular ions in groundwater of the Cherchen River Basin

类型	统计量	рН	K ⁺	Na ⁺	Ca ²⁺	${\rm Mg}^{2+}$	Cl-	SO ₄ ²⁻	HCO ₃	NO ₃	F^{-}	TDS
	最大值/mg·L ⁻¹	8.43	722	5 906	562	1 095	10 310	6 236	757	32.0	14.7	25 407
	最小值/mg·L ⁻¹	7.14	4.96	39.2	16.7	9.57	35.4	105	45.5	0.36	0.34	423
单一结构潜水	平均值/mg·L ⁻¹	8.05	40.2	483	111	102	711	667	221	9.08	1.69	2 345
	标准偏差/mg·L ⁻¹	0.31	119	989	117	183	1 703	1 098	166	8.84	2.52	4 240
	变异系数	0.04	2.96	2.05	1.06	1.79	2.40	1.65	0.75	0.97	1.50	1.81
	最大值/mg·L ⁻¹	8.34	17 898	115 621	1 102	25 366	172 922	54 148	1 428	138	8.27	359 376
多层结构潜水- 承压水	最小值/mg·L ⁻¹	6.30	5.33	109	36.9	34.3	149	211	96.5	0.03	0.07	829
	平均值/mg·L-1	7.72	734	16 293	318	1 499	22 337	9 288	295	16.8	1.38	50 781
	标准偏差/mg·L-1	0.49	2 530	30 880	292	3 815	43 642	15 267	230	24.8	1.40	94 091
	变异系数	0.06	3.45	1.90	0.92	2.54	1.95	1.64	0.78	1.47	1.02	1.85
	最大值/mg·L ⁻¹	8.43	17 898	115 621	1 102	25 366	172 922	54 148	1 428	138	14.7	359 376
	最小值/mg·L ⁻¹	6.30	4.96	39.2	16.7	9.57	35.4	105	45.5	0.03	0.07	423
地下水	平均值/mg·L-1	7.85	458	9 969	235	940	13 687	5 840	266	13.53	1.50	31 407
	标准偏差/mg·L ⁻¹	0.45	1 983	25 078	258	3 025	35 339	12 543	209	19.9	1.92	76 476
	变异系数	0.06	4.33	2.52	1.10	3.22	2.58	2.15	0.79	1.47	1.28	2.43
地下水质量	量标准(Ⅲ类)	6.5 ~ 8.5	/	200	/	/	250	250	/	88.57	1.00	1 000
生活饮用	水卫生标准	6.5 ~ 8.5	/	200	/	/	250	250	/	44.29	1.00	1 000

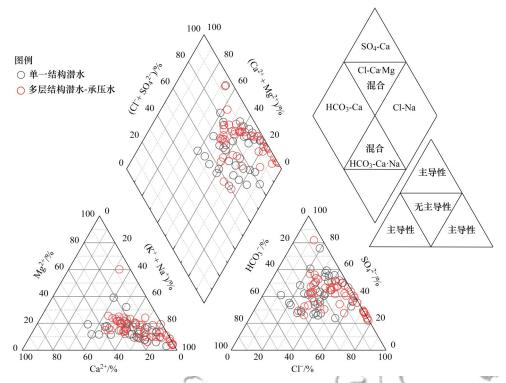


图 2 车尔臣河流域地下水主要离子 Piper 图

Fig. 2 Piper diagram of the major ionic concentrations in groundwater of the Cherchen River Basin

程度较高,其次为 Ca²⁺和 Mg²⁺. 因此,矿物溶解和阳离子交换可能是控制研究区地下水化学类型的重要因素^[21].

2.2 地下水污染评价

车尔臣河流域地下水污染评价结果见图 3. 由 OW 值可知[图 3(a)],引起地下水污染的主要水化学 指标有TDS、K+、Na+、Mg2+、Cl-、SO42和F,其中,部 分多层结构潜水-承压水受 Cl⁻(67.8, OW 值)、Na⁺ (56.7,OW 值)、K*(43.0,OW 值)、Mg²⁺(41.4,OW 值) 和 SO₄²⁻(26.4,OW 值)的极大值影响,表现为重污染 等级,多层结构潜水-承压水中这些水化学指标的 OW 值比单一结构潜水高 8.7~24.9倍.单一结构潜 水 PIG 值介于 0.45~19.1,平均值为 1.88[图 3(b1) 和 3(b2)],属无明显污染、低污染、中污染、高污染 和重污染等级的样品占比分别为38.9%、33.3%、 11.1%、11.1%和5.6%. 多层结构潜水-承压水PIG 值介于 0.82~266, 平均值为 32.3 [图 3(c1)和 3 (c2)],属无明显污染、低污染、中污染、高污染和重 污染等级的样品占比分别为11.1%、24.1%、 11.1%、5.6%和48.2%.显然,多层结构潜水-承压水 水质较单一结构潜水差. 据水文地质调查结果,承压 区位于冲洪积平原区,而潜水区则分布于山前倾斜 平原区(图1),Cl、Na*和K*而非具有典型农业和生活 排放特征的NO;是导致劣质多层结构潜水-承压水的 主要化学组分[23]. 因此,劣质多层结构潜水-承压水

可能由矿物溶解、蒸发浓缩和阳离子交换等水文地球化学作用综合导致.

3 讨论

3.1 多元统计分析

研究区地下水主要离子相关性分析结果见图 4 (a). 单一结构潜水 TDS、 K^* 、 Na^* 、 Mg^{2*} 、 $C\Gamma$ 、 SO_4^{2-} 和 F^* 间的相关系数 r 值介于 $0.91 \sim 1.00$,呈强正相关关系,多层结构潜水-承压水 TDS、 Na^* 、 $C\Gamma$ 和 SO_4^{2-} 以及 K^* 和 Mg^{2*} 间的 r 值介于 $0.97 \sim 1.00$,呈强正相关关系,说明这些离子具有相同的来源和迁移转化途径 [24,25]. 多层结构潜水-承压水中 NO_3^* 与 K^* 呈一定正相关关系 (r=0.72),这在一定程度反映出农业种植区地下水可能受农业施肥的影响 [19].

主成分分析(PCA)通过降低自变量或数据集的维数,以解释大数据集内相关变量的方差[26.27],该方法已被广泛用于地下水化学成因分析研究中[28-30].为进一步揭示影响车尔臣河流域绿洲带地下水化学的控制因素,分别对单一结构潜水[图4(b)]和多层结构潜水-承压水[图4(c)]主要离子进行PCA.单一结构潜水和多层结构潜水-承压水均获得2个PCs,单一结构潜水中,PC1解释度为74.1%,PC2解释度为10.3%,累计解释度为84.4%;多层结构潜水-承压水中,PC1解释度为59.5%,PC2解释度为13.8%,累计解释度为73.3%.PCA分析结果与相关性分析结果基

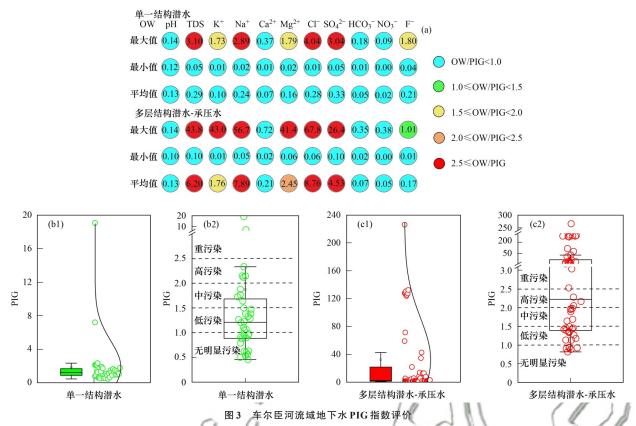


Fig. 3 Evaluation of PIG index of groundwater in the Cherchen River Basin

本一致,总体看,多层结构潜水-承压水化学成因机制相对单一结构潜水更为复杂.就单一结构潜水而言,PC1可解释pH、TDS、K⁺、Na⁺、Ca²⁺、Mg²⁺、Cl⁻、SO₄²⁻和 F⁻,这主要反映出地质背景对地下水化学的控制;PC2可解释NO₃和 HCO₅,且二者呈相互抑制关系,这反映出人类活动对地下水化学的影响.就多层结构潜水-承压水,PC1可解释pH、TDS、K⁺、Na⁺、Ca²⁺、Mg²⁺、Cl⁻和SO₄²⁻;PC2可解释NO₃、HCO₃⁻和 F⁻,反映出多层结构潜水-承压水化学组分受地质背景和人类活动双重控制,且位于农业活动相对集中的冲洪积平原区,这类水受农业活动排放影响更大.

3.2 地下水化学成因解析

3.2.1 水岩作用

Gibbs模型最早由 Gibbs 于 1970年提出[31,32],此模型主要依据地表水中 Cl⁻、HCO₃⁻、Na⁺和 Ca²⁺与 TDS 间的内联关系,将地表水化学控制机制划分为蒸发-浓度作用、水岩作用和降水作用这 3 大类.相对地表水,地下水水岩作用过程更为充分,为更真实地反映出地下水化学控制机制,将 Gibbs模型中反映水岩作用的控制区域进行适当优化,改进后的 Gibbs模型在地下水化学成因控制解析中得以广泛应用[33-35].由图 5(a)和 5(b)可知,车尔臣河流域绿洲带地下水化学主要受水岩作用和蒸发-浓缩作用双重控制,而研究区极低的降水量(年均 18.6 mm)对地下水化学成

因几乎不产生影响,这与塔里木盆地其它流域研究结果是一致的^[7,35],结合主成分分析结果和研究区水文地质条件,PC1中多数离子[图 4(b)和 4(e)]主要来源于可溶性盐岩(如蒸发盐岩、硅酸盐岩和碳酸盐岩等)溶解.在干旱炎热气候条件下,加之区内地下水位埋深普遍较浅,较大的蒸发量(年均 2 506.9 mm)对地下水多数离子具有明显的浓缩作用,特别是冲洪积平原多层结构潜水-承压水表现更为明显.车尔臣河流域地下水中TDS、K*、Na*、Cl⁻和 SO₄²⁻浓度普遍高于我国东北平原、华北平原、云贵高原和内蒙古高原等地^[4,22,36,37],水岩作用和强烈的蒸发-浓缩作用是导致这一差异的重要原因.

水岩作用是一个自然发生的化学风化和矿物溶解过程,可溶性盐岩因矿物组成差异释放出的水化学组分各不相同,如盐岩(Na⁺和 Cl⁻)、石膏(Ca²⁺和 SO₄²⁻)、方解石(Ca²⁺和 HCO₃⁻)、白云石(Ca²⁺、Mg²⁺和 HCO₃⁻)^[4,38].含水层中不同矿物溶解释放出的离子具有各自的比例关系^[19],因此可利用这一关系来解析不同盐岩对地下水化学的控制作用.由图 5(c)和 5(d)可知,绝大多数单一结构潜水和多层结构潜水-承压水样点接近全球硅酸盐岩平均风化端元,说明车尔臣河流域含水层中发生的水岩作用主要表现为硅酸盐岩化学风化.部分多层结构潜水-承压水样点接近蒸发盐岩化学风化。部分多层结构潜水-承压水样点接近蒸发盐岩化学风化端元,说明这类水受蒸发盐岩

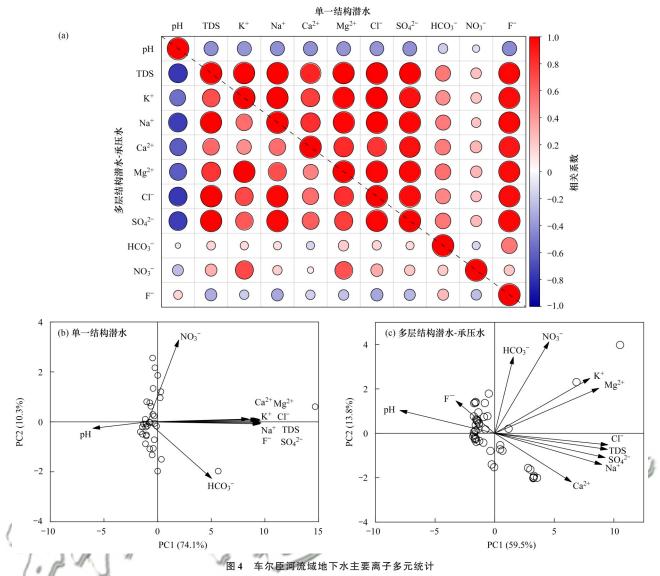


Fig. 4 Statistical analysis of major compositions in groundwater of the Cherchen River Basin

化学溶解影响比单一结构潜水强,这与多层结构潜水-承压水中 K*(734 mg·L⁻¹)、Na*(16 293 mg·L⁻¹)和 Cl⁻(22 337 mg·L⁻¹)浓度平均值均高于单一结构潜水中 K*(40.2 mg·L⁻¹)、Na*(483 mg·L⁻¹)和 Cl⁻(711 mg·L⁻¹)的结果是一致的(表 2).同时,冲洪积平原区多层结构潜水-承压水咸化与强烈的蒸发-浓缩作用和较慢的地下水循环速率也存在一定关系.此外,部分山前平原区单一结构潜水样点偏向碳酸盐岩溶解端元,说明含水层中少量的碳酸盐岩(如方解石和白云石)也参与了化学风化过程.

3.2.2 阳离子交换

在地下水化学演化过程中,阳离子交换作用是影响地下水化学组成的另一个重要因素[35],当阳离子发生正向交换时,含水层中矿物颗粒将原本吸附的 K^+ 和 Na^+ 解吸,而电荷较强的 Ca^{2+} 和 Mg^{2+} 被吸附,从而提高地下水 K^+ 和 Na^+ 浓度.利用 $(Ca^{2+}+Mg^{2+}-SO_4^{2-}+HCO_3^-)$ 与 (Na^+-Cl^-) 间的线性拟合关系可解释地下水

中阳离子交换作用发生的水平,即(Ca2++Mg2+-SO4--HCO,⁻)和(Na⁺-Cl⁻)值均为0 meg·L⁻¹时,说明地下水 主要离子来源于矿物溶解,(Ca²⁺+Mg²⁺-SO₄²⁻-HCO₃-) 值低于(Na⁺-CI⁻)值时,说明地下水发生阳离子正向 交换, 若两者相反则发生阳离子负向交换, 当(Ca2++ Mg²⁺-SO₄²⁻-HCO₃-)/(Na⁺-Cl⁻)值接近1时,说明阳离子 交换作用是控制地下水化学成因的重要因素[39,40]. 车 尔臣河流域地下水中(Ca²⁺+Mg²⁺-SO₄²⁻-HCO₄-)与(Na⁺ -CI⁻)拟合关系如图 6(a),全部的单一结构潜水样点 和72.2%的多层结构潜水-承压水样点分布于0点附 近,表明地下水中主要离子来源于矿物化学溶解,这 与 Gibbs 分析结果一致[图 5(a)和 5(b)]. (Ca2+ + Mg2+ - $SO_4^{2-}-HCO_5^{-}$)与(Na⁺-Cl⁻)的拟合斜率为-0.76(R^2 = 0.98),表明研究区地下水已经发生了明显的阳离子 交换作用,且22.2%的多层结构潜水-承压水样点受 阳离子正向交换影响而富集低电荷离子(K*和 Na*), 而阳离子逆向交换作用对研究区地下水化学的控制

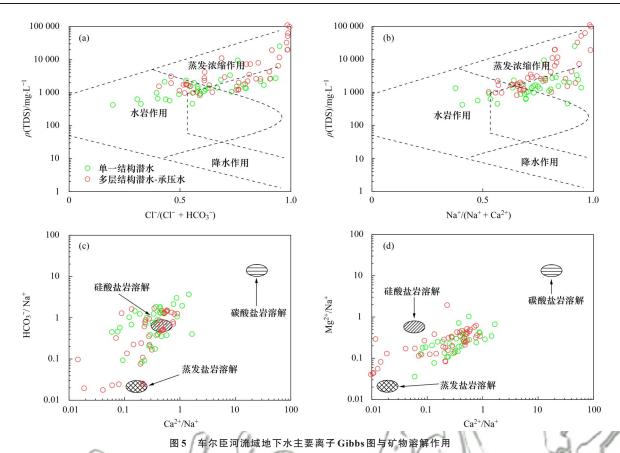


Fig. 5 Gibbs diagram of the major compositions and processes of water-rock interaction in groundwater of the Cherchen River Basin

并不明显.

CAI 指数(chloro-alkaline index, meq·L⁻¹)可用于 分析地下水中阳离子发生的交替吸附作用^[35,41],计算 公式如下:

$$CAI1 = \frac{CI^{-} - (K^{+} + Na^{+})}{CI^{-}}$$

$$CAI2 = \frac{CI^{-} - (K^{+} + Na^{+})}{SO_{4}^{2} + HCO_{3}^{-}}$$
(7)

当地下水中 Ca2+和 Mg2+被吸附, K+和 Na+被解吸 时,CAI均表现为负值,反之亦然,当CAI值接近0时, 说明阳离子交换作用对地下水化学影响较小,而CAI 的绝对值越大说明阳离子交换程度越高[42]. 由图 6 (b)可知,80.6%的单一结构潜水样点和66.7%的多 层结构潜水-承压水样点 CAI1 和 CAI2 值小于 0, 表明 在地下水演化过程中阳离子发生了正向交换作用, 结果导致地下水中 K⁺和 Na⁺浓度升高, Ca²⁺和 Mg²⁺因 被矿物颗粒吸附而浓度降低. 研究区地下水中 CAII 和 CAI2 值域分布较大,这可能与山前至平原的不同 水文地质条件相关[35],含水层由山区至冲洪积平原, 粉砂颗粒逐渐增加,地下水流速随之缓慢,地下水位 埋深变浅,冲洪积平原地下水具有更长的水岩作用 时间,也更易发生蒸发-浓缩和阳离子交换作用,从而 明显提高了冲洪积平原区多层结构潜水-承压水中 K⁺和 Na⁺浓度,这是导致研究区地下水咸化的重要

原因.

3.2.3 人类活动影响

据水文地质调查结果,车尔臣河流域绿洲带共 涉及且末县和若羌县2个行政区,区内主要以农业活 动为主,工业活动分布较少. 地下水中NO;浓度超过 《生活饮用水卫生标准》(GB 5749-2022)的采样点仅3 组[18],均位于若羌县农业集中区. 因此,有必要利用 地下水中具有典型农业排放特征的 NO; 和相对稳定 且具有人类活动排放特征的CIT建立关系[23,43-45],以 分析人类活动对地下水污染的影响.由图7(a)可知, 研究区地下水受自然来源和人类活动排放综合影 响.利用CI/Na+和NOJ/Na+间的关系进一步验证人类 活动排放对地下水化学的影响[图7(b)],尽管多数 地下水样点集中在可溶性盐岩控制端元,但仍然存 在部分样点偏向农业活动端元,这说明除蒸发盐岩、 硅酸盐岩和碳酸盐岩溶解对车尔臣河流域地下水化 学起主控作用外,农业活动排放也在一定程度影响 着地下水水质,通常,农业集约区农业化肥、生活污 水、人畜粪便和土壤有机氮是地下水 NO; 的典型污 染来源[23,46,47],然而本研究中3组NO。超标样点均位 于若羌县农田集中区,同时该区域多层结构潜水-承 压水中NO3与K+呈一定正相关关系[图4(a)],这反 映出地下水很可能受农业施肥影响[19]. 此外,且末县

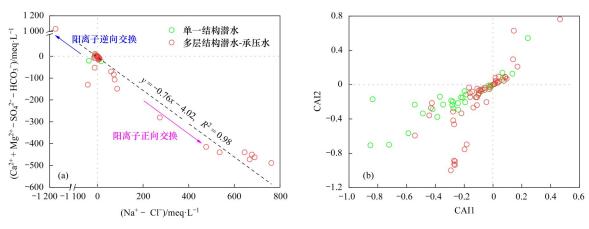


图 6 车尔臣河流域地下水阳离子交换过程

Fig. 6 Processes of cation exchange in groundwater of the Cherchen River Basin

和若羌县农业种植区主要分布在冲洪积平原的承压区(图1),冲洪积平原多层结构潜水-承压水中NO₃浓度[(16.8 ± 24.8)mg·L⁻¹]整体高于山前平原单一结构潜水中NO₃浓度[(9.08 ± 8.84)mg·L⁻¹](表 2),且冲洪积平原多层结构潜水-承压水中NO₃浓度

[(16.8 ± 24.8) $mg \cdot L^{-1}$] 也高于地下水 NO_3 浓度背景值(13.53 $mg \cdot L^{-1}$) [8],而山前平原单一结构潜水中 NO_3 浓度[(9.08 ± 8.84) $mg \cdot L^{-1}$]接近地下水 NO_3 浓度背景值,这进一步说明农业施肥是冲洪积平原区地下水 NO_3 污染的主要来源.

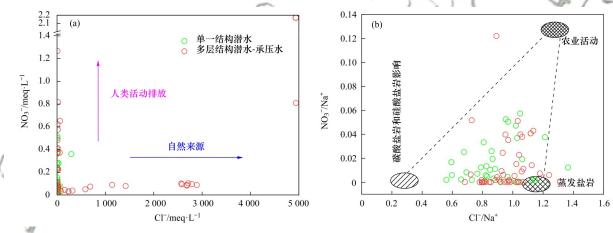


图 7 人类活动对车尔臣河流域地下水主要离子影响 Fig. 7 Influence of anthropogenic activities on major compositions in groundwater of the Cherchen River Basin

4 结论

(1)车尔臣河流域地下水多属中性至弱碱性水,阳离子浓度平均值顺序为: $Na^+ > Mg^{2+} > K^+ > Ca^{2+}$,阴离子浓度平均值顺序为: $Cl^- > SO_4^{2-} > HCO_3^- > NO_3^- > F^-$,主导阳离子为 $K^+ + Na^+ \mathbb{Z}(84.4\%)$,主导阴离子为无 主导 型 (55.6%)、 $Cl^- \mathbb{Z}(22.2\%)$ 和 $SO_4^{2-} \mathbb{Z}(22.2\%)$,地下水化学类型为: Cl-Na 水(87.8%) > $Cl-Ca\cdot Mg$ 水(10.0%) > SO_4-Ca 水(2.2%). 地下水咸化主要由 Na^+ (均值 9 969 $mg\cdot L^{-1}$)、 Cl^- (均值 13 687 $mg\cdot L^{-1}$)和 SO_4^{2-} (均值 5 840 $mg\cdot L^{-1}$)引起.

(2)据 PIG 评价结果,研究区单一结构潜水和多层结构潜水-承压水属于无明显污染和低污染等级,占比分别为 72.2% 和 35.2%,单一结构潜水水质明显优于多层结构潜水-承压水水质.引起多层结构潜

水-承压水污染的主要水化学指标为 $C\Gamma$ 、 Na^{\dagger} 和 K^{\dagger} ,而非 NO_3^{\bullet} ,说明自然的水文地球化学过程而非人类活动排放是导致车尔臣河流域绿洲带地下水水质恶化的主要原因.

(3)水岩作用和蒸发-浓缩作用是控制研究区地下水化学的主要因素.硅酸盐岩和蒸发盐岩的矿物溶解是地下水化学组分的重要来源,且蒸发盐岩溶解对冲洪积平原区多层结构潜水-承压水影响更大.由于炎热干旱的气候条件,蒸发-浓缩作用和阳离子交换作用进一步提高了冲洪积平原区多层结构潜水-承压水 K*和 Na*浓度,很大程度加剧了这类地下水咸化程度.此外,部分农业区地下水受农业施肥影响而出现 NO。超标现象.

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