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目 次

南京北郊冬季挥发性有机物来源解析及苯系物健康评估 ··················· 张玉欣,安俊琳,林旭,王俊秀,师远哲,刘静达南京冬季重污染过程中黑碳气溶胶的混合态及粒径分布····································	(1)
基于 GAM 模型分析影响因素交互作用对 PM _{2.5} 浓度变化的影响 ····································	£(13) £(22) £(33)
画江口短叶茳芏+芦苇沼泽湿地大、小潮日土壤间隙水溶解性 CH ₄ 与 CO ₂ 浓度日动态	(41)
河南鸡冠洞 CO, 季节和昼夜变化特征及影响因子比较 ··· 张萍,杨琰,孙喆,梁沙,张娜,田宁,李建仓,凌新有,张志钦 蠡湖表层沉积物荧光溶解性有机质(FDOM)荧光光谱特征 ············· 陈俊伊,王书航,姜霞,黄晓峰,赵丽 滇池沉积物有机质沉积特征与来源解析 ·············· 韩秀秀,黄晓虎,余丽燕,杨浩,黄昌春,黄涛,余艳红,罗玉紫色土小流域浅层井水中胶体颗粒的季节变化 ····································	(60) (70) (78) (87) (95)
基于 LDI 的土地利用类型与亚地尔质的相关性:以亦州太闹三田岛国家建地公园为例 杨朝辉,苏群,陈志辉,白俊武,钱新强,张志敏基于分位数回归的洱海藻类对氮、磷及水温的响应特征 陈小华,李小平,钱晓雍,胡双庆深水型水库藻形态功能组(MBFG)的季节演替特征 杨毓,卢金锁,张颖三峡库区水体中可溶性 C、N 变化及影响因素 范志伟,郝庆菊,黄哲,柴雪思,江长胜三峡库区古夫河小流域氮磷排放特征 华玲玲,李文超,翟丽梅,崔超,刘宏斌,任天志,张富林,雷秋良黄土高原坝系流域干湿季交替下氮输出特征及其源解析:以羊圈沟为例	
5 种沉水植物的氮、磷吸收和水质净化能力比较····································	(156)
大冶湖表层沉积物-水中多环芳烃的分布、来源及风险评价	(102)
pH 和络合剂对五价锑在水钢锰矿和水铁矿表面吸附行为的影响	(180) (188) (195) (201)
一一一一一一一一一一一一一一一一一一一一一一一一一一一一一一一一一一一一一	(212) (220) (229) (238)
硫自养填充床反应器降解水中高浓度高氯酸盐的特性及菌群分析	(247) (253) (260) (269) (276) (283) (294) (301) (309) (318)
不同温度制备的生物质炭对土壤有机碳及其组分的影响:对土壤活性有机碳的影响 ····································	(327)
水分管理模式与土壤 Eh 值对水稻 Cd 迁移与累积的影响	(333)
花、镉单一及复合污染胁迫下土壤生态功能稳定性的影响机制 ········· 陈成瑶,杨惠子,李敏,牛晓丛,苏雨轩,张园生物炭、蒙脱石及其混合添加对复合污染土壤中重金属形态的影响 ····· 高瑞丽,唐茂,付庆灵,郭光光,李晓,胡红青某电镀厂六价铬污染土壤还原稳定化试剂筛选与过程监测········· 李培中,吕晓健,王海见,杨苏才,魏文侠,宋云名对, SPA 15 颗粒对(√/ Ⅱ) 的吸附缝合及其对土壤(√/ Ⅱ) 的像复数力	(352) F(361) F(368)
少化 3BA-13 枫起 MCd(ll / lh / lb / lb / lb / lb / lb / lb /	(3/4) (382) (389) (399) (405) (412)

黄土高原坝系流域干湿季交替下氮输出特征及其源解析:以羊圈沟为例

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摘要:本研究通过对干湿季氮湿沉降过程、降雨过程及基流过程进行动态监测,分析干湿季交替下降雨及基流过程的各形态氮浓度变化,探讨黄土高原坝系流域氮湿沉降对流域氮输出的影响并利用同位素方法解析氮输出来源.结果表明,2015年湿季(7、8月)共11场降雨,产生的氮湿沉降负荷约达814.18 kg,氮沉降通量约为4.26 kg·hm⁻²;干季(9月)共3场降雨,产生的湿沉降负荷约达155.58 kg,氮沉降通量为0.83 kg·hm⁻²,呈现出极大的季节变异性.通过对其中4场降雨过程进行动态分析发现,不同降雨强度对水体氮输出过程影响不同,4场降雨对流域水体的氮贡献量为16.94 kg.降雨径流氮输出负荷占流域水体氮输出负荷的比率为14.45%~64.84%,说明降雨对流域水体氮输出贡献很大.羊圈沟坝系流域δ¹⁵N变化范围较大,为-0.844% ~12.791% ο,δ¹⁸O值在8.166% ~15.115% 范围内波动.

关键词:黄土高原; 坝系流域; 氮湿沉降; 输出负荷; 同位素

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Characteristics of Nitrogen Transport and Its Source Trace in Loess Plateau's Dam Watershed in Alternating Wet and Dry Seasons: A Case Study of Yangjuangou Watershed

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Abstract: In this study, we wanted to explore the impacts of N wet deposition on N export and trace the N source by isotopic method through monitoring N wet deposition in Loess Plateau's Dam Watershed in alternating wet and dry seasons, through measuring N wet deposition, rainfall-runoff and base flow process and analyzing concentration change of different forms of N. The results showed that there were 11 rainfall events in the 2015 wet season, in which N wet deposition load reached 814. 18 kg and N deposition flux was about 4. 26 kg·hm⁻², while there were three rainfall events in the 2015 dry season, in which N wet deposition load reached 155. 58 kg and N deposition flux was 0. 83 kg·hm⁻², so it presented a great seasonal variability. By collecting the dynamic process of four rainfall events, we found that, different rainfall intensity had different influence on N export process. The contribution of N wet deposition of four rainfall events to N export in watershed was 16. 94 kg. The ratio that N output load of rainfall and river in watershed N output load was 14. 45%-64. 84%, which showed that the contribution of rainfall to watershed N transport was big. The variation range of δ^{15} N in Loess Plateau's Dam Watershed was big, which was -0.844%c-12.791%c, and the δ^{18} O was within the range of 8. 166 ‰-15. 115 ‰ in the dam watershed.

Key words: Loess Plateau; dam watershed; nitrogen wet deposition; output load; isotope

随着社会经济的发展和人类活动的加剧,空气中含氮化合物日益增多^[1,2].有研究表明 1860 年全球人为氮输入量为 16 Tg·a⁻¹,而 2005 年全球人为氮输入量达到 210 Tg·a^{-1[3]}.大气氮沉降增加引发了生物多样性降低、氮素过饱和、土壤酸化和水体富营养化等一系列生态环境问题^[4].河流氮输出对流域的生态环境有着深远影响,近些年频繁报道的湖库、近海水域富营养化便是例证^[5,6].由于近年

来工农业迅速发展,废水中 N 的运载增多,河流中 N 浓度升高,δ¹⁵N值也相对较高^[7]. 为了解析水体中 N 来源,稳定同位素作为一种有效的示踪技术在识别

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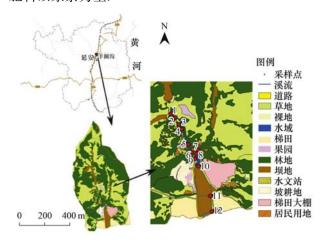
地表水中硝酸盐的来源及迁移转化过程中得到了广 泛的应用^[8]. 1971 年 Kohl 等^[9] 首次利用硝酸盐 N 同位素(15N)评估农田化肥对河流中硝酸盐污染; Johannsen 等[10]报道林地流域水体硝酸盐δ¹⁸O值为 2‰,与实测值相近,硝酸盐主要来源于流域土壤的 硝化作用; 在以农业为主的流域,水体中硝酸盐来 源主要包括化肥及有机肥施用、大气沉降、N。固定 和土壤氮[8];目前,国内 Chen 等[11]研究发现夏季 太湖水体中硝酸盐主要来自大气沉降与生活污水, 冬季北太湖主要来自污水和有机肥,东太湖主要来 自土壤有机氮的硝化作用; 邢萌等[12] 通过对浐河、 涝河硝酸盐δ¹⁵N的研究发现不同来源的硝酸盐 N 同 位素表现出明显的差别,工业排污可能是该河流氮 主要来源. 利用δ¹⁵N和δ¹⁸O同位素结合水化学方法 能较好地分析水体 NO; 污染来源[13].

黄土高原地区是我国西北地区重要的生态屏 障,对我国的生态环境建设以及全球变化的研究,具 有重要作用[14]. 黄土高原坝系流域是主要粮食生 产基地,由于长期以农业耕作为主,较高的垦殖度, 不合理的土地利用导致该地区水土流失严重. 同时 农业耕作中化肥的不合理使用又导致了土壤氮元素 输出的污染变得日益严重[15]. 本文通过对位于陕 西省延安市羊圈沟典型坝系流域进行研究,探讨黄 土高原氮湿沉降季节性动态特征及氮输出特征,揭 示氮湿沉降对坝系流域氮输出的影响,并利用同位 素方法解析氮来源,以期为该流域氮污染排放的预 防和控制及加强流域管理提供科学依据.

1 材料与方法

1.1 研究区概况

本项目主要研究依托中国科学院生态环境研究 中心的黄土高原丘陵沟壑区羊圈沟野外观测站,该 站位于陕西省延安市羊圈沟小流域(36°42′N, 109°31′E),见图 1. 该区域属于典型黄土高原丘陵 沟壑地貌,地形破碎,水土流失严重. 研究流域的海 拔1 050~1 295 m, 总面积为 187. 69 hm², 属于半干 旱干旱大陆性季风气候,气候变化剧烈,年平均气温 为 9. 4 ℃, 7 月多年平均为 22. 9 ℃, 1 月多年平均为 -6.5℃. 降水多集中在7~9月,多年平均降水量 为535 mm. 土壤类型以黄绵土为主,抗蚀性差,土 质疏松,质地均一[16]. 黄绵土是由黄土母质经直接 耕种而形成的一种幼年土壤,其土体疏松、软绵,土 色浅淡. 黄绵土的颗粒组成以细砂粒(0.05~0.25 mm)和粉粒(0.005~0.05 mm)为主,有机质含量耕 地一般在 3~10 g·kg⁻¹之间,草地 10~30 g·kg⁻¹, 腐殖质组成以富里酸为主,保肥能力较弱[17]. 该地 区主要作物有马铃薯和玉米,4月初施用基肥,以农 家肥(羊粪、鸡粪等)、氮肥(碳酸氢铵)和磷肥(过 磷酸钙)为主,6月底到7月初进行一次追肥,此时 肥料以尿素为主.



流域采样点分布及土地利用方式

Fig. 1 Sampling point distribution of watersheds and different land-use types

1.2 研究方法

1.2.1 样点布设

以沿流域内左支沟为典型,根据不同土地利用 方式均匀布设采样点,共计12个,供常规采样,其中 样点6号、12号处为水文站所在地,样点1以上沟 段平时无水流,见表1. 为了直观地反映不同土地利 用方式流域氮浓度变化趋势,采样点划分成4段,分 别是 A、B、C、D 段.

表 1 流域采样点地理位置及描述 Table 1 Location and description of sampling points

1111直	术件点	土地利用力式	地埋似直	奋 社
	1	草地	常流源头	无居民居住
A 段	2	草地	大转弯	附近种植玉米
	3	草地	核桃树	放羊
	4	林地	核桃树 + 石磨	
B段	5	林地	左支沟口前 50 m	居民区
	6	林地	左支沟水文站	

工事利用子子

	-	1	110 010 00 1. 2.	75/H F 4/H III
A 段	2	草地	大转弯	附近种植玉米
	3	草地	核桃树	放羊
	4	林地	核桃树 + 石磨	
B段	5	林地	左支沟口前 50 m	居民区
	6	林地	左支沟水文站	
	7	水塘	右支沟水文站	不在河流里
C 段	8	水塘	水塘上	
C AX	9	水塘	水塘中	
	10	水塘	水塘下	
D段	11	耕地	耕地排水口	试验站
D 4X	12	耕地	淤地坝水文站	

1.2.2 采样方法

(1) 常规采样 2015 年 7~9 月进行采样,每月 10 日、20 日采集样品,其中7、8 月为湿季,共采集

(1)

到24个常规样,9月为干季,采集到23个常规样. 采样时将采样瓶置于水体剖面中部,沟道较浅处直接用瓶口不接触放置沟道底部使水流流进采样瓶, 防止底部沉积物进入瓶中.

- (2)自动采样 2015年7~9月进行采样,其中湿季采集到202个过程样,干季采集到48个过程样.在6号和12号样点采用ISCO6712全尺寸便携式水质自动采样器采集水样. 当明渠中水位或者流速达到所设定值时(水位0.06 m,流速0.001 m·s⁻¹),触发采样器启动采集水样,采至24个样则结束采样.
- (3)湿沉降采样 2015年7~9月进行采样,其中湿季采集到13场降雨,干季采集到4场降雨,共采集到17场降雨. 其中湿季以小雨居多,干季以暴雨居多(日降雨量小于10 mm 为小雨,10~25 mm为中雨,25~50 mm为大雨,50 mm以上为暴雨). 在流域范围内选择3个点位(左支沟、村支部、试验站站顶)为放置雨水采集器收集每场降雨雨水样品,待降雨停止后回收雨水样,见图2.
- (4)稳定同位素采样 按月份在所布设的 12 个采样点分别采集 2 L 水样,相应采集不同季节雨 水样品 2 L,用于稳定同位素检测(由于羊圈沟 7、8 月为雨季,故采样过程在 9 月结束,本研究只将 7~ 9 月的样品作为研究对象).
- (5)试验流量资料 通过 ISCO 6712 全尺寸便携式水质自动采样器实时监测水位、流量及流速的变化.



图 2 雨水采样点分布示意

Fig. 2 Distribution of rain sampling points

1.2.3 分析方法

取适量水样过径 0.45 μm 有机微孔滤膜(经过 80℃水浴 12 h 处理),4℃冷藏保存,一部分通过连 续流动注射分析仪(法国 Futura 型号)测定溶解性

总氮(DTN)、硝态氮(NO_3^--N)、氨氮(NH_4^+-N),另一部分用于 $\delta^{15}N-NO_3$ 的测定. 另取适量水样通过国家标准碱性过硫酸钾的消解方法处理, 4° 冷藏保存,通过连续流动注射分析仪测定 TN.

溶解性有机氮(DON) === DTN - (NO₃ + NH₄ +)

同位素样品采用改进的阴离子交换树脂法进行 处理. 根据NO; -N浓度,取一定体积的水样,通过阴 离子交换树脂柱(Bio-Rad AG1-X8 型树脂)进行离 子交换. 取 8 mL 3 mol·L-1盐酸洗脱吸附在树脂柱 上的 NO,, 向洗脱液中逐次加入 Ag,O,每次加入约 1 g进行反应,共加入约3.3 g Ag,O. 最后用 pH 试 纸检验,pH值要在5.5~6.0之间. 用过滤方法除 去 AgCl 沉淀,将含有 AgNO。的滤液收集在容积为 50 mL 的烧杯中进行冷冻干燥,将冷冻干燥后得到 的 AgNO。 样品用去离子水溶解后转移入尖底离心 管中,再次进行冷冻干燥,使样品均匀地浓缩至较 小体积,最后将 AgNO, 样品用适量的去离子水重 新溶解,把溶解后的溶液转移到5 mm×9 mm 的银 杯中,将银杯放入特制的铝制模具中. 将模具部分 浸入到液氮中,直到 AgNO。溶液冷冻. 将银杯上 部合上,同时用胶模封住模具,进行冷冻干燥. 最 后按照常规方法将银杯压褶,采用 MAT253 稳定 同位素质谱仪(德国 Finnigan 公司)进行质谱分 析[12].

1.3 数据分析

1.3.1 氮湿沉降通量的计算方法^[18] 采用式(1)对氮湿沉降通量进行计算:

$$F = \sum \left(\frac{p \times c}{100} \right)$$

式中,F 为沉降通量($kg \cdot hm^{-2}$),p 为逐月降雨量(mm),c 为雨水中总氮浓度($mg \cdot L^{-1}$).

1.3.2 地表径流量及输出负荷的计算方法^[18] 地表径流量的计算公式:

$$x = \int_{0}^{t} q_{i}(t) dt \approx \sum_{i=1}^{n-1} Q_{i} \Delta t_{i}$$
$$= \sum_{i=1}^{n-1} \frac{q_{i} + q_{i+1}}{2}$$
(2)

输出负荷计算方法:

$$y_{i} = \int_{0}^{t} c_{i}(t) q_{i}(t) dt \approx \sum_{i=1}^{n-1} c_{i} Q_{i} \Delta t_{i}$$

$$= \sum_{i=1}^{n-1} \Delta t_{i} \frac{c_{i} + c_{i+1}}{2} \times \frac{q_{i} + q_{i+1}}{2}$$
(3)

式中,x 为径流量(\mathbf{m}^3); q_i 为样本 i 在监测时的流量($\mathbf{m}^3 \cdot \mathbf{s}^{-1}$); γ_i 为第 j 种污染物的排放负荷(\mathbf{g}); c_i

为 t 时刻径流中第 j 种污染物的浓度($\operatorname{mg} \cdot \operatorname{L}^{-1}$); q_t 为 t 时刻的流量($\operatorname{m}^3 \cdot \operatorname{s}^{-1}$); c_i 为第 j 种污染物在样本 i 监测时的浓度($\operatorname{mg} \cdot \operatorname{L}^{-1}$).

1.3.3 监测点流域输出负荷^[19]

流域氮平均浓度的计算公式为:

$$c_{N} = \sum c_{i} \times w_{i} \tag{4}$$

式中, c_N 为流域氮平均浓度($mg \cdot L^{-1}$); c_i 为在 i 时刻流域氮浓度($mg \cdot L^{-1}$); w_i 为按径流量大小取得的加权系数.

监测点流域输出量的计算公式为:

$$P = c_{N} \cdot Q \tag{5}$$

式中,P 为监测点氮的输出量(g); c_N 为流域氮平均浓度($mg \cdot L^{-1}$); Q 为流域总径流量(m^3).

1.3.4 流域氮沉降对水体氮贡献量的估算[20]

利用径流输出系数(表2)计算氮湿沉降对流域 水体氮的贡献,计算公式为:

$$M = c \cdot A \cdot q \tag{6}$$

式中,M 为降雨输入的氮量(g); c 为雨水氮浓度(g·m⁻³); A 为降雨量(m³); q 为径流系数.

表 2 径流输出系数[21,22]

Table 2 Export coefficient of runoff

土地利用类型	林地	草地	耕地	水体
径流系数	0. 023 ~ 0. 03	0. 032 ~ 0. 04	0. 036 ~ 0. 101	1

2 结果与分析

2.1 流域氮湿沉降干湿季特征

本研究在雨季一共采集 14 场降雨. 图 3 为 7 ~ 9 月雨水中各形态氮浓度的月平均值变化规律,从中可以看出,7 月硝态氮、溶解性总氮和总氮浓度均为最大,铵态氮浓度最低,其中硝态氮浓度明显超过了铵态氮浓度; 而 9 月硝态氮、溶解性总氮和总氮浓度均为最低,硝态氮浓度接近于铵态氮浓度. 其中 7 月铵态氮浓度处于最低值 0. 42 mg·L⁻¹,8 月铵态氮浓度达最大值 0.91 mg·L⁻¹; 7 月硝态氮浓度达最大,为 4.38 mg·L⁻¹,9 月硝态氮浓度最低为 0.50 mg·L⁻¹. 如图 3 所示,羊圈沟小流域湿季的雨水中硝态氮浓度较高,干季的雨水中硝酸氮浓度较低;湿季的雨水中硝态氮浓度存在较大的差异,干季的雨水中硝态氮浓度和铵态氮浓度

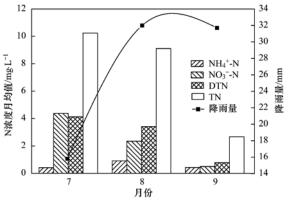


图 3 氮湿沉降浓度月变化及降雨量

Fig. 3 Monthly variation of nitrogen wet deposition concentration and precipitation

几乎接近,说明雨水中的硝态氮和铵态氮可能在外界条件影响下进行了相互转化. 羊圈沟小流域试验区雨水中的主要成分是无机氮,通过对各个月中各场降雨 TN 浓度计算平均值得到平均沉降及平均沉降通量,计算各月中每场雨 TN 浓度之和,得到总沉降负荷列于表 3. 从中可知,7~9月分别采集到 4、7和 3场降雨,其中 8月氮沉降负荷比其他两月稍高,可能是由降雨频繁引起的. 湿季的雨水氮总沉降负荷和氮平均沉降通量远高于干季,可能与降雨量和降雨次数有关.

表 3 氮平均沉降负荷及沉降通量的月变化

Table 3 Mean monthly variation of nitrogen deposition

fluxes and associated load 项目 9月 7月 8月 平均沉降负荷/kg 67.77 77.58 51.86 平均沉降通量/kg·hm⁻² 1.46 2.80 0.83 总沉降负荷/kg 271.10 543.08 155. 58

2.2 流域氮输出特征

从图 4 中可以看出 DTN、TN 和NO₃-N浓度变化规律相似:DTN、TN 和NO₃-N浓度呈现随径流流向先增高后降低的趋势. 在流域的 A 段硝酸盐氮浓度稍高. 该区域无居民生活,几乎无需考虑人类生活垃圾、生活污水造成的人为影响; 土地利用方式为草地,河流附近种植大量玉米,可能施用无机氮肥,农田用水和雨水汇入河流造成硝酸盐氮浓度升高. 3 号采样点有居民放羊,动物粪便中含有尿素和氨氮,随着雨水的冲刷淋溶进入河流中,使得 3 号采样点硝酸盐氮浓度最高.

在流域的 B 段,硝酸盐氮浓度变化十分明显,

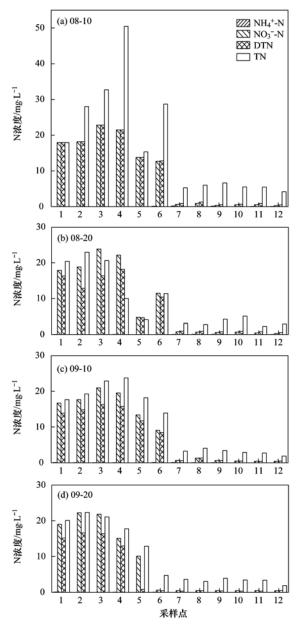


图 4 流域不同形态氮干湿季变化特征

Fig. 4 Concentration variation of different forms of nitrogen in alternating dry and wet seasons

呈现陡降的趋势,该区域土地利用方式为林地,4号采样点附近种植核桃树,施用农肥,使得硝酸盐氮浓度略微降低.5和6号采样点为居民生活区,产生的生活污水、生活垃圾和人类粪便中含有硝酸盐氮和氨氮,故硝酸盐氮浓度并未降低到河流本身氮浓度.

在流域的 C 段,水塘中存在大量水生植物,水体中水生植物吸收硝态氮比铵态氮更容易,水生植物对硝酸盐的吸收会导致硝酸盐浓度降低,同时,水塘水流速度减缓,水体处于静止缺氧状态,微生物在缺氧过程中发生反硝化作用,反硝化细菌将硝酸盐氮还原为 N₂、N₂O,可能使硝酸盐浓度降低,故水塘

中硝酸盐浓度本身很低. 7 号采样点为水塘的另一端,8 号采样点为水塘人口,6 和 7 号采样点附近的水体均汇人8 号采样点处,由于水塘人口处存在河流的流动,削弱了微生物的反硝化作用,故8 号采样点的硝酸盐氮浓度高于7 号采样点.

在流域的 D 段, 硝态氮浓度变化基本稳定, 保持在较低范围内. 11 和 12 号采样点处附近建有梯田大棚, 施用有机肥料, 微生物发生硝化反应, 将部分氮有机物转化为硝酸盐, 但是 11 号采样点附近建有试验站, 居民生活污水排入河流中, 生活污水的排放造成了河流淤泥堆积, 构成了厌氧微生物进行反硝化作用的有利条件, 故硝酸盐浓度很低. 12 号采样点的河流流速缓慢, 存在更多淤泥堆积, 导致硝酸盐浓度进一步降低. 另外, 水生植物的吸收也可能导致硝酸盐浓度的下降.

从图 4 中氨氮的浓度分析得知,羊圈沟流域氨 氮浓度在整个流域内变化不大,变化范围为 0.01~ 0.18 mg·L⁻¹. 从 6 号采样点处氨氮浓度开始升高, 可能与居民生活排放的生活污水有一定的关系. 7 号采样点处氨氮浓度出现最高点,该点附近存在垃 圾置放处,生活垃圾中含有氨氮和有机氮,导致峰值 出现在该处. 9 号采样点为钓鱼处,鱼的饲料中有 含氮有机物,微生物进行脱氨作用,使氨氮浓度高于 8号采样点. 11和12号采样点排放生活污水,污水 中含有大量的氨氮和有机氮,造成了氨氮浓度升高 的现象. 从整体上看,硝态氮浓度远高于氨氮浓度, 是由于土壤和水体中的微生物发生了硝化作用,使 得铵态氮转化为硝态氮. 另外 8 月 10 日 TN 浓度远 高于其它 3 次常规采样,最高值甚达 50.45 mg·L⁻¹, 而 DTN 浓度与其它 3 次常规采样相差不 大,推测可能原因是8月10日羊圈沟小流域外界施 加了氮源且氮源含量不稳定.

2.3 同位素解析流域氮输出来源

如果反应环境中没有大量的 NH_4^+ 积累,矿化和硝化作用产生的硝酸盐的 $\delta^{15}N$ 值与初始反应物质的 $\delta^{15}N$ 值一致,具有较小的同位素分馏 $^{[12]}$. 由图 4 可知,样品中 NH_4^+ 浓度均为 0.1 $mg \cdot L^{-1}$ 左右,可以近似地认为所采集水样的同位素分馏不显著, $\delta^{15}N$ 即可表明河流中 NO_3^- 的来源. 由于不同硝酸盐来源的氮同位素之间具有重叠性 $^{[23]}$,而 $\delta^{18}O$ 值往往存在较大差异,所以单纯使用 $\delta^{15}N$ 不能够有效地识别硝酸盐来源,而硝酸盐 $\delta^{18}O$ 值也是一种识别硝酸盐来源的指标,可减少氮同位素在识别硝酸盐来源时的不确定性 $^{[24]}$.

图 5 为 8 月 10 日常规采样点δ¹⁵N和δ¹⁸O值变化 情况. 从中可知,羊圈沟小流域硝酸盐δ¹⁵N无明显变 化规律,变化起伏波动较大,变化范围为 - 0.844‰ ~12.791‰,平均值为 5.974‰, δ¹⁸O值大体呈现逐 渐下降的趋势,在8.166‰~15.115‰范围内波动. 羊圈沟小流域 A、B 段硝酸盐平均氮浓度很高, A 段 δ¹⁵N值在 6.704‰~12.791‰范围内变动, B 段δ¹⁵N 值变化范围在 3.514‰~8.015‰, δ^{15} N受到的影响 较复杂. 根据表 4 可知,该地区δ¹⁵N值处于土壤氮、 大气降水、有机肥料、牲畜粪便和生活污水δ¹5N特 征值范围内,而且采样当天刚下过雨,说明该地区 δ¹⁵N可能受到土壤氮、降雨、人为作用共同影响. 而 监测结果表明δ¹δO值满足硝化作用的同位素特征值 范围,不满足降雨中的硝酸盐δ¹⁸Ο值的范围(25‰~ 75‰),说明降雨中的硝酸盐不是该地区硝酸盐氮 主要来源. 结合该地区种植、放牧和居民生活状 况,可判断 A 段硝酸盐主要来自土壤氮和牲畜粪 便: B 段硝酸盐主要来自生活污水. 许多研究表明, 土壤中的硝化反应是林地流域地表水中硝酸盐的主 要来源[29~31].

表 4 硝酸盐潜在来源的同位素范围[23,25~28]

Table 4 Isotope ranges of potential sources of nitrate

Table + Isoto	pe ranges or potential sou	ices of intrate
硝酸盐来源	$\delta^{15}N/\%o$	$\delta^{18}O/\%$ o
土壤氮	0 ~8	-5 ∼7
大气降水中的硝酸盐	− 10 ~ 9	25 ~ 75
牲畜粪便	10 ~ 20	-5 ∼7
硝酸盐肥料	-5 ∼5	18 ~ 24
铵肥	-5 ∼0	- 5 ∼ 7
有机肥料	2 ~ 30	- 5 ∼ 7
生活污水	4 ~ 19	- 5 ∼ 7
硝化作用	− 12 ~ 29	− 10 ~ 15
吸收作用	-27 ~0	_

流域 C 段池塘内,硝酸盐 $\delta^{15}N$ 分布在 -0.844% ~ 10.345% 范围内, $\delta^{18}O$ 的变化范围为 8.247% ~ 10.97%,处于硝化作用同位素特征值的范围内. 但是 7 和 8 号采样点 NO_3^- -N浓度非常低, $\delta^{15}N$ 值较高,造成这种现象的原因可能是氮的轻同位素 ^{14}N 较 ^{15}N 优先发生氨 (NH_3) 挥发、反硝化作用和植物吸收等过程. 有研究表明,反硝化能引起同位素大范围的分馏 $(-40\% \sim -5\%)^{[8]}$,导致 $\delta^{15}N$ 和 $\delta^{18}O$ 均在残余硝酸盐中富集,且二者以接近 2:1的比例富集 $^{[23]}$.但是羊圈沟流域 C、D 段 $\delta^{15}N$ 与 $\delta^{18}O$ 的比值不存在 2:1的关系(图 6),且它们的富集系数不在 $1.3 \sim 2.1$ 范围内 $^{[25]}$,说明硝酸盐浓度降低而 $\delta^{15}N$ 值升高可能不是反硝化过程导致的,也可能是由于污水和有机

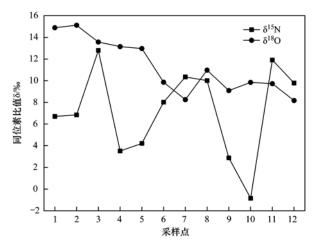


图 5 雨季基流δ¹⁵N及δ¹⁸O值变化

Fig. 5 Variation of $\delta^{15}N$ and $\delta^{18}O$ in base flow in rainy season

肥污染的输入掩盖了反硝化过程的结果. 另外,浮游植物对硝酸盐的吸收同化作用也可能会导致硝酸盐浓度降低. 9 和 10 号采样点硝酸盐δ¹⁵N值骤降至2.88‰和 – 0.844‰,而硝酸盐浓度很低,故推测该地区硝酸盐来源是水体植物吸收过程和硝化反应共同作用.

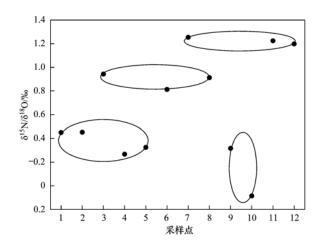


图 6 雨季基流δ¹⁵N/δ¹⁸O值

Fig. 6 The $\delta^{15} N/\delta^{18} O$ of base flow in rainy season

流域 D 段,硝酸盐 δ¹⁵N 值稍高 (9.784‰~11.907‰),δ¹³O均值为 8.944‰. 该地区为居民生活区,有大量生活污水、垃圾排放,并且在附近建有梯田大棚,施用大量的农肥. 对照表 4,该地区的硝酸盐δ¹⁵N处于有机肥料和生活污水δ¹⁵N特征值范围,δ¹³O接近其δ¹³O特征值的上限,推测该地区硝酸盐δ¹⁵N受到人为污染和农业污染的影响. 故 D 段硝酸盐来源主要是有机肥料和生活污水. 综上所述,羊圈沟坝系流域的硝酸盐氮来源主要是土壤氮、有机肥料和生活污水,而大气降水来源较少.

3 讨论

根据流域不同土地利用方式(图1)将果园、梯田大棚、坡耕地均定为耕地,其面积为13.85 hm²,林地面积为92.15 hm²,草地面积70.62 hm²,水域面积为0.44 hm².根据式(6)计算,得出羊圈沟小流域内次降雨氮湿沉降对河流贡献量(图7).从中可以看出,7月采集到4场降雨的氮沉降通过土壤截留之后向流域内河道输入的总氮量为9.32 kg,8月采集到7场降雨氮湿沉降对流域水体氮贡献量为17.89 kg,氮的总输出负荷为129.82 g,降雨径流氮输出负荷与流域水体氮输出负荷比值为64.84%.9月采集到3场降雨氮湿沉降对河流贡献量为5.28 kg,氮的总输出负荷为20.08 g,降雨径流氮输出负荷占流域水体氮输出负荷14.45%.氮贡献量取决于降雨量而与降雨频次无关.地表径流主要与降雨

量有很大关系,卢龙彬等^[32]研究表明,产流必要条件是表层土壤达到饱和且雨强≥7.2 mm·h⁻¹. 黄土高原属于干旱半干旱地区,气温较高,降雨较少,雨强较小时,不足以形成径流,故该地区径流输出负荷很小. 据了解,该流域土地利用方式有林地、耕地等,是影响流域氮人为输入的主要因素,故推测该流域内水体中湿沉降输入的氮量相对于农肥的输入量,其贡献很小.

对 4 场不同强度降雨进行分析,结果见表 5. 从中可知,4 场降雨的总径流量为 13.01 m³,总氮的输出负荷为 94.36 g,其中,可溶性总氮输出负荷占总氮输出量的 46.99%,氨氮和硝态氮的输出负荷分别为 0.44 g 和 39.51 g,分别占可溶性总氮的 1% 和 89.11%.4 场典型降雨对河流贡献量在 0.93~9.4 kg 之间.因此,羊圈沟小流域湿沉降对水体氮的贡献比农业肥料小而且水体中主要以硝态氮的形式输出.

表 5 不同降雨事件下各形态氮的输出负荷

Table 5 Output load of different N forms in different rainfall events

日期(月-日)	降雨量/mm	总径流量/m³	TN/g	DTN/g	NH ₄ -N/g	NO ₃ -N/g
08-02	10	4. 76	33. 04	16. 44	0. 34	13. 60
08-25	4. 2	2. 70	15. 84	10. 29	0. 03	10. 13
09-04	13	0. 77	1.88	1. 52	0. 01	1.34
09-09	4. 8	4. 78	43. 60	16. 09	0.06	14. 44
总计	32	13. 01	94. 36	44. 34	0. 44	39. 51

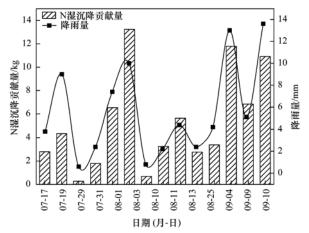


图 7 7~9 月流域内大气氮湿沉降对河流贡献量

Fig. 7 Amount of contribution of the atmosphere N wet deposition in watersheds to the river from July to September

4 结论

(1) 在羊圈沟小流域湿季期间,产生的湿沉降负荷约达 814.18 kg,氮沉降通量约为 4.26 kg·hm⁻²,其中铵态氮沉降通量为 0.26 kg·hm⁻²,硝态氮沉降通量为 0.90 kg·hm⁻²;干季期间,产生的湿沉降负荷约达 155.58 kg,氮沉降通量为 0.83

- $kg \cdot hm^{-2}$,其中铵态氮沉降通量为 0. 11 $kg \cdot hm^{-2}$,硝 态氮沉降通量为 0. 14 $kg \cdot hm^{-2}$,呈现出极大的季节变异性.
- (2)流域内干湿季大气氮湿沉降对河流贡献量范围在5.28~17.89 kg,降雨径流氮输出负荷占河流氮输出负荷的比率为14.45%~64.84%;不同强度的降雨对水体氮输出过程影响不同,水体中氮主要以硝态氮的形式输出.
- (3) 羊圈沟坝系流域 δ^{15} N变化范围较大,为 -0.844%。 ~ 12.791‰, δ^{18} O 值 在 8.166‰ ~ 15.115‰范围内波动,流域的硝酸盐氮来源主要是有机肥料、生活污水和土壤氮,大气降水来源较少. 参考文献·
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HUANJING KEXUE

Environmental Science (monthly)

Vol. 38 No. 1 Jan. 15, 2017

CONTENTS

Size Distribution and Mixing State of Black Carbon Aerosol in Nanjing During a Heavy Winter Pollution Event JIANG Lei, TANG Li-li, PAN Liang-bao, et al. (Interactive Effects of the Influencing Factors on the Changes of PM _{2.5} Concentration Based on GAM Model HE Xiang, LIN Zhen-shan (2) Particle Size Distribution and Diffusion for Simulated Cooking Fume LI Shuang-de, XU Jun-bo, MO Sheng-peng, et al. (3) Greenhouse Gas Fluxes at Water-Air Interface in Small Pond Using Flux-Gradient Method Based on Spectrum Analyzer ZHAO Jia-yu, ZHANG Mi, XIAO Wei, et al. (4) Diurnal Variations of Concentration of Porewater Dissolved CH ₄ and CO ₂ in a Brackish Marsh Dominated by Cyperus malaccensis and Phragmites australis During Neap and Spring Tidal Days in the Minjiang River Estuary TAN Li-shan, YANG Ping, HE Lu-lu, et al. (5) Comparisons Between Seasonal and Diurnal Patterns of Cave Air CO ₂ and Control Factors in Jiguan Cave, Henan Province, China ZHANG Ping, YANG Yan, SUN Zhe, et al. (6) Fluorescence Spectral Characteristics of Fluorescent Dissolved Organic Matter (FDOM) in the Surface Sediments from Lihu Lake CHEN Jun-yi, WANG Shu-hang, JIANG Xia, et al. (7)	(22) (33) (41)
Particle Size Distribution and Diffusion for Simulated Cooking Fume	(33) (41) (52)
Greenhouse Gas Fluxes at Water-Air Interface in Small Pond Using Flux-Gradient Method Based on Spectrum Analyzer	(41)
Diurnal Variations of Concentration of Porewater Dissolved CH ₄ and CO ₂ in a Brackish Marsh Dominated by <i>Cyperus malaccensis</i> and <i>Phragmites australis</i> During Neap and Spring Tidal Days in the Minjiang River Estuary TAN Li-shan, YANG Ping, HE Lu-lu, et al. (5 Comparisons Between Seasonal and Diurnal Patterns of Cave Air CO ₂ and Control Factors in Jiguan Cave, Henan Province, China ZHANG Ping, YANG Yan, SUN Zhe, et al. (6 Fluorescence Spectral Characteristics of Fluorescent Dissolved Organic Matter (FDOM) in the Surface Sediments from Lihu Lake CHEN Jun-yi, WANG Shu-hang, JIANG Xia, et al. (7)	(52)
Days in the Minjiang River Estuary	
Comparisons Between Seasonal and Diurnal Patterns of Cave Air CO ₂ and Control Factors in Jiguan Cave, Henan Province, China	
Fluorescence Spectral Characteristics of Fluorescent Dissolved Organic Matter (FDOM) in the Surface Sediments from Lihu Lake CHEN Jun-yi, WANG Shu-hang, JIANG Xia, et al. (7)	((0)
	(60)
Sedimentary Characteristics and Sources of Organic Matter in Sediments of Dianchi Lake	
Seasonal Variation of Colloid Particles in the Shallow Well Water of a Small Watershed of Purple Soil	87)
Deposition Characteristics of Suspended Solids and the Response of Dissolved Nutrients in Spring in the Western Lakeside of Taihu Lake	
QI Chuang, WANG Guo-xiang, WU Xin-ting, et al. (9	
Correlation Between LDI-based Land Use Types and Water Quality in Sanshan Island of Taihu Lake National Wetland Park, Suzhou	(104)
Response Characteristics of Algal Chlorophyll-a to Nitrogen, Phosphorus and Water Temperature in Lake Erhai Based on Quantile Regression	
CHEN Xiao-hua, LI Xiao-ping, QIAN Xiao-yong, et al. (1	(113)
Seasonal Succession Characteristics of the Morphologically-based Functional Groups (MBFG) in Deep-water Reservoir	(121)
Change and Influencing Factors of Dissolved Carbon and Dissolved Nitrogen in Water of the Three Gorges Reservoir	(129)
Characteristics of Nitrogen and Phosphorus Emissions in the Gufu River Small Watershed of the Three Georges Reservoir Area	(138)
Characteristics of Nitrogen Transport and Its Source Trace in Loess Plateau's Dam Watershed in Alternating Wet and Dry Seasons; A Case Study of Yangjuangou Watershed	
JIA Jun-jie, GAO Yang, CHEN Wei-liang, et al. (1-	
Comparison of Nitrogen and Phosphorus Uptake and Water Purification Ability of Five Submerged Macrophytes	(156)
Pollution Level, Distribution Characteristics and Risk Assessment of 32 PPCPs in Surface Water of Luomahu Lake	
Distribution, Sources and Risk Assessment of the PAHs in the Surface Sediments and Water from the Daye Lake	
Effects of pH and Complexing Agents on Sb(V) Adsorption onto Birnessite and Ferrihydrite Surface	(180)
POC. CHURD P. C. A. C. L. D. C. C. P. LI. L. M. L. WANGOL C. MILLIE J. J. A.	
Effects of UV Radiation on the Coagulation Process of Kaolin and Involved Mechanisms WANG Wen-dong, WANG Chang-xin, LIU Hui, et al. (1)	(188)
Effects of UV Radiation on the Coagulation Process of Raolin and Involved Mechanisms WANG Wen-dong, WANG Chang-xin, LiU Hui, et al. (19) Effect of Dissolved Humic Acid on Thyroid Receptor Antagonistic Activity of Zinc in Aquatic Environment AI Yang, KONG Dong-dong, YU Chang, et al. (19)	
	(195)
Effect of Dissolved Humic Acid on Thyroid Receptor Antagonistic Activity of Zinc in Aquatic Environment	(195) (201)
Effect of Dissolved Humic Acid on Thyroid Receptor Antagonistic Activity of Zinc in Aquatic Environment	(195) (201)
Effect of Dissolved Humic Acid on Thyroid Receptor Antagonistic Activity of Zinc in Aquatic Environment Al Yang, KONG Dong-dong, YU Chang, et al. (19) Optimization and Validation of the Analytical Method to Detect Common Illicit Drugs in Wastewater Cathode Electric Field Enhanced Removal of Nitrobenzene from Aqueous Solution Based on Activated Carbon Fibers (ACF)-Ozone Technique ZHAO Chun, ZHANG Shuai, ZHOU Yu, et al. (2) Effect of PVDF Hollow Fiber Ultrafiltration Membranes Modification with Carbonnanotube on Membrane Fouling Control During Ultrafiltration of Sewage Effluent	(195) (201) (212)
Effect of Dissolved Humic Acid on Thyroid Receptor Antagonistic Activity of Zinc in Aquatic Environment	(195) (201) (212)
Effect of Dissolved Humic Acid on Thyroid Receptor Antagonistic Activity of Zinc in Aquatic Environment	(195) (201) (212) (220) (229)
Effect of Dissolved Humic Acid on Thyroid Receptor Antagonistic Activity of Zinc in Aquatic Environment Al Yang, KONG Dong-dong, YU Chang, et al. (2) Optimization and Validation of the Analytical Method to Detect Common Illicit Drugs in Wastewater Cathode Electric Field Enhanced Removal of Nitrobenzene from Aqueous Solution Based on Activated Carbon Fibers (ACF)-Ozone Technique ZHAO Chun, ZHANG Shuai, ZHOU Yu, et al. (2) Effect of PVDF Hollow Fiber Ultrafiltration Membranes Modification with Carbonnanotube on Membrane Fouling Control During Ultrafiltration of Sewage Effluent WANG Li-ying, SHI Jie, WANG Kai-lun, et al. (2)	(195) (201) (212) (220) (229)
Effect of Dissolved Humic Acid on Thyroid Receptor Antagonistic Activity of Zinc in Aquatic Environment Al Yang, KONG Dong-dong, YU Chang, et al. (19 Optimization and Validation of the Analytical Method to Detect Common Illicit Drugs in Wastewater GAO Ting-ting, DU Peng, XU Ze-qiong, et al. (20 Cathode Electric Field Enhanced Removal of Nitrobenzene from Aqueous Solution Based on Activated Carbon Fibers (ACF)-Ozone Technique ZHAO Chun, ZHANG Shuai, ZHOU Yu, et al. (20 Effect of PVDF Hollow Fiber Ultrafiltration Membranes Modification with Carbonnanotube on Membrane Fouling Control During Ultrafiltration of Sewage Effluent WANG Li-ying, SHI Jie, WANG Kai-lun, et al. (20 Enhanced Treatment of Petrochemical Secondary Effluent by Biological Aerated Filter (Fe ²⁺)-Ozonation Process XU Min, WU Chang-yong, ZHOU Yue-xi, et al. (20 Specificity of Intact Ladderane Lipids in Anaerobic Ammonium Oxidizing Bacteria from Four Reactors WANG Han, FANG Fang, Li Kai, et al. (20 Characteristics of Perchlorate Reduction and Analysis of Consortium Structure in a Sulfur-Based Reactor at a High Perchlorate Concentration	(195) (201) (212) (220) (229) (238)
Effect of Dissolved Humic Acid on Thyroid Receptor Antagonistic Activity of Zinc in Aquatic Environment Al Yang, KONG Dong-dong, YU Chang, et al. (19 Optimization and Validation of the Analytical Method to Detect Common Illicit Drugs in Wastewater GAO Ting-ting, DU Peng, XU Ze-qiong, et al. (20 Cathode Electric Field Enhanced Removal of Nitrobenzene from Aqueous Solution Based on Activated Carbon Fibers (ACF)-Ozone Technique ZHAO Chun, ZHANG Shuai, ZHOU Yu, et al. (20 Effect of PVDF Hollow Fiber Ultrafiltration Membranes Modification with Carbonnanotube on Membrane Fouling Control During Ultrafiltration of Sewage Effluent WANG Li-ying, SHI Jie, WANG Kai-lun, et al. (20 Enhanced Treatment of Petrochemical Secondary Effluent by Biological Aerated Filter (Fe ²⁺)-Ozonation Process XU Min, WU Chang-yong, ZHOU Yue-xi, et al. (20 Specificity of Intact Ladderane Lipids in Anaerobic Ammonium Oxidizing Bacteria from Four Reactors WANG Han, FANG Fang, Li Kai, et al. (20 Characteristics of Perchlorate Reduction and Analysis of Consortium Structure in a Sulfur-Based Reactor at a High Perchlorate Concentration	(195) (201) (212) (220) (229) (238)
Effect of Dissolved Humic Acid on Thyroid Receptor Antagonistic Activity of Zinc in Aquatic Environment Al Yang, KONG Dong-dong, YU Chang, et al. (2) Optimization and Validation of the Analytical Method to Detect Common Illicit Drugs in Wastewater Cathode Electric Field Enhanced Removal of Nitrobenzene from Aqueous Solution Based on Activated Carbon Fibers (ACF)-Ozone Technique ZHAO Chun, ZHANG Shuai, ZHOU Yu, et al. (2) Effect of PVDF Hollow Fiber Ultrafiltration Membranes Modification with Carbonnanotube on Membrane Fouling Control During Ultrafiltration of Sewage Effluent WANG Li-ying, SHI Jie, WANG Kai-lun, et al. (2) Enhanced Treatment of Petrochemical Secondary Effluent by Biological Aerated Filter (Fe ²⁺)-Ozonation Process XU Min, WU Chang-yong, ZHOU Yue-xi, et al. (2) Specificity of Intact Ladderane Lipids in Anaerobic Ammonium Oxidizing Bacteria from Four Reactors WANG Chao, TAO Hua-qiang, SONG Yuan-yuan, et al. (2) Characteristics of Perchlorate Reduction and Analysis of Consortium Structure in a Sulfur-Based Reactor at a High Perchlorate Concentration ZHANG Chao, TAO Hua-qiang, SONG Yuan-yuan, et al. (2)	(195) (201) (212) (220) (229) (238)
Effect of Dissolved Humic Acid on Thyroid Receptor Antagonistic Activity of Zinc in Aquatic Environment Al Yang, KONG Dong-dong, YU Chang, et al. (2) Optimization and Validation of the Analytical Method to Detect Common Illicit Drugs in Wastewater Cathode Electric Field Enhanced Removal of Nitrobenzene from Aqueous Solution Based on Activated Carbon Fibers (ACF)-Ozone Technique ZHAO Chun, ZHANG Shuai, ZHOU Yu, et al. (2) Effect of PVDF Hollow Fiber Ultrafiltration Membranes Modification with Carbonnanotube on Membrane Fouling Control During Ultrafiltration of Sewage Effluent WANG Li-ying, SHI Jie, WANG Kai-lun, et al. (2) Enhanced Treatment of Petrochemical Secondary Effluent by Biological Aerated Filter (Fe ²⁺)-Ozonation Process XU Min, WU Chang-yong, ZHOU Yue-xi, et al. (2) Specificity of Intact Ladderane Lipids in Anaerobic Ammonium Oxidizing Bacteria from Four Reactors WANG Chao, TAO Hua-qiang, SONG Yuan-yuan, et al. (2) Fast Start-up of SBAF System Assisted CANON Process and the Microbial Analysis LIU Zhu-han, YUE Xiu, YU Guang-ping, et al. (2)	(195) (201) (212) (220) (229) (238) (247) (253)
Effect of Dissolved Humic Acid on Thyroid Receptor Antagonistic Activity of Zinc in Aquatic Environment Al Yang, KONG Dong-dong, YU Chang, et al. (2) Optimization and Validation of the Analytical Method to Detect Common Illicit Drugs in Wastewater Cathode Electric Field Enhanced Removal of Nitrobenzene from Aqueous Solution Based on Activated Carbon Fibers (ACF)-Ozone Technique ZHAO Chun, ZHANG Shuai, ZHOU Yu, et al. (2) Effect of PVDF Hollow Fiber Ultrafiltration Membranes Modification with Carbonnanotube on Membrane Fouling Control During Ultrafiltration of Sewage Effluent WANG Li-ying, SHI Jie, WANG Kai-lun, et al. (2) Enhanced Treatment of Petrochemical Secondary Effluent by Biological Aerated Filter (Fe ²⁺)-Ozonation Process XU Min, WU Chang-yong, ZHOU Yue-xi, et al. (2) Specificity of Intact Ladderane Lipids in Anaerobic Ammonium Oxidizing Bacteria from Four Reactors WANG Chao, TAO Hua-qiang, SONG Yuan-yuan, et al. (2) Characteristics of Perchlorate Reduction and Analysis of Consortium Structure in a Sulfur-Based Reactor at a High Perchlorate Concentration ZHANG Chao, TAO Hua-qiang, SONG Yuan-yuan, et al. (2)	(195) (201) (212) (220) (229) (238) (247) (253) (260)
Effect of Dissolved Humic Acid on Thyroid Receptor Antagonistic Activity of Zinc in Aquatic Environment Al Yang, KONG Dong-dong, YU Chang, et al. (19 Optimization and Validation of the Analytical Method to Detect Common Illicit Drugs in Wastewater Cathode Electric Field Enhanced Removal of Nitrobenzene from Aqueous Solution Based on Activated Carbon Fibers (ACF)-Ozone Technique ZHAO Chun, ZHANG Shuai, ZHOU Yu, et al. (2) Effect of PVDF Hollow Fiber Ultrafiltration Membranes Modification with Carbonnanotube on Membrane Fouling Control During Ultrafiltration of Sewage Effluent WANG Li-ying, SHI Jie, WANG Kai-lun, et al. (2) Enhanced Treatment of Petrochemical Secondary Effluent by Biological Aerated Filter (Fe ²⁺)-Ozonation Process XU Min, WU Chang-yong, ZHOU Yue-xi, et al. (2) Specificity of Intact Ladderane Lipids in Anaerobic Ammonium Oxidizing Bacteria from Four Reactors WANG Han, FANG Fang, LI Kai, et al. (2) Characteristics of Perchlorate Reduction and Analysis of Consortium Structure in a Sulfur-Based Reactor at a High Perchlorate Concentration ZHANG Chao, TAO Hua-qiang, SONG Yuan-yuan, et al. (2) Fast Start-up of SBAF System Assisted CANON Process and the Microbial Analysis LIU Zhu-han, YUE Xiu, YU Guang-ping, et al. (2) Mechanism for Effects of High Free Ammonia Loadings on Biological Nitrification JI Min, LIU Ling-jie, ZHAI Hong-yan, et al. (2) Impact of Biodegradable Organic Matter on the Functional Microbe Activities in Partial Nitrification Granules WANG Shu-yong, QIAN Fei-yue, WANG Jian-fang, et al. (2)	(195) (201) (212) (220) (229) (238) (247) (253) (260) (269)
Effect of Dissolved Humic Acid on Thyroid Receptor Antagonistic Activity of Zinc in Aquatic Environment	(195) (201) (212) (220) (229) (238) (247) (253) (260) (269) (276)
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