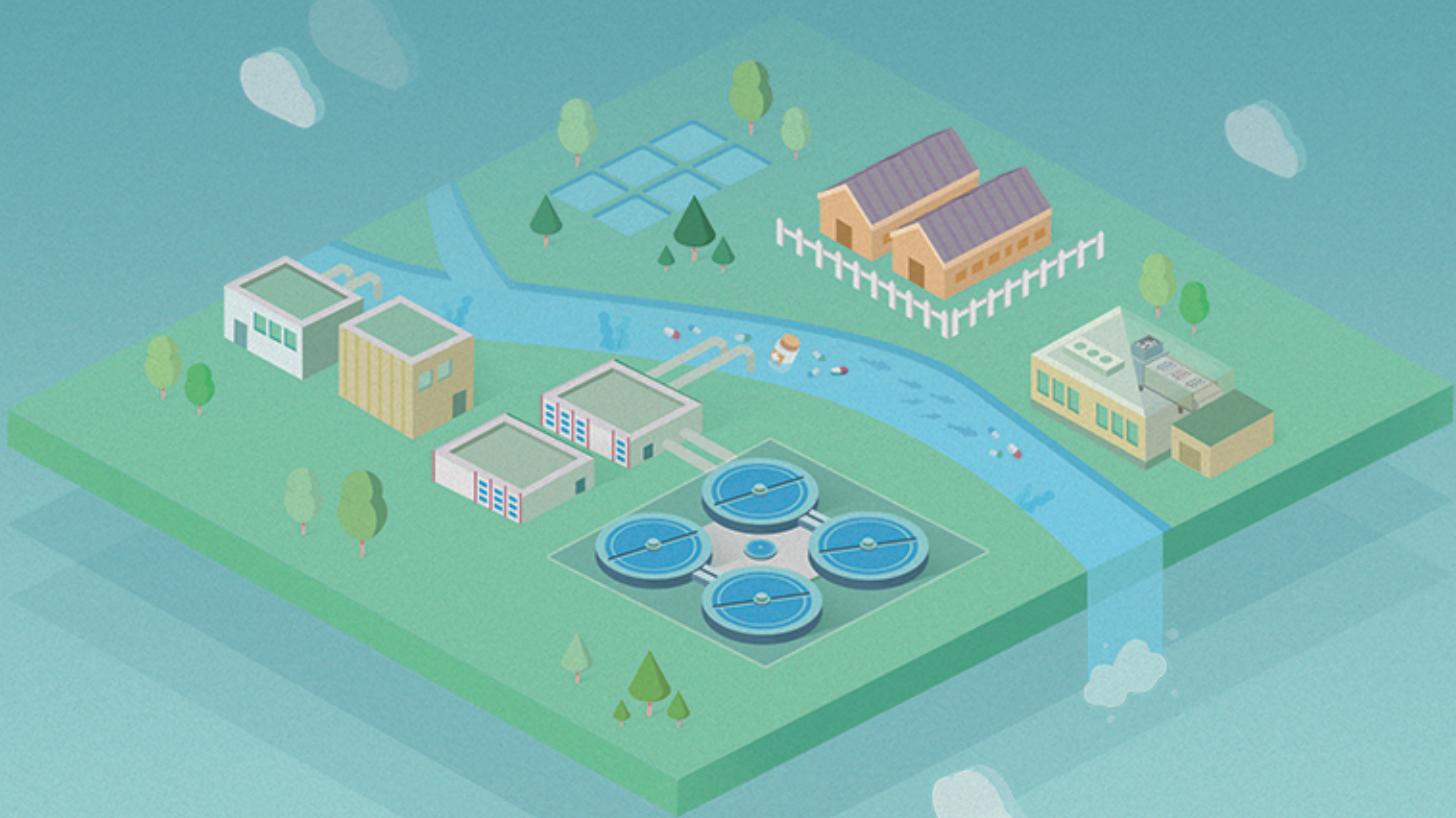


长江中游典型饮用水水源中药物的时空分布及风险评价

武俊梅，魏琳，彭晶倩，何鹏，施鸿媛，汤冬梅，吴振斌



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重庆农田土壤有机碳稳定性同位素空间分布特征

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摘要: 对重庆 182 个典型农田土壤剖面有机碳稳定性同位素组成($\delta^{13}\text{C}_{\text{SOC}}$)的测定结果表明, 所有剖面土壤 $\delta^{13}\text{C}_{\text{SOC}}$ 值均随采样深度增加逐渐趋正, 表、中和底层均值分别为 $(-23.63 \pm 1.53)\text{\%e}$ 、 $(-22.43 \pm 1.59)\text{\%e}$ 和 $(-21.42 \pm 1.90)\text{\%e}$ 。就地域而言, 渝东北土壤 $\delta^{13}\text{C}_{\text{SOC}}$ 值偏负程度最高, 渝中土壤则偏正。水田 $\delta^{13}\text{C}_{\text{SOC}}$ 值明显偏负, 旱地偏正, 水旱轮作则居中; 三者表层土壤 $\delta^{13}\text{C}$ 均值分别为 $(-25.32 \pm 0.93)\text{\%e}$ 、 $(-23.17 \pm 1.37)\text{\%e}$ 和 $(-24.75 \pm 1.28)\text{\%e}$; 不同类型土壤表层 $\delta^{13}\text{C}$ 均值依序为: 水稻土 < 潮土 < 紫色土 < 石灰(岩)土 < 黄壤。回归树分析表明, 表层土壤 $\delta^{13}\text{C}_{\text{SOC}}$ 值主要受作物类型控制, 中底层则主要与土壤类型有关; 其它因素如土壤性质(总氮、SOC 和 pH)和气象条件(降雨量和气温)等也有一定的影响。换言之, 表层土壤稳定性碳同位素分布主要受碳源的影响, 中底层主要与土壤碳循环过程有关。

关键词: 农田; 土壤有机碳(SOC); 稳定性同位素组成; 空间分布; 主控因素

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Spatial Characterization of Stable Isotope Composition of Organic Carbon from Farmland Soils in Chongqing

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Abstract: Soil was sampled from 182 profiles in typical farmlands of Chongqing and analyzed for the stable carbon isotope composition of organic matter ($\delta^{13}\text{C}_{\text{SOC}}$)。The results showed that the values of $\delta^{13}\text{C}_{\text{SOC}}$ for each soil profile were gradually increasing with increasing soil depth, and the mean values were $(-23.63 \pm 1.53)\text{\%e}$, $(-22.43 \pm 1.59)\text{\%e}$, and $(-21.42 \pm 1.90)\text{\%e}$ for surface, middle, and bottom layers, respectively。The $\delta^{13}\text{C}_{\text{SOC}}$ values in the northeastern region of Chongqing tended to be more negative, whereas those in central Chongqing were less negative。Paddy fields showed the most negative values of $\delta^{13}\text{C}_{\text{SOC}}$, followed by rice-upland rotating fields and upland fields, with the average being $(-25.32 \pm 0.93)\text{\%e}$, $(-23.17 \pm 1.37)\text{\%e}$, and $(-24.75 \pm 1.28)\text{\%e}$ for the surface layers, respectively。For different soil types, the $\delta^{13}\text{C}$ values in the surface layers were in the order of paddy soil < alluvial soil < purplish soil < calcareous soil < yellow soil。According to the regression-tree analysis, the crop types predominantly influenced the variation in $\delta^{13}\text{C}_{\text{SOC}}$ in surface soils, and soil types mainly affected that in the middle- and bottom-layer soils。Other factors, such as soil properties (TN, SOC, and pH) and meteorological conditions (precipitation and air temperature) played only minor roles in the variation of $\delta^{13}\text{C}_{\text{SOC}}$ 。In short, the stable isotope composition of organic carbon in the surface soils was primarily controlled by the input carbon source, whereas that in the deeper layers was closely linked with carbon cycling processes within the soils。

Key words: farmland; soil organic carbon (SOC); stable isotope composition; spatial variation; controlling factors

土壤作为大气 CO_2 的天然碳汇, 其有机碳储量约为 1500 Pg, 高于植物碳库(约 550 Pg)和大气碳库(约 850 Pg)的总和^[1~3]。土壤有机碳(SOC)的累积量和周转速率因而显著影响着全球碳收支和气候系统的稳定性^[4~6]。植物残体是土壤有机碳最重要的来源, 具有不同光合途径的植物在光合作用过程中存在同位素的分馏效应, 从而使植株稳定性碳同位素组成($\delta^{13}\text{C}$)差异较大。研究表明, C3 植物的 $\delta^{13}\text{C}$ 值介于 $-20\text{\%e} \sim -35\text{\%e}$ 之间, 均值为 -27\%e ; C4 植物 $\delta^{13}\text{C}$ 值范围为 $-9\text{\%e} \sim -17\text{\%e}$, 均值为 -13\%e ^[7]。因此, 植被凋落物与根系分泌物直接影响着表层土壤的 $\delta^{13}\text{C}$ 值。土壤剖面 $\delta^{13}\text{C}_{\text{SOC}}$ 的分布主要归因于微生物分解有机物时对 ^{12}C 的优先选择, 导致分解底物中 ^{13}C 的富集^[8]。然而, 土壤性质和气象条件等因素的不同, 可能会影响微生物的分解效率, 进而导致土壤

稳定性碳同位素组成不同。

稳定性碳同位素组成能有效反映环境中碳的动态变化, 国内外许多学者常将其用于研究环境中碳的源汇关系, 定量评价碳组分的循环周转, 反演气候变化和人文干扰下土壤的历史变化等^[9,10]。例如, Atere 等^[10]的研究利用 $\delta^{13}\text{C}_{\text{SOC}}$ 研究了不同施肥制度下水稻土有机碳积累和碳流动途径; Wang 等^[11]的研究通过碳同位素质量平衡模型和文献荟萃分析验证了 $\delta^{13}\text{C}_{\text{SOC}}$ 与土壤有机碳周转之间的关系; Li 等^[12]的研究利用 $\delta^{13}\text{C}_{\text{SOC}}$ 揭示了青藏高原东北部冰丘土壤

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有机碳的动态变化。然而,土壤稳定性碳同位素组成是多种因素共同作用下的结果,其分布可能随研究区域和空间尺度的变化而变化^[13~16]。有研究表明,东北典型漫岗坡耕地表层土壤 $\delta^{13}\text{C}$ 值与地形坡度、土壤黏粒含量和pH都显著相关^[17];阿拉尔绿洲 $\delta^{13}\text{C}_{\text{SOC}}$ 受土壤类型和土壤含水量的影响最甚^[4];阿根廷南部的森林生态系统 $\delta^{13}\text{C}_{\text{SOC}}$ 的控制因素是年均降雨量,而不是土壤性质^[18];气温和SOC则是高寒地区 $\delta^{13}\text{C}_{\text{SOC}}$ 的控制因素^[11]。

目前,土壤有机碳稳定性同位素组成在碳循环研究中的应用在日益增加^[19,20],但由于同位素分析仪器精密贵重,分析成本高,故其在大尺度空间下分布的基础数据仍极为匮乏,影响了人们对其的深入探讨。因此,本研究以重庆地区农田为研究对象,拟通过对采自182个典型剖面的546个土壤样品有机碳 $\delta^{13}\text{C}$ 值的测定,探讨不同土壤层次、耕作制度、作物和土壤类型下有机碳稳定性同位素组成的差异和影响因素,从多个角度揭示重庆农田土壤有机碳稳定性同位素的空间分布特征,以期为未来缓解气候变化、农业土壤碳汇管理和实施碳中和战略提供科学基础。

1 材料与方法

1.1 研究区概况

重庆位于中国西南部的长江上游地区,地跨东

经 $105^{\circ}11' \sim 110^{\circ}11'$ 、北纬 $28^{\circ}10' \sim 32^{\circ}13'$,总体地势为东南部和东北部高、中部和西部低,东部海拔一般在 $1000 \sim 2796$ m之间,西部海拔在 $168 \sim 900$ m之间,形成了西部丘陵、中部平行岭谷和东部中低山为主的地貌组合。区域山地占76%,丘陵占22%,河谷平原仅占2%。重庆属亚热带季风性湿润气候,年均气温为 $15 \sim 18^{\circ}\text{C}$,常年降雨量为 $1000 \sim 1450$ mm,年均相对湿度接近80%,属高湿区。由于复杂的地质地貌条件,加上区域内气候和生物等成土因素具有明显的垂直变化和空间差异,土壤多样,分布较广的主要有黄壤、紫色土、水稻土、石灰(岩)土和潮土。全市耕地面积为237.05万hm²,是重庆第二大土地利用类型,主要耕作制度为旱作、水田和水旱轮作,主要种植水稻、玉米、油菜和烟叶等粮食与经济作物。按传统习惯,可将重庆分为渝西(北碚、璧山、大足、合川、江津、南川、綦江、荣昌、铜梁、永川、潼南和万盛)、渝中(巴南、长寿、涪陵、九龙坡、沙坪坝和渝北)、渝东北(城口、垫江、奉节、开县、梁平、万州、忠县、巫溪、巫山和云阳)和渝东南(丰都、彭水、石柱和武隆)。

1.2 样品采集

以土壤类型分布面积比例为基础,选择远离工业区和城区的典型农田作为样地(图1)。于2016年7~8月间农作物收获后人工挖掘土壤剖面,按表层

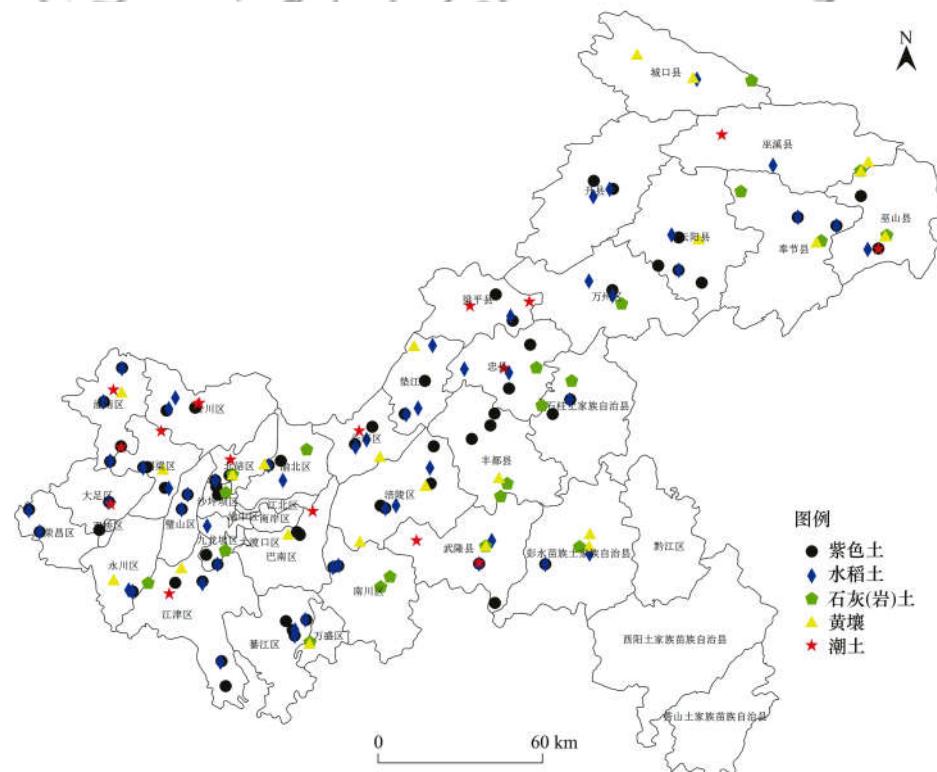


图1 重庆农田采样点分布示意

Fig. 1 Distribution map of farmland sampling sites in Chongqing

(0~30 cm)、中层(30~60 cm)和底层(60~100 cm)进行分层采样。每个剖面同一层次、不同位置分别采集4个样品后归并为该层混合样。每块样地近20年来的耕作制度和作物种植历史则通过向农户口头询问获取。本研究共计182块样地(土壤剖面),546个土样。样地按土壤类型分,包括紫色土68块、水稻土53块、石灰(岩)土21块、黄壤23块和潮土17块;按作物类型分,长期种植C3作物的72块、C4作物12块、C3和C4混种的98块。

1.3 气象参数

采样点的经纬度和海拔高度信息由GPS仪采集记录,年均气温和年均降雨量由覆盖研究区域的59个气象站点的多年(1986~2015年)均值数据经克里格插值后提取获得。

1.4 土样测定

采集的土壤经自然风干后除去石砾和动植物残体等杂质,按四分法取样,依次研磨制备过10和200目筛的样品。

称取过10目筛的土样测定pH值(1:5土水比,雷磁PHS-3E型)。对于pH>6.5的土壤,称取0.500 g过200目筛的样品,加入过量2 mol·L⁻¹的盐酸,反应6 h以去除土壤中的碳酸盐,然后用去离子水洗涤至中性,低温烘干重新磨细后用于SOC含量和同位素组成分析;至于pH<6.5的土样则不经过盐酸处理,直接称样测定^[21]。土样SOC、TN和稳定性碳同位素比均用元素分析仪——连续流同位素比质谱联用系统测定(农业部西南耕地保育重点实验室,Vario Pyro Cube-IsoPrime 100 system)。

稳定性碳同位素比的表示方法: $\delta^{13}\text{C}$ 值表示样品中两种碳同位素比值相对于某一标准对应比值的相对千分差,是描述样品与标准样品相比较时, $\delta^{13}\text{C}$ 天然丰度变异程度的指标,其计算公式为:

$$\delta^{13}\text{C} = \left[\frac{(^{13}\text{C}/^{12}\text{C})_{\text{样品}}}{(^{13}\text{C}/^{12}\text{C})_{\text{标准}}} - 1 \right] \times 1000\text{‰}$$

式中,标准选用美国南卡罗莱纳州白垩系皮狄组地层中的美洲拟箭石(PDB),其 $^{13}\text{C}/^{12}\text{C} = 0.011\,237\,2$ ^[22]。

表1 重庆农田土壤基本情况¹⁾

Table 1 Basic situation of farmland soil in Chongqing

土层	样品数	pH	$\omega(\text{TN})$ /g·kg ⁻¹	$\omega(\text{SOC})$ /g·kg ⁻¹	$\delta^{13}\text{C}_{\text{SOC}}$ /‰
表层	182	6.17 ± 1.26	1.23 ± 0.43	8.69 ± 3.84	-23.63 ± 1.53
中层	182	6.57 ± 1.08	0.84 ± 0.34	4.96 ± 2.57	-22.43 ± 1.59
底层	182	6.64 ± 1.05	0.69 ± 0.27	3.19 ± 1.81	-21.42 ± 1.90

1) 数值为“mean ± SD”

2.2 不同耕作制度下 $\delta^{13}\text{C}_{\text{SOC}}$ 值的差异

如图3所示,水田剖面 $\omega(\text{SOC})$ 最高,表、中和

1.5 统计分析

统计分析皆利用SPSS 22.0软件完成。利用单因素方差分析(One-way ANOVA)结合多重比较法(LSD)对土壤剖面不同层次间或不同耕作制度、作物和土壤类型等变量的均值差异性进行比较(显著性水平为0.05)。为避免自变量之间的多重线性问题,土壤pH、TN、气象条件(气温和降雨量)、SOC、海拔、作物类型和土壤类型等因素对表、中和底层土壤 $\delta^{13}\text{C}$ 值的影响用回归树(regression tree)进行分析。这种方法是通过递归方式构建二叉决策树,并经过剪枝形成优化模型的过程。其优点是结果透明,便于评估自变量的相对重要性^[23]。

2 结果与分析

2.1 SOC 及其 $\delta^{13}\text{C}$ 分布概况

由表1可知,重庆农田剖面土壤 $\delta^{13}\text{C}_{\text{SOC}}$ 范围从上至下分别为-26.81‰~-19.25‰、-26.23‰~-18.78‰和-26.15‰~-15.36‰,平均值分别为(-23.63 ± 1.53)‰、(-22.43 ± 1.59)‰和(-21.42 ± 1.90)‰, $\delta^{13}\text{C}_{\text{SOC}}$ 值随土壤深度增加而逐渐趋正;表、中和底层 $\omega(\text{SOC})$ 均值分别为(8.69 ± 3.84)、(4.96 ± 2.57)和(3.19 ± 1.81) g·kg⁻¹,SOC含量随土壤深度增加而减少。

图2为重庆农田SOC及其 $\delta^{13}\text{C}$ 值的区域分布。就区域而言, $\delta^{13}\text{C}_{\text{SOC}}$ 值整体表现为渝东北<渝西<渝东南<渝中。渝中地区土壤表、中和底层 $\delta^{13}\text{C}_{\text{SOC}}$ 范围分别为-26.19‰~-19.25‰、-25.53‰~-17.87‰和-25.38‰~-16.30‰,平均值分别为(-22.97 ± 1.65)‰、(-21.68 ± 1.79)‰和(-20.59 ± 2.26)‰;渝西剖面从上至下 $\delta^{13}\text{C}_{\text{SOC}}$ 范围分别为-26.42‰~-20.41‰、-26.23‰~-19.65‰和-26.15‰~-15.36‰,平均值分别比渝中同层次偏负0.86‰、1.02‰和0.97‰;渝东北剖面从上至下 $\delta^{13}\text{C}_{\text{SOC}}$ 范围分别为-26.63‰~-20.97‰、-25.87‰~-18.78‰和-24.89‰~-17.07‰,平均值分别比渝中同层次偏负0.96‰、1.04‰和1.29‰。

底层平均值分别为(9.88 ± 1.49)、(5.99 ± 2.12)和(3.55 ± 1.93) g·kg⁻¹,旱地SOC含量最低,各层分

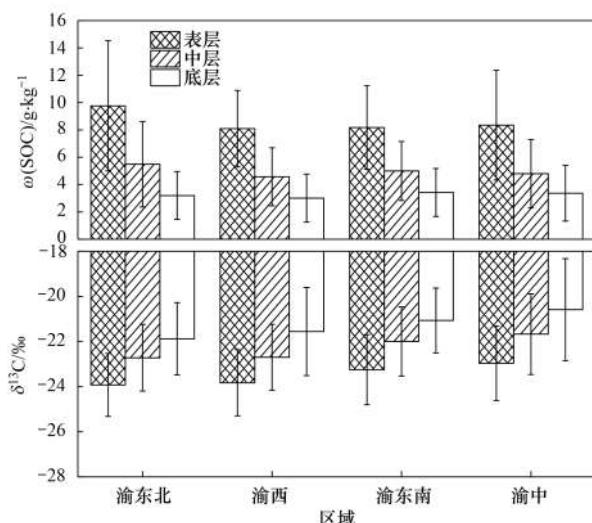
图2 重庆农田SOC及其 $\delta^{13}\text{C}$ 值的区域分布

Fig. 2 Regional distribution of soil organic carbon and its $\delta^{13}\text{C}$ value in Chongqing

别比水田低15%、21%和14%，水旱轮作则居于两者之间。

如图3所示,表层和中层土壤 $\delta^{13}\text{C}$ 值由负趋正且表现为:水田<水旱轮作<旱地;底层土壤 $\delta^{13}\text{C}$ 值由负趋正,依序为:水旱轮作<水田<旱地,整体表现出旱地明显偏正。就表层而言,旱地、水旱轮作和水田 $\delta^{13}\text{C}_{\text{SOC}}$ 值范围分别为 $-26.54\text{\textperthousand} \sim -19.25\text{\textperthousand}$ 、 $-26.81\text{\textperthousand} \sim -21.75\text{\textperthousand}$ 和 $-26.23\text{\textperthousand} \sim -23.87\text{\textperthousand}$,平均值分别为 $(-23.17 \pm 1.37)\text{\textperthousand}$ 、 $(-24.75 \pm 1.28)\text{\textperthousand}$ 和 $(-25.32 \pm 0.93)\text{\textperthousand}$,水旱轮作和水田表层 $\delta^{13}\text{C}_{\text{SOC}}$ 值分别比旱地表层偏负 $1.58\text{\textperthousand}$ 、 $2.15\text{\textperthousand}$ 。

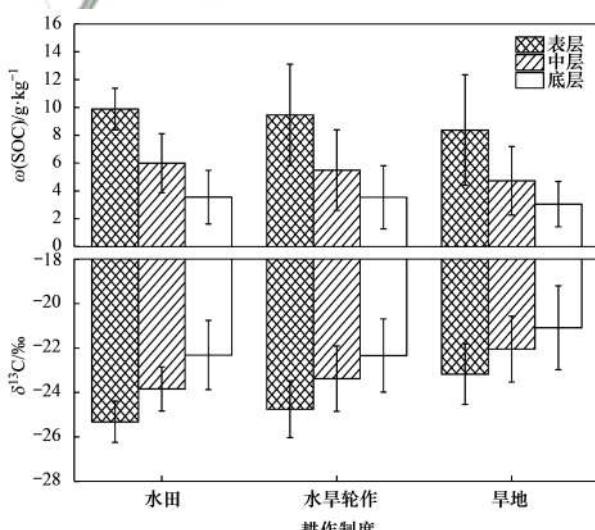
图3 不同耕作制度下SOC及其 $\delta^{13}\text{C}$ 值的分布特征

Fig. 3 Distribution characteristics of soil organic carbon and its $\delta^{13}\text{C}$ value in different farming systems

2.3 不同作物类型下 $\delta^{13}\text{C}_{\text{SOC}}$ 值的差异

图4为不同作物类型下SOC及其 $\delta^{13}\text{C}$ 值的分布。从中可知,表层SOC含量表现为:C3>C3和C4

混种>C4;中层和底层SOC含量均表现为:C3>C4>C3和C4混种。就表层而言,种植C3作物的 $\omega(\text{SOC})$ 在 $2.23 \sim 23.18\text{ g}\cdot\text{kg}^{-1}$ 之间,平均值为 $(9.35 \pm 3.59)\text{ g}\cdot\text{kg}^{-1}$;种植C4作物的 $\omega(\text{SOC})$ 在 $3.50 \sim 11.26\text{ g}\cdot\text{kg}^{-1}$ 之间,平均值为 $(7.18 \pm 2.03)\text{ g}\cdot\text{kg}^{-1}$ 。

如图4所示,剖面 $\delta^{13}\text{C}_{\text{soc}}$ 值由负趋正且表现为:C3<C3和C4混种<C4。以表层为例,种植C3作物的土壤 $\delta^{13}\text{C}$ 值在 $-26.81\text{\textperthousand} \sim -21.35\text{\textperthousand}$ 之间,平均值为 $(-24.56 \pm 1.26)\text{\textperthousand}$;种植C4作物的在 $-24.32\text{\textperthousand} \sim -20.53\text{\textperthousand}$ 之间,平均值比C3作物偏正 $2.03\text{\textperthousand}$;C3和C4混种的在 $-26.54\text{\textperthousand} \sim -19.25\text{\textperthousand}$ 之间,平均值比C3作物偏正 $1.49\text{\textperthousand}$ 。

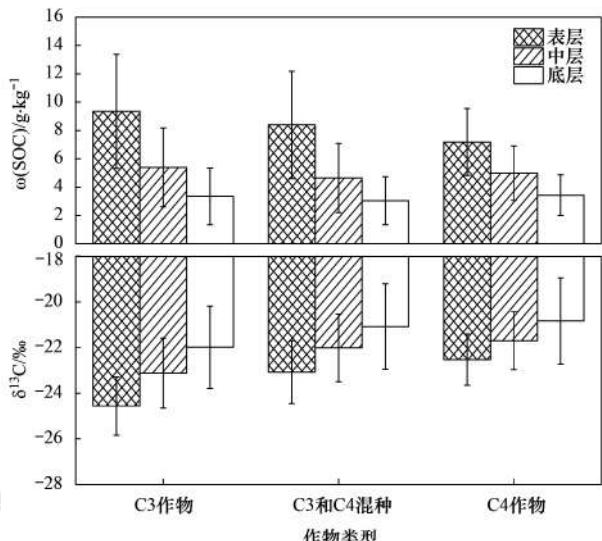
图4 不同作物类型下SOC及其 $\delta^{13}\text{C}$ 值的分布特征

Fig. 4 Distribution characteristics of soil organic carbon and its $\delta^{13}\text{C}$ value in different crop types

2.4 不同土壤类型下 $\delta^{13}\text{C}_{\text{SOC}}$ 值的差异

不同土壤类型下,SOC及其 $\delta^{13}\text{C}$ 值的分布情况如图5所示。SOC含量整体表现为:石灰(岩)土>水稻土>黄壤>潮土>紫色土,石灰(岩)土 $\omega(\text{SOC})$ 最高,其表层最为明显,平均值为 $(12.23 \pm 4.49)\text{ g}\cdot\text{kg}^{-1}$,是紫色土表层的1.79倍。

如图5所示,表层土壤 $\delta^{13}\text{C}_{\text{SOC}}$ 值由负趋正的次序为:水稻土<潮土<紫色土<石灰(岩)土<黄壤,中层依序为:水稻土<潮土<石灰(岩)土<紫色土<黄壤,底层则为:水稻土<石灰(岩)土<潮土<紫色土<黄壤。其中,水稻土表层土壤 $\delta^{13}\text{C}$ 均值为 $(-24.77 \pm 1.23)\text{\textperthousand}$, $\Delta\delta^{13}\text{C}$ (即底层与表层土壤 $\delta^{13}\text{C}$ 之差)平均值为 $2.46\text{\textperthousand}$;紫色土表层土壤 $\delta^{13}\text{C}$ 均值为 $(-23.20 \pm 1.29)\text{\textperthousand}$, $\Delta\delta^{13}\text{C}$ 平均值为 $2.30\text{\textperthousand}$;而黄壤 $\delta^{13}\text{C}$ 值总体上在所有土壤中最高,其表层均值为 $(-22.61 \pm 1.29)\text{\textperthousand}$, $\Delta\delta^{13}\text{C}$ 平均值为

2.11‰。整体表现出黄壤和紫色土 $\delta^{13}\text{C}_{\text{SOC}}$ 偏正程度高于水稻土。

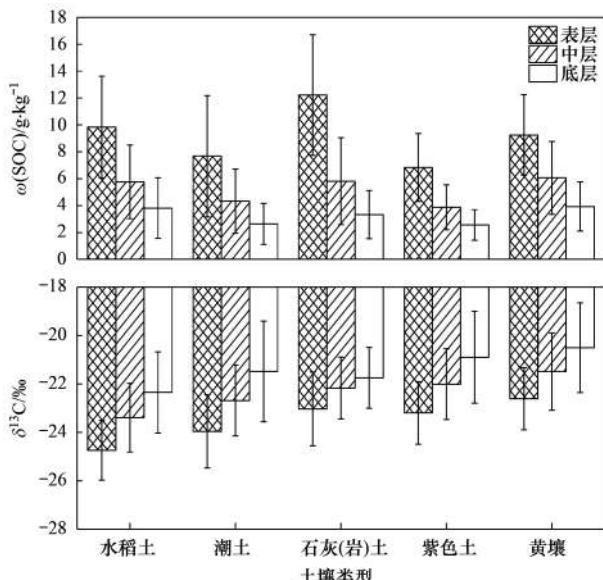


图5 不同土壤类型下SOC及其 $\delta^{13}\text{C}$ 值的分布特征

Fig. 5 Distribution characteristics of soil organic carbon and its $\delta^{13}\text{C}$ value in different soil types

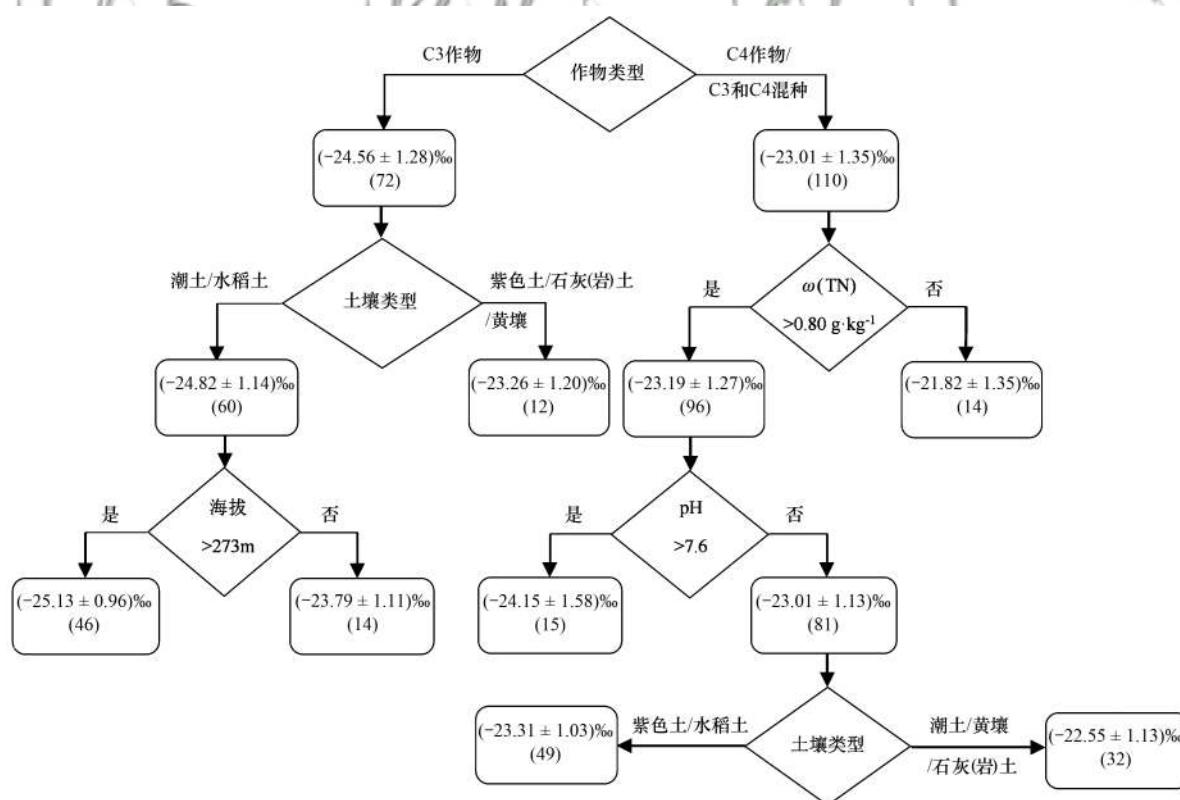
2.5 回归树

回归树分析表明,作物、土壤类型、TN、海拔和pH皆可影响表层土壤的 $\delta^{13}\text{C}_{\text{SOC}}$ 值,其中作物类型

是首要因素(图6)。当种植作物为C3类型,土壤类型为潮土或水稻土,且海拔>273m时, $\delta^{13}\text{C}_{\text{SOC}}$ 值偏负程度最高,均值为(-25.13 ± 0.96)‰;当作物为C4或C3和C4作物混种,且 $\text{TN} \leq 0.80 \text{ g} \cdot \text{kg}^{-1}$ 时, $\delta^{13}\text{C}_{\text{SOC}}$ 值偏正程度最高,为(-21.82 ± 1.35)‰。

与表层土壤不同,中层土壤 $\delta^{13}\text{C}_{\text{SOC}}$ 值的主控因素为土壤类型,其他因素如TN、降雨量和海拔也发挥着重要作用(图7)。当潮土或水稻土的 $\omega(\text{TN}) > 0.89 \text{ g} \cdot \text{kg}^{-1}$,且年均降雨量>1111.0 mm时, $\delta^{13}\text{C}_{\text{SOC}}$ 值偏负,为(-24.54 ± 0.82)‰;当土壤类型为紫色土、黄壤或石灰(岩)土,降雨量介于1053.5~1162.5 mm, $\omega(\text{TN}) \leq 0.89 \text{ g} \cdot \text{kg}^{-1}$,且海拔≤359.5 m时, $\delta^{13}\text{C}_{\text{SOC}}$ 偏正,其值为(-19.95 ± 1.28)‰,比最负值偏负4.59‰。

对于底层土壤 $\delta^{13}\text{C}_{\text{SOC}}$ 值的主控因素仍然是土壤类型,气温、SOC和pH则为次要因素(图8)。当土壤类型为石灰(岩)土或水稻土, $\omega(\text{SOC}) > 6.26 \text{ g} \cdot \text{kg}^{-1}$ 时, $\delta^{13}\text{C}_{\text{SOC}}$ 值偏负,均值为(-23.88 ± 1.20)‰;当土壤类型为紫色土、黄壤或潮土,气温>17.6℃, $\omega(\text{SOC}) \leq 2.47 \text{ g} \cdot \text{kg}^{-1}$ 时, $\delta^{13}\text{C}_{\text{SOC}}$ 值偏正程度最高,比最负值偏负4.98‰。



$\delta^{13}\text{C}_{\text{SOC}}$ 值为“mean ± SD”,括号中为样本数,下同

图6 表层土壤 $\delta^{13}\text{C}$ 值回归树

Fig. 6 Regression tree for top-layer soils

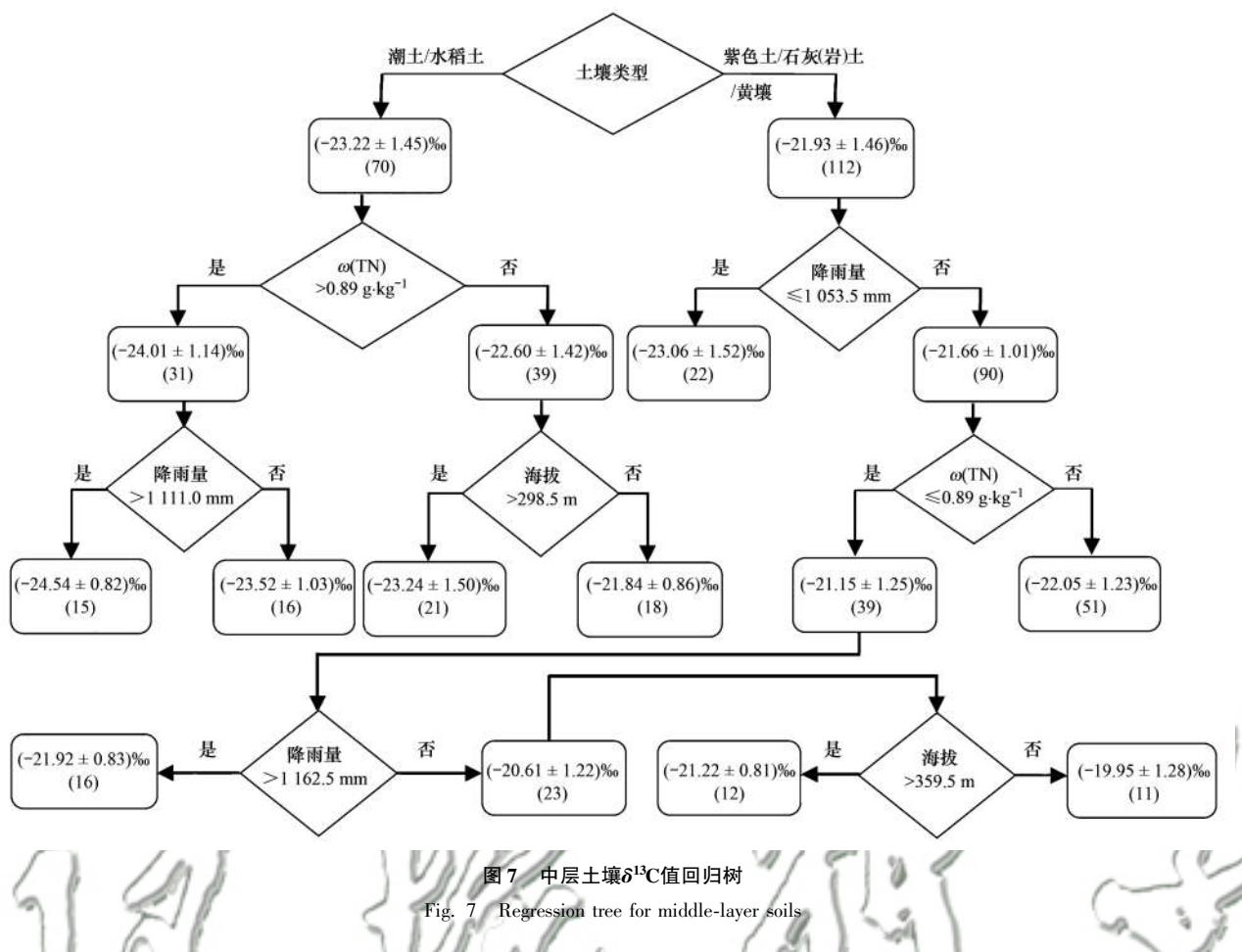


图7 中层土壤 $\delta^{13}\text{C}$ 值回归树
Fig. 7 Regression tree for middle-layer soils

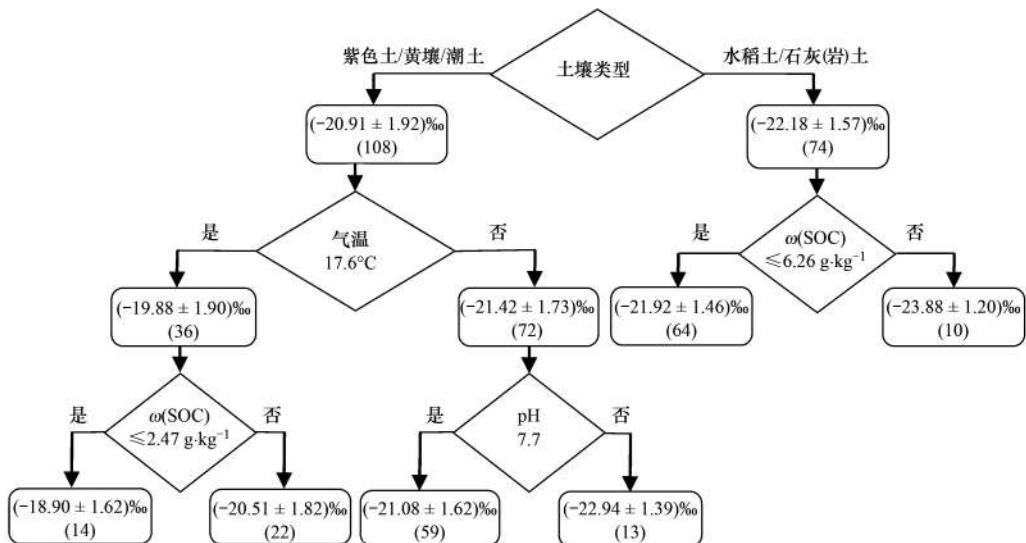


图8 底层土壤 $\delta^{13}\text{C}$ 值回归树
Fig. 8 Regression tree for bottom-layer soils

3 讨论

3.1 $\delta^{13}\text{C}_{\text{SOC}}$ 值空间分布特征

重庆农田表层土壤 $\delta^{13}\text{C}$ 平均值为 -23.63\% , 偏负于华北平原地区高产农田(表层为 -21\% 左右)^[10]和东北松花江流域坡耕地(表层为 -19\% 左右)^[17], 偏正于广东省西南部森林土壤(表层为

-27.31\%)^[24]; 中层和底层土壤的 $\delta^{13}\text{C}$ 平均值分别为 -22.43\% 和 -21.42\% 。与已有的研究结果一样^[25~28], 本研究也发现 $\delta^{13}\text{C}_{\text{SOC}}$ 值均随土壤深度加深而趋正。这可能与土壤表层有机碳向下淋溶迁移过程中, 发生的有机碳分解转化和同位素分馏作用有关^[29]。在SOC分解过程中, 微生物倾向于利用较轻的 ^{12}C 组分, ^{13}C 重组则在分解底物中得以积累^[19],

故 $\delta^{13}\text{C}_{\text{SOC}}$ 值随剖面加深逐渐趋正。除此之外, Suess 效应也可能发挥了重要作用, 即化石燃料的燃烧导致大气 CO_2 的 $\delta^{13}\text{C}$ 值比 200 年前偏负了约 1.99‰, 使得地表植被 $\delta^{13}\text{C}$ 也随着偏负, 从而导致土壤上层新输入有机碳 $\delta^{13}\text{C}$ 值与下层古老有机碳 $\delta^{13}\text{C}$ 值相比更加趋负^[21]。

就重庆不同区域而言, 渝中 $\delta^{13}\text{C}_{\text{SOC}}$ 值偏正, 渝西和渝东北则偏负。渝中地区海拔较低, 气温较高, 且为重庆主要社会经济活动区域, 土地集约化程度较高^[30], 导致有机质循环周转速度较快, 因此 $\delta^{13}\text{C}_{\text{SOC}}$ 值整体表现为偏正。渝西地区广泛分布着稻田, 水田样品采集多源于此区域。有研究表明, 水田有机碳周转速度低于旱地, 具有更高的有机碳储量和封存潜力^[31,32], 故 $\delta^{13}\text{C}_{\text{SOC}}$ 值较偏负。渝东北属于高海拔地区, 气温较低且降雨量较大, 制约着有机碳的循环周转^[33], 而环境中分布的植物大多为 C3 类型, 因此土壤 $\delta^{13}\text{C}$ 值整体偏负。

3.2 环境因子对 $\delta^{13}\text{C}_{\text{SOC}}$ 组成的影响

通过回归树分析, 本研究发现不同土层的影响因素重要性排序和划分阈值均有差异。表层土壤(图 6)的首要影响因素为作物类型, 中层(图 7)和底层(图 8)均为土壤类型, 这一结果与 Jia 等^[34]的研究结果相似。作物类型是控制表层 $\delta^{13}\text{C}_{\text{SOC}}$ 值最重要的环境变量, 其凋落物和根系分泌物直接影响着表层土壤 $\delta^{13}\text{C}$ 值, 这归因于植物本身的 $\delta^{13}\text{C}$ 值差异, C3 植物 $\delta^{13}\text{C}$ 值较 C4 相比更趋向负值, 因此种植 C3 作物的土壤表层 $\delta^{13}\text{C}$ 值更趋负^[8, 34]。当种植作物为 C3 类型时, 海拔较高的潮土和水稻土的 $\delta^{13}\text{C}_{\text{SOC}}$ 值更趋负值。这可能是由于海拔越高, 气温越低, 降雨量越大, 导致土壤中有机碳分解速率降低^[33]。此外, 潮土和水稻土均具有较高的含水率, 抑制着大多数微生物的活性^[4,35]。

然而, 在中层和底层中, 土壤类型则为最主要的控制因素, 土壤类型的影响可能与有机质分解过程中同位素分馏有关, 不同土壤类型间微生物群落组成和理化性质有很大差异, 这就导致有机质在不同土壤环境中的分解速度和分解程度差异显著^[36]。中层和底层土壤的主控因素虽均为土壤类型, 但具体划分也有所差异。中层土壤中, 潮土和水稻土的 $\delta^{13}\text{C}_{\text{SOC}}$ 值偏负, 而底层则将石灰(岩)土和水稻土划分为一类。这可能由于石灰(岩)土剖面有机碳同位素组成随土壤深度加深变化轻微, 底层与表层之差仅为 1.28‰(图 5), 因此随着其他类型土壤同位素分馏效应的增强, 石灰(岩)土底层 $\delta^{13}\text{C}_{\text{SOC}}$ 值则较偏负, 已有研究发现了同样的规律^[37~40]。

除了作物类型和土壤类型等主要因素的影响,

土壤 $\delta^{13}\text{C}$ 值还受 TN、pH、海拔、气温和降雨量的影响(图 6~8)。氮是控制陆地生态系统的关键元素之一, 较高的氮含量通过影响微生物活动, 从而影响 SOC 循环周转速度, 利于有机碳的积累, 导致 $\delta^{13}\text{C}_{\text{SOC}}$ 值更趋负^[41,42]。有研究表明, 弱酸环境有助于土壤微生物群落多样性的提高, 随着微生物活动增强, SOC 分解速率加快, 使 $\delta^{13}\text{C}_{\text{SOC}}$ 值更趋正^[43,44]。此外, 海拔的变化也会造成气候(气温和降雨量)的差异, 从而影响 SOC 的矿化速率和植被的分布与生长。Papathodorou 等^[45]的研究发现气温的升高会增加微生物种类、数量和活性, 从而在一定程度上加速 ^{12}C 的分解, 导致 ^{13}C 的富集。

4 结论

(1) 重庆农田所有剖面 $\delta^{13}\text{C}_{\text{SOC}}$ 值随深度加深逐渐趋正, 表、中和底层均值分别为(-23.63 ± 1.53)‰、(-22.43 ± 1.59)‰ 和 (-21.42 ± 1.90)‰。其中, 渝东北 $\delta^{13}\text{C}_{\text{SOC}}$ 值偏负程度最高, 渝中 $\delta^{13}\text{C}_{\text{SOC}}$ 值偏正程度最高。

(2) 不同耕作制度下, 旱地 $\delta^{13}\text{C}_{\text{SOC}}$ 值明显偏正, 水田偏负, 水旱轮作则居于两者之间, 三者表层土壤 $\delta^{13}\text{C}$ 均值分别为(-23.17 ± 1.37)‰、(-25.32 ± 0.93)‰和(-24.75 ± 1.28)‰。与种植 C4 及 C3、C4 作物混种相比, 种植 C3 作物的土壤 $\delta^{13}\text{C}_{\text{SOC}}$ 值明显偏负; 不同类型土壤表层 $\delta^{13}\text{C}_{\text{SOC}}$ 值依序为: 水稻土 < 潮土 < 紫色土 < 石灰(岩)土 < 黄壤。

(3) 影响土壤有机碳稳定性碳同位素空间分布的主控因素有作物类型、土壤类型、TN、海拔、pH、气温和降雨量, 其中关键因素为作物类型(表层)和土壤类型(中层和底层)。

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