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紫色土丘陵区坡地柑橘园土壤碳氮的空间分布特征

李子阳1.2, 陈露1.2, 赵鹏2, 周明华1*, 郑静1, 朱波1

(1. 中国科学院水利部成都山地灾害与环境研究所山地表生过程与生态调控重点实验室,成都 610041; 2. 中国科学院大学, 北京 100049)

摘要:20世纪90年代以来,西南紫色土丘陵区大量坡耕地转变为果园,提高了农民的经济收益,但这一土地利用变化对土壤碳(C)、氮(N)空间分布特征的影响仍然缺乏研究.为探究紫色土丘陵区坡耕地转变为果园后土壤C、N的空间分布特征及其主要影响因 盆地中部紫色土丘陵区代表性柑橘园为研究对象,分析了由坡耕地转变为柑橘园后,土壤C、N空间分布特征及其主要影响因 素.结果表明,坡面位置(坡位)对土壤总氮(TN)、硝态氮(NO₃⁻-N)和可溶性有机碳(DOC)含量均有显著影响(P < 0.05),而对土 壤总有机碳(SOC)和铵态氮(NH₄⁺-N)的含量没有显著影响(P > 0.05).在0~30 cm 土层,土壤NO₃⁻-N含量沿坡面的变化趋势为: 上坡位 < 中坡位 < 下坡位,而TN和DOC含量沿坡面的变化趋势为:上坡位 > 中坡位 > 下坡位.各坡位土壤C、N含量随深度(0~30 cm)增加呈现整体降低趋势,其中土层深度对土壤TN、SOC、NO₃⁻-N和DOC含量的影响显著(P < 0.05).坡面土壤TN储量 (0~30 cm)随坡位自上而下逐渐降低(P < 0.05),其中坡上、坡中和坡下的土壤TN储量分别为2.37、1.89和1.62 t⁺hm⁻²(以 N 计).SOC储量沿坡面的分布差异不显著,变化范围(以 C ⁺)为56.12~58.48 t⁺hm⁻²,整个坡面土壤TN与SOC的储量呈显著正相 关关系.结果表明,在预测土壤C、N含量及储量对土地利用转变响应时不能忽视土壤养分的空间分布规律,研究可为理解紫色 土丘陵区耕地转变为果园后土壤C、N空间分布特征及主控因子提供研究案例.

关键词:紫色土;柑橘园;坡位;氮(N);碳(C);空间分布

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Spatial Distribution Characteristics of Soil Carbon and Nitrogen in Citrus Orchards on the Slope of Purple Soil Hilly Area

LI Zi-yang^{1,2}, CHEN Lu^{1,2}, ZHAO Peng², ZHOU Ming-hua^{1*}, ZHENG Jing¹, ZHU Bo¹

(1. Key Laboratory of Mountain Surface Process and Ecological Regulation, Institute of Mountain Hazards and Environment, Chinese Academy of Sciences, Chengdu 610041, China; 2. University of Chinese Academy of Sciences, Beijing 100049, China)

Abstract: Since the 1990s, a large area of sloping farmland in a purple soil hilly region of southwest China was converted into an orchard to prevent soil erosion, increase soil fertility, and elevate economic benefits for farmers. In order to explore the spatial distribution of soil carbon (C) and nitrogen (N) fractions on the slope of returning arable lands to citrus orchards in purple soil hilly areas, a soil sampling event was carried out in a citrus orchard at the Yanting Agro-ecological Experimental Station of Purple Soil, Chinese Academy of Sciences, to examine the differences in soil C and N fractions and their influencing factors. The results showed that the slope position had significant effects on the contents of soil total nitrogen (TN), nitrate nitrogen (NO₃⁻-N), and dissolved organic carbon (DOC) (P < 0.05), but the effects were not obvious regarding the total organic carbon (SOC) and ammonia nitrogen (NH₄⁺-N) of the soil (P > 0.05). For topsoil (0-30 cm), the variation trend of soil NO₃⁻-N content along the slope was upper slope < middle slope < lower slope, whereas the TN and DOC contents along the slope exhibited the trend of upper slope > middle slope > lower slope. The contents of soil C and N in each slope position generally showed a downward trend with increasing soil depth (0-30 cm). The contents of soil TN, SOC, NO₃⁻-N, and DOC were significantly affected by soil depth (P < 0.05). The TN storage (0-30 cm) significant difference in SOC reserves along the slope , with a value of 2. 37, 1. 89, and 1. 62 t · hm⁻² (reported as N) for the upper slope, middle slope, and lower slope, respectively. There was no significant difference in SOC reserves along the slope , with a range from 56. 12 to 58. 48 t · hm⁻² (reported as C). Our results provide scientific basis for understanding the spatial distribution of soil nutrients of the restored farmland in purple soil hilly areas. Our research suggests that the spatial distribution of soil carbon and nitrogen storage sho

Key words: purple soil; citrus orchards; slope position; soil nitrogen (N); soil carbon (C); spatial variation

碳(C)、氮(N)是植物生长必需的营养元素,也 是土壤养分的重要组成部分^[1].作为生物圈中最大的 陆地C库,土壤的C储存量远远高于植物和大气C储 存量的总和,而土壤中约有2/3的C被认为是土壤有 机碳(SOC)^[2].SOC可以直接影响土壤肥力和土壤结 构,其微小的变化都将会对大气中的二氧化碳(CO₂) 浓度产生极大的影响^[3].而土壤中的N素为生物体的 生命元素,通常是陆地生态系统初级生产力的营养 限制因子^[4],同时还直接影响SOC在土层中的累积速 率^[5,6].土壤C、N的累积受多种因素的影响,如气候 条件^[7]、地形^[8]和土地利用变化^[9]等.其中,土地利用 变化会导致植物种类、数量及生物量的改变,植物生 理变化又会影响土壤C、N含量动态变化,进而影响 土壤C、N的固存^[10].在过去的几十年里,许多国内外 研究在不同尺度上探讨了土地利用变化对土壤C、N

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- **作者简介:** 李子阳(1996~),男,博士研究生,主要研究方向为土地 利用和土壤生态,E-mail:liziyang18@mails.ucas.ac.cn
 - * 通信作者,E-mail: mhuazhou@imde. ac. cn

含量分布的影响.例如, 雷利国等^[11]在重庆缙云山的 研究发现, 林地转变为耕地后, 大大降低了 SOC 含 量; Kooch 等^[12]在伊朗北部的研究发现, 森林向牧场 的转变增加了土壤总氮(TN)含量.在全球气候变化 背景下, 研究土地利用转变下土壤 C、N空间分布变 化具有重要意义^[13-15].

紫色土坡耕地作为川中丘陵地区的主要耕地类型,具有结构差和有机质含量低等特点^[16].20世纪90年代以来,为提高山区农民的经济效益,大量坡耕地被人工种植果树建成果园^[17].对于川中丘陵区坡地而言,坡面结构与位置作为该地区的主要地形因子,影响着温度、水分和光照等在坡面的空间分布,从而影响土壤坡面C、N累积与分解^[18].此外,不同坡面位置(坡位)具有不同的土壤侵蚀与泥沙沉积,也显著影响着土壤C、N在坡面上的重新分配^[19,20].目前,关于紫色土丘陵区地形条件对坡地柑橘园土壤C、N含量与储量及其空间分布的研究不足.

因此,本研究选取川中丘陵紫色土区种植13 a的 柑橘园为对象,采用野外调查采样与室内分析相结 合的方法,分析了坡地柑橘体系不同坡位土壤C、N 含量与组分的空间分布特征与影响因素,以期为川 中丘陵紫色土地区柑橘园土壤肥力长期可持续管理 提供科学依据.

1 材料与方法

1.1 研究区域概况

研究区域位于中国科学院盐亭紫色土农业生态 试验站附近的万安小流域(E105°31′,N31°16′).该 流域位于四川盆地中北部盐亭县大兴回族乡,地处 嘉陵江流域和涪江流域的分水岭^[21],为典型的川中 丘陵沟渠集水区,平均海拔在400~600 m之间.区内 属亚热带季风气候,年均气温为17.3℃,年均降水量 为826 mm(1981~2006年),且65%的年降水量集中 在5~9月^[22].本研究的试验土壤被当地称为"紫色 土",是我国川中丘陵地区一种主要的土壤类型^[23]. 大部分紫色土区域土层较浅薄,平均土层厚度为30 ~80 cm,保水保肥差,有机质和土壤TN含量低^[24].近 20年来,川中丘陵地区大面积不适宜耕种的陡坡耕 地被人工恢复为森林和果园,而该地区作为石灰性 紫色土土地利用转变的一个极具代表性的研究地 点,备受关注^[25,26].

本研究主要在万安流域选取了一个典型的柑橘 园作为试验坡地,坡长为140m,坡宽为45m,平均坡 度夹角14.5°.该坡地于2008年由农田退耕,并休耕 一年,2009年开始种植柑橘.农田退耕前由当地村民 耕种,至少在过去的20年里,施行7个月小麦和4.5 个月玉米的轮作制度.其中小麦年施氮量为130 kg·hm⁻²(以N计,下同)、磷肥量为90 kg·hm⁻²(以P 计,下同)和钾肥36 kg·hm⁻²(以K计,下同),玉米年施 氮量为150 kg·hm⁻²、磷肥量为90 kg·hm⁻²和钾肥36 kg·hm⁻².退耕后,种植"爱媛38号"果冻橙,种植密度 为1350株·hm⁻².每年施肥3次,分别是春肥、夏肥以 及冬肥,春肥主要施用高氮复合肥(含氮量26%、含 磷量10%及含钾量10%),夏肥主要施用平衡型复合 肥(含氮、磷和钾量各16%),冬肥主要施用普通复合 肥(含氮量18%、含磷量10%以及含钾量10%).春夏 肥每株用量1500g复合肥,冬肥每株用量800g复合 肥.施用复合肥的同时,每株还配合施用5kg左右有 机肥(猪粪发酵,氮磷钾平均总含量8%).以上所有 信息均为询问当地农民和查阅官方记录获取.具体 试验选点信息见图1.



Fig. 1 Location of study area and sampling sites

1.2 样品采集与测定

采样时间为2022年10月11日,根据样地的具体 地形,结合Brubaker坡面划分原则,将研究坡位分成3 个部分:坡上、坡中和坡下位置[27].在每个坡位随机 选取3块2m×2m的样方,每个样方均去除地表覆盖 植被和枯枝落叶层.由于紫色土壤具有明显的土壤 "浅"的特征,坡上和坡中大部分采样点在30 cm 以下 混入了大块石块和母质,而坡下采样点在70 cm 左右 即到达紫色土母岩,所以本研究坡上、坡中和坡下3 个坡位的采样深度分别为0~30、0~30和0~70 cm. 使用土钻(直径5 cm)分别采集0~10、10~20、20~ 30、30~50和50~70 cm 土壤样品,每个土层取3个 样品重复,共得到土壤样品99份[3个坡位×3个地 点×(3~5)个土层×3个样品重复],同时采集对应 土样的标准环刀样品(环刀容积100 cm³). 土壤样品 在室温下风干20d以上,研磨后通过2mm筛网测定 粒径组成、pH值、硝态氮(NO₃⁻-N)、铵态氮(NH₄⁺-N) 和可溶性有机碳(DOC)含量.通过0.25 mm筛网测 定土壤有机碳(SOC)、全氮(TN)和全碳(TC)含量^[28].

采用 105 ℃ 环刀烘干法测定各垂直土层土壤容 重(BD)和土壤含水量(SWC).利用激光粒度分析仪 (Mastersizer 2000,英国)分析土壤粒径分布(PSD)^[29], 利用 pH 计(Delta 320,中国上海)测定土壤(土:水 = 1:5)的 pH 值^[30].采用连续流动分析仪(AA3,德国 SEAL)分析氯化钾浸提液中的 NO₃⁻-N、NH₄⁺-N 和 DOC 浓度.土壤 TC 和 TN 含量用碳氮分析仪测定 (CNS Elementar vario MAX,德国)^[31].采用高温外热 重铬酸钾氧化-容量法测定 SOC 的含量^[32].

1.3 数据统计与分析

SOC 和 TN 储量(t·hm⁻²,分别以 C 和 N 计,下同) 计算如下^[33]:

> SOC $ideal = \sum SOC_i \times BD_i \times D_i \times 0.1$ TN $ideal = \sum TN_i \times BD_i \times D_i \times 0.1$

式中,SOC_i和TN_i为第*i*层采样土壤的SOC和TN含量 (g·kg⁻¹),BD_i为第*i*层采样土壤的容重(g·cm⁻³), D_i 为 第*i*层采样土壤的土层厚度(cm),0.1为单位换算 因子.

采用 Kolmogorov-Smirnov 检验原始数据是否服从 正态分布(P=0.05),采用 Levene's 检验方差齐性 (P>0.05),所有数据均以平均数±标准差(SD)表示. 不同处理之间采用 LSD 最小显著差数法进行多重比 较,显著水平为0.05.采用双因素方差定量分析坡位 (因子1)和土层厚度(因子2)及其交互作用对土壤理 化性质的影响.使用 SPSS Statistics 23.0(SPSS Inc,美 国)进行相关性分析和线性回归分析,P<0.05认为 结果具有显著相关性.使用 Origin 9.0(Originlab,美 国)生成所有描述参数变化的图表.此外,使用 CANOCO 5.0(Micro Power,美国)软件进行部分冗余 分析(RDA)和方差分解分析(VPA),分别计算各因素 对土壤TN储量和SOC储量的影响.

2 结果与分析

2.1 紫色土丘陵区柑橘园不同坡位土壤物理参数的空间分布特征

退耕13a后,不同坡位的柑橘园土壤物理参数分 布如表1和表2所示. 土壤 pH 值在8.54~8.77之间, 属于典型的石灰性紫色土.所有坡位0~30cm土层, 均出现pH值随土壤深度的增加而逐渐增加的现象. 其中上坡位和下坡位0~10 cm 土层的 pH 值均显著 低于 20~30 cm 土层(P<0.05). 但是,同一土层厚度 的不同坡位间pH值没有显著差异(P>0.05).不同 坡位剖面土壤 BD 在 0~10 cm 土层最低, 为 1.15~ 1.23 g·cm⁻³,并随土层深度增加而显著增大.下坡位 表层土壤(0~20 cm)BD显著高于中坡位和上坡位土 壤.不同坡位最大SWC均出现在0~10 cm 土层(上、 中和下坡位分别为:17.04%,15.74%和15.80%),且 上坡位土壤的SWC显著低于中、下两坡位.不同坡 位的ω(粉粒)均高于50%,ω(黏粒)和ω(沙粒)的变 化范围分别为7.74~16.95%和20.49~37.41%,表 明该地区土壤质地以黏壤土和粉壤土为主(表2).如 表3所示,双因素方差分析表明坡位和土层厚度显著 影响土壤砂、粉、黏粒含量(P<0.01),且两因素交 互作用(坡位×土层深度)对上述机械组成因子也有 显著影响(P<0.05).同样地,坡位和土层厚度也显 著影响了土壤 BD(P<0.01),但对土壤 pH 值和 SWC 均无显著影响(P>0.05).

	表1 不同坡位土壤 pH值、BD和 SWC 的空间分布特征 ¹⁾
Table 1	Spatial distribution characteristics of soil pH, BD, and SWC in different slope positions

土层厚度 pH				$BD/Mg \cdot m^{-3}$		SWC /%			
/cm	坡上	坡中	坡下	坡上	坡中	坡下	坡上	坡中	坡下
0 ~ 10	8.61±0.10Aa	8.65±0.16Aa	8.54±0.07Aa	1.15±0.02Aa	1.17±0.05Aa	1.23±0.02Ba	$12.04 \pm 1.64 \mathrm{Ab}$	15.74±2.57Ba	$15.80{\pm}2.81{\rm Bb}$
$10 \sim 20$	8.63±0.14Aa	8.71±0.17Aa	8.62±0.17Aab	1.16±0.03Aa	1.19±0.04Aa	1.27±0.02Bab	13.03±0.98Aa	15.30±1.32Ba	15.10±1.41Bab
$20 \sim 30$	8.71±0.13Ab	8.77±0.17Aa	$8.68{\pm}0.15{\rm Ab}$	$1.23{\pm}0.01{ m Ab}$	$1.29{\pm}0.03\rm{Ab}$	$1.28 \pm 0.04 \mathrm{Aab}$	11.40±2.35Aa	15.63±2.25Ba	14.93±2.07Bab
30 ~ 50	—	—	$8.71{\pm}0.08{\rm b}$	_	—	$1.29{\pm}0.05{\rm b}$	—	—	12.34±0.51a
$50 \sim 70$	_	_	8.64±0.23ab	_	_	$1.31 \pm 0.03 \mathrm{b}$	—	_	14.32±1.22ab

1)不同大写字母表示同一土层不同坡位间差异显著(P < 0.05),不同小写字母表示同一坡位不同土层间的差异显著(P < 0.05);BD和SWC分别表示土壤容重和土壤含水量;"一"表示该土层未取得土样

2.2 紫色土丘陵区柑橘园不同坡位土壤碳、氮养 分含量的空间分布特征

如图2所示,3种坡位土壤NH₄⁺⁻N含量在0~ 30 cm 土层差异较小,其中上坡位0~10 cm 土层 ω (NH₄⁺⁻N)最高,为3.29 mg·kg⁻¹,并随着土层加深 呈现先降低后增加的趋势.中坡位 ω (NH₄⁺⁻N)的 变化趋势与上坡位相反,呈现先增加后降低的趋势,在 10 ~ 20 cm 土层达到最高值,为 2.44 mg·kg⁻¹. 下坡位 ω (NH₄⁺-N)在 0 ~ 30 cm 土层的变化趋势与中坡位相同,但在 50 cm 土层达到最低值 1.58 mg·kg⁻¹后,趋势转为增加,在 70 cm 土层达到 下坡位的最高值 2.61 mg·kg⁻¹[图 2(a)].上、中两

表2 不同坡位土壤机械组成的空间分布特征¹⁾

Table 2 Spatial distribution characteristics of soil mechanical composition in different slope positions

土层厚度	$\omega(\tilde{t})$	b粒)(2 ~ 0.02 mm)/	<i>ω</i> (粉粒	≿)(0.002 ~ 0.02 r	nm)/%	ω(黏粒)(<0.002 mm)/%			
/cm	坡上	坡中	坡下	坡上	坡中	坡下	坡上	坡中	坡下
0 ~ 10	24.39±1.47Aab	27.67±2.3Aab	26.98±0.56Aa	58.66±3.26Aa	59.53±3.69Aa	$58.04{\pm}4.04\mathrm{Ab}$	16.95±1.27Ba	$12.80{\pm}2.16{\rm Ab}$	14.99±1.15ABab
$10 \sim 20$	$26.60{\pm}0.98{\rm Ab}$	23.25±1.36Aa	25.39±1.47Aa	58.01±4.11Aa	$68.29{\pm}4.24\mathrm{Bb}$	$58.16{\pm}4.56{\rm Ab}$	15.39±0.69Ba	8.46±1.67Aa	16.45±1.29Bb
20 ~ 30	20.49±1.23Aa	24.78±2.10Aa	37.41±2.58Bb	62.56±2.09Ba	67.75±2.27Bb	51.65±2.17Aa	16.95±2.01Ca	7.74±0.99Aa	10.94±0.86Ba
30 ~ 50	—	—	29.15±0.87a	—	_	56.79±2.98ab	_	_	14.06±2.12ab
50 ~ 70	—	—	27.74±3.85a	—	—	$59.84 \pm 3.49 \mathrm{b}$	_	_	11.42±1.39a

1)不同大写字母表示同一土层不同坡位间差异显著(P<0.05),不同小写字母表示同一坡位不同土层间的差异显著(P<0.05);"一"表示该 土层未取得土样

表3 坡位和土层深度对土壤物理性质影响的双因素方差分析1)

Table 3	Two-way ar	alvsis of	variance to	test the	e effects	of slope	positions and	l soil la	vers on soil	nhysica	l nroi	pertie
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	项目	pН	BD	SWC	砂粒	粉粒	黏粒
	坡位	0.400	0.000^{**}	0.043	0.000^{**}	0.000^{**}	0.001**
	土层深度	0.574	0.001**	0.205	0.004^{**}	0.000^{**}	0.002^{**}
_	坡位×土层深度	0.918	0.099	0.441	0.000^{**}	0.000^{**}	0.012^{*}

1)*表示显著性水平 P < 0.05,**表示极显著性水平 P < 0.01

坡位之间 0 ~ 30 cm 土层 NO₃⁻-N 含量并无显著差异 (P < 0.05),且随土层的变化趋势较小,0~10 cm 土层 NO₃⁻-N 含量略高于 20~30 cm 土层.下坡位 0 ~ 30 cm 土层 NO₃⁻-N 含量均高于上、中两坡位,其 中 0~10 cm 土层 ω (NO₃⁻-N)在整个坡面最高,为 11.74 mg·kg⁻¹,并随着土层深度的加深,先降低后 增加[图 2(b)].上、中两坡位 0~30 cm 土层 DOC 含量同样没有显著差异(P < 0.05),且变化趋势一致,均随着土层加深而降低.与 NO_3^--N 含量在整个坡面的分布相反,下坡位土壤DOC含量均低于上、中两坡位、虽然 ω (DOC)随深度的变化趋势在 $0 \sim 30 \text{ cm}$ 土层与上、中两坡位一致,但是在30 cm 土层以下基本保持不变,维持在20.76~21.43 mg·kg⁻¹的波动范围内[图2(c)].



NH₄⁺-N、NO₃⁻-N和DOC分别为土壤铵态氮、硝态氮和可溶性有机碳 图 2 不同坡位土壤NH₄⁺-N,NO₃⁻-N和DOC含量沿土层的空间分布 Fig. 2 Spatial distribution of NH₄⁺-N, NO₃⁻-N, and DOC contents along soil layers at different slope positions

整体来看,0~30 cm 土层的不同坡位,土壤TN 含量分布呈现出上坡位>中坡位>下坡位的规律. 整个土壤剖面0~10 cm 土层 ω (TN)均是最高值,不 同坡位的波动范围为0.54~0.83 g·kg⁻¹,且随着土层 加深,土壤TN含量呈现出降低的趋势[图3(a)].对 于下坡位30 cm 以下的土层, ω (TN)维持在0.28~ 0.34 g·kg⁻¹的波动范围内,土层间差异并不显著(P> 0.05). 上、中两坡位 0 ~ 30 cm 土层ω(SOC) 几乎保持 一致,且随着土层深度的加深而降低,变化范围为 13.92 ~ 17.88 g·kg⁻¹. 下坡位 0 ~ 20 cm 土层ω(SOC) 略低于上、中两坡位,但 20 cm 以下的各土层间并无 明显差异,基本维持在 14.02 ~ 14.57 g·kg⁻¹的范围内 上下波动[图 3(b)]. 土壤 C/N 值同样也是反映应土 壤 C、N养分含量的重要指标.如图 3(c)所示,0~20 cm 土层上坡位土壤 C/N 较低,分别为 22.50(0~10 cm 土层)和 25.29(10~20 cm 土层).对于 20~30 cm

土层,下坡位土壤C/N较高,并在30 cm以下的土层中 呈现先降低后增加的趋势.



IN、SOC和C/N分別为主集忌氮、主集有机碳和主集碳氮口 图3 不同坡位土壤TN、SOC和C/N沿土层的空间分布

Fig. 3 Spatial distribution of TN and SOC contents, and C/N along soil layers at different slope positions

层加深,土层TN储量呈现降低的趋势.0~30 cm土 土壤TN和SOC储量(0~30 cm)沿坡向的空间分 层 SOC 储量在不同坡位的变化范围在 56.12~58.48 布如图4所示.整体来看,0~30 cm 柑橘林土壤剖面 TN储量随坡面呈自上而下降低的趋势,其中坡上、 t·hm⁻². 其中,中坡位SOC储量最高,但3个坡位的差 坡中和坡下的土壤TN储量分别为2.37、1.89和 异并不显著(P>0.05). 与TN储量结果相似,不同坡 1.62 t·hm⁻², 坡下土壤 TN 储量显著低于坡上(P < 位 SOC 储量的最大值均出现在 0~10 cm 土层,分别 0.05),二者相差0.75 t·hm⁻²[图4(a)].但同层土壤不 为20.13、20.92和19.74 t·hm⁻²(坡上、坡中和坡下) 同坡位间TN储量差异较小,仅仅在10~20 cm 土层, 随着土层加深,SOC储量呈现微弱降低的趋势,但土 坡上土壤TN储量显著高于坡下.同-·坡位,随着土 层间的差异同样不显著[图4(b)]. 80 (b) (a) 3 A 60 AB 2 40 SOC储量/t·hm⁻² IN储量/t·hm⁻² 1 20 0 0 883 200 Ba Aa 20 Bab Ab 1 Aa Aa Aa Aa Aa 40 $0 \sim 30 \text{ cm}$ 2 $10 \sim 10 \text{ cm}$ 10~20 cm 60 3 20~30 cm 80 坡上 坡中 坡下 披上 坡中 坡下 图 4 不同坡位 0~30 cm 土层 TN 和 SOC 储量的空间分布



2.3 不同因素对柑橘园坡地土壤碳、氮变化影响 及其贡献

坡位和土层深度对柑橘园坡地土壤C、N养分的 影响见表4.土壤NO₃⁻-N、DOC和TN含量受到坡位 和土层深度的显著影响(P < 0.05),且两因素交互作 用(坡位×土层深度)对土壤NO₃⁻-N含量也有显著影 响(P < 0.05).但是土壤NH₄⁺-N含量与坡位和土层深 度之间没有明显的相关关系(P > 0.05).此外,土层 深度对SOC含量影响显著(P < 0.05),但SOC的含量 不受坡位的影响(P>0.05).无论是坡位和土层深度,以及两因素交互作用,均对土壤的C/N影响较小(P>0.05).土壤物理参数与土壤C、N养分含量的相关矩阵如表5所示,在整个柑橘园坡地,土壤NO₃⁻-N含量与土壤pH呈极显著负相关关系(P<0.01),其余土壤C、N组分含量与土壤pH均无显著相关关系(P>0.05).除土壤NO₃⁻-N和NH₄⁺-N外,DOC、TN、SOC含量分别与土壤C/N值和BD呈极显著负相关关系(P<0.01).其余土壤物理参数与土壤C、N组分含

表4 坡位和土层深度对土壤C、N养分含量影响的双因素方差分析¹⁾

Table 4 To	wo-way analysis of v	ariance to test the ef	fects of slope positi	ons and soil layers on	soil C and N nutrient c	ontents
项目	TN	SOC	C/N	NH_4^+-N	NO ₃ ⁻ -N	DOC
坡位	0.037*	0.697	0.510	0.394	0.000^{**}	0.003**
土层深度	0.045^{*}	0.042^{*}	0.189	0.468	0.000^{**}	0.011^{*}
坡位×土层深度	0.880	0.934	0.986	0.258	0.011^{*}	0.794

1)*表示显著性水平 P < 0.05,**表示极显著性水平 P < 0.01

表5 坡面土壤物理性质和C、N养分含量的相关矩阵¹⁾

Table 5 Correlation matrix of physical properties and C and N nutrient contents of slope soils												
	$_{\rm pH}$	SWC	BD	砂粒	粉粒	黏粒	$\mathrm{NH_4^+}$ -N	NO ₃ ⁻ -N	DOC	TN	SOC	C/N
рН	1											
SWC	-0.380	1										
BD	0.161	-0.245	1									
砂粒	-0.117	0.112	0.375	1								
粉粒	0.495	0.077	-0.102	-0.736**	1							
黏粒	-0.558	-0.256	-0.339	-0.241	-0.479	1						
$\mathrm{NH_4}^+$ -N	-0.245	0.424	-0.297	-0.203	-0.005	0.269	1					
NO ₃ ⁻ -N	-0.802^{**}	0.305	-0.023	0.113	-0.357	0.367	0.168	1			~	
DOC	-0.202	0.388	-0.887^{**}	-0.577	0.376	0.209	0.279	0.070-	1	\sim	m	لم
TN	-0.375	0.347	-0.949**	-0.324	-0.006	0.429	0.317	0.117	0.884^{**}	1	12	1
SOC	-0.363	0.592	-0.853**	-0.098	-0.010	0.142	0.289	0.138	0.784^{**}	0.873**	/1_	1
C/N	0.335	-0.373	0.878^{**}	0.538	-0.217	-0.386	-0.249	-0.236	-0.901**	-0.905**	-0.738**	31

1)**表示极显著性水平 P < 0.01

量均无明显的相关关系.此外,为了 直观地研究相 橘园坡地土壤 TN含量与 SOC含量、TN 储量与 SOC 储量之间的关系,采用线性拟合方法(图5).研究 呈极显著正相 结果发现,土壤TN含量和SOC含 瞐

0.01),不同土层的TN储量与SOC储量 关关系(P< 也呈显著正相关关系(P<0.05),回归斜率分别为 6.6和3.8, R²分别为0.76和0.41, P值均小于0.05, 拟合效果较好



图 5 坡地土壤 TN 与 SOC 含量; TN 与 SOC 储量的相关性拟合

Fig. 5 Relationships between SOC and TN contents, and between SOC and TC storages on the slope

根据各环境变量(坡位、土壤深度及土壤理化因 子:BD、SWC、pH、粉粒和黏粒含量)和土壤TN和 SOC含量的RDA结果可见(图6),第一轴土壤TN和 SOC含量与环境因子相关性为0.94,第二轴土壤TN 和 SOC 含量与环境相关性为 0.93, 排序结果可靠. RDA1和RDA2解释率分别为85.3%、3.2%,两轴能 解释 88.6% 的差异信息.SWC、土层深度和土壤 BD 对 TN 和 SOC 含量变异的解释度分别为 55.6% (P < (0.01)、39.2% (P < 0.05)和32.5% (P < 0.05), 是影

响土壤TN和SOC含量变化3个主要的环境因子.此 外, VPA分析表明(图7), 采样位置(坡位和土层深 度)和土壤物理因子(BD、pH、SWC、黏粒和粉粒含 量)可以解释土壤TN储量97.6%的变异[图7(a)]和 SOC储量42.2%的变异[图7(b)]. 样地位置和土壤 物理因子是影响土壤 TN 储量的主要因素. 它们分别 解释了 16.6% (P=0.07)和4.4% (P=0.25)的变异, 2个因素的交互作用解释了 76.6% 的变异 (P < 0.05). 采样位置对 SOC 储量变化不能提供相应的解

释,但是土壤物理因子可以解释 SOC 储量 15.9% (P=0.45)的变异,且2个因素的交互作用解释了 40.6%的变异(P<0.05).



照辐射^[54]均具有较大的差异,这些条件会影响植被 长势,土壤微生物的活性以及凋落物的分解程度^[35], 从而共同决定了坡地土壤养分含量的差异以及分布 的空间异质性^[36].本研究发现,坡位显著影响紫色土 柑橘园剖面土壤物理参数和C、N养分分布,且坡位 对不同土壤参数的影响差异较大(表3和表4).土壤 NO₃⁻N含量在坡面的分布表现为:下坡位>中坡 位>上坡位[图2(b)],这与Zhang等^[28]在黄土高原丘



此外,本研究还发现,0~30 cm 土壤 NO₃⁻-N、 NH₄⁺-N和DOC含量随土壤深度的变化趋势各异,但 是整体呈现出随土层加深波动下降的趋势(图2).这 与汝海丽等^[45]对黄土高原坡地的研究结果一致.除 坡形和坡位外,影响养分在土壤剖面垂直分布的主 要因素是径流强度、与土壤入渗能力相关的土壤结 构、土壤溶液中各养分的吸附能力以及植物对养分 的吸收和利用^[46,47].土壤 NO₃⁻-N、NH₄⁺-N和DOC均



土壤物理性质包括:pH,容重,土壤含水量,黏粒含量,粉粒含量;采样位置包括:坡位和土层深度;数字表示每个解释变量所解释的方差的比例 图 7 基于方差分解分析(VPA)的土壤物理性质和采样位置对土壤 TN和 SOC 储量的影响

Fig. 7 Effects of soil physical properties and sampling positions on TN and SOC storages based on variation partitioning analysis (VPA)

属于土壤C和N的速效养分,易溶于水且容易被植物 吸收利用.随自然降雨入渗到下层后,被大量果树的 根系生长消耗,从而造成下层土壤C和N的速效养分 低于上层的现象^[40].而柑橘园下坡位拥有更深的土 层,土壤NO₃⁻N和NH₄⁺-N含量在50 cm土层往下出 现含量升高的趋势,侧面证明了更深层土壤在没有 植物利用的情况下,会出现速效N素累积的现象.

3.2 影响紫色土丘陵区柑橘园土壤TN含量、SOC 含量和土壤C/N变化的主要因素

土壤 TN 和 SOC 的含量不仅是体现土壤肥力的 直接指标,还可以指示土壤C、N养分矿化输入和损 失平衡的结果. 与速效 C、N 养分变化相似, 柑橘园 土壤 TN 和 SOC 含量在坡面的分布也会受到上述多 种因素的共同影响.一般研究认为,在土壤侵蚀或雨 水冲刷作用下,坡位由上至下TN和SOC含量会逐渐 增加^[48,49]. 而本研究发现,土壤TN含量在坡面的分布 表现为:上坡位>中坡位>下坡位[图3(a)],且受到 坡位的显著影响(表4). 虽然SOC含量在坡面上的分 布不受空间位置的影响,表现为不同坡位间差异不 显著,但下坡位0~20 cm 土层 SOC 含量仍略低于上、 中两坡位,此处借助RDA分析来解释上述结果产生 的原因. 在众多土壤物理因素中,SWC和BD是影响 坡面土壤TN和SOC的含量分布的关键因素(图6), 其中 TN 和 SOC 含量与 BD 均呈显著负相关关系(表 5). 前人研究表明,坡面土壤随着 BD 增大,产流时间 提前,径流系数增大,入渗率减小,泥沙侵蚀和养分 流失严重^[50].研究区域内0~20 cm 土层,上坡位 BD 显著低于下坡位(表1),可合理推测下坡位的土壤侵 蚀更严重,从而造成果园下坡位TN和SOC含量会明 显减少.另有前人研究发现,林地坡面枯枝落叶的分 解系数(枯枝落叶量/枯枝落叶层)由坡下到坡上逐渐 减小,分解率明显降低[41]. 柑橘园上坡位土壤含水量 显著低于中坡位和下坡位(表1),水分丧失较快无法 为微生物分解提供足够的水分,土壤微生物对有机 C、N的分解速率低,导致上坡位TN和SOC的含量要 高于下坡位.同时,下坡位地势较平,靠近路基村道, 人为活动更加频繁,翻耕土壤使土壤有机C、N分解 速度更快,积累困难^[51].土壤C/N是土壤质量变化的 重要指标,其大小会影响植物的养分利用效率,对生 态系统中C、N循环过程起着决定作用^[52]. 柑橘园坡 地的C/N表现出下坡位高于上坡位,中坡位在二者间 徘徊的现象,这与TN和SOC沿坡面的分布相反.谢 柠枍等[53]在紫色土植烟坡地也发现了同土层中土壤 C/N变化规律与土壤C、N基本相反的现象,表明整个 坡面上TN和SOC积累并非同步,土壤TN和SOC因坡 位和土壤物理性质以及微生物的差异而发生变化,

最终导致土壤 C/N 在不同坡位呈现出一定的差异性.

柑橘园土壤剖面 TN 和 SOC 含量分布具有明显 的表聚性,并在0~30 cm 土层呈现随土层深度的加 深,含量均在降低的现象.双因素方差分析表明,土 壤TN和SOC含量均受到土层深度的显著影响,同时 RDA分析进一步证明,土层深度对土壤TN和SOC含 量变异解释度高达39.2%,仅次于SWC的解释度.刘 延坤等[54]对长白山人工林的研究也得到了类似的结 果. 这是因为,土壤表层累积了大量的枯枝落叶和腐 殖质,加上光、水、热和通气状况等良好,为微生物和 土壤动物生长提供良好环境,加速了地表凋落物分 解为土壤有机C、N,并且富集在土层表面[55]. 之后随 水分或其他介质向土壤下层迁移扩散,形成土壤C、 N含量从表层向下层逐渐降低的分布格局[56].整体来 看,整个坡面的C/N不受土层深度的影响(表4),这与 梁珂等[57]在重庆三峡紫色土区的研究结果一致,进 一步证明了紫色土坡耕地退耕后土壤 C/N 仍然相对 稳定的结论。

3.3 土地利用方式转变后紫色土坡地土壤TN和 SOC储量变化的主要影响因素

目前已有大量研究表明,植被恢复导致的生物 量增加会影响土壤C、N的积累^[58],而牧场或林地的 开垦会使土壤物理性质退化,从而造成更多的土壤 C、N损失^[59],本研究表明,紫色土坡耕地转变为柑橘 园 13 a 后, 坡上、坡中和坡下的土壤 TN 储量(0~30 cm)分别为2.37、1.89和1.62 t·hm⁻²,坡面TN储量呈 现出自上而下显著降低的趋势[图4(a)];SOC储量沿 坡面的分布差异不显著,变化范围在56.12~58.48 t·hm⁻²[图4(b)]. 由于储量是根据含量计算而来,因 此土壤储量的变化趋势与导致原因和土壤含量基本 一致,此处不加赘述.值得注意的是,本研究只探讨 了 0~30 cm 土层的土壤 TN 和 SOC 的储量,原因是 上、中两坡位的土层厚度只有 30 cm, 如果计算整个 土层的总储量,下坡位土壤TN和SOC的储量是最高 的. 基于回归分析, 土壤 TN和 SOC 的含量及储量均 呈现高度正相关关系(图5),说明土地利用转变过程 中存在较强的土壤C、N相互作用^[60],熊杏等^[61]在南 方红壤以及张春华等[62]在东北黑土的研究均得出了 相同的结论.分析认为,出现上述结果的原因是由于 土壤中大部分氮素和碳素均以有机状态存在于土壤 有机质中,其动态变化均受土壤有机质分解的影 响^[63],导致TN和SOC储量在坡面的高低分布也呈现 出相同的趋势. VPA分析结果表明,柑橘园坡面土壤 TN和SOC储量受到坡位、土层深度和土壤物理性质 的影响,且这3个因素的共同影响对土壤TN和SOC 储量变异的解释度最大(图7).因此,坡耕地转变为

柑橘园后会通过影响土壤 C、N 的输入和输出平衡, 以及改变土壤结构来影响土壤 C、N 储量^[64].此外,微 生物在土壤的生物地球化学循环和生态系统功能中 发挥着重要作用^[65].未来在研究土地利用方式对紫 色土丘陵生态系统不同地形条件和不同恢复年限植 被恢复后土壤-植物系统碳、氮的动态变化的影响机 制时,还应综合考虑土壤微生物的相关特性.

4 结论

(1)紫色土丘陵区柑橘园坡面土壤不同 C、N 组 分在空间上的分布具有较大差异,并受到坡位和土 层深度的显著影响.其中,土壤 NO₃⁻-N含量从上坡到 下坡逐渐增加,TN 和 DOC 含量从上坡到下坡逐渐降 低,NH₄⁺-N 和 SOC 含量在坡面上的分布没有显著差 异.各坡位表层土壤的 C 和 N 养分含量均大于底层 (NH₄⁺-N 含量除外).

(2)土壤物理性质也会显著影响坡面土壤C、N 养分含量.其中,土壤NO₃⁻-N含量与土壤pH呈极显 著负相关关系,除土壤NO₃⁻-N含量和NH₄⁺-N含量外, DOC含量、TN含量、SOC含量和土壤C/N与土壤BD 呈极显著负相关关系.

(3)整体来看,紫色土柑橘园坡地土壤TN储量 (0~30 cm)沿坡面的变化趋势为:上坡位>中坡位> 下坡位,且上坡位土壤TN储量显著高于下坡位,而 SOC储量沿坡面的分布差异不显著,整个坡面TN与 SOC的储量呈显著正相关关系.综上所述,土壤TN 和SOC储量受到坡位、土层深度和土壤物理性质的 共同影响.因此,紫色土丘陵区坡耕地向果园的转变 应重视土地利用类型及其转变后土壤C、N在坡面的 迁移与重新分配规律,并据此优化养分管理措施,有 效提升资源利用效率和减少农业面源污染问题.

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