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长江下游沿江平原土壤发育过程中碳库分配动态

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摘要:土壤发育和土地利用过程中土壤碳库的分配动态是揭示碳循环过程的关键.为了明晰土壤碳库分配及其变化趋势,在长 江下游沿江平原典型区建立土壤围垦时间序列(围垦0、60、160、280、1000和1500a),对不同土地利用方式下表层土壤有机碳 (SOC)、无机碳(SIC)、颗粒态(POC)和矿物结合态有机碳(MAOC)的含量、密度及土壤固碳潜力(CSP)等指标进行测定和估算. 结果表明,围垦1500a后,由长江冲积物母质发育的SOC含量经过围垦初期的下降后上升4.9%,而SIC经过快速的淋失,含量已 由初期占总碳含量的25.8%普遍降至0.2%.MAOC含量总体上高于POC,对SOC积累贡献率达48.0%~79.7%.区内有机碳密 度(SOCD)占总碳密度的57.4%~100%,土壤碳饱和水平(CSL)为18.6%~56.1%,水旱轮作的CSP相较于光滩增长了20.8%.碳 氮比和全氮含量是解释土壤碳积累过程的关键因素,围垦年限对评价土壤碳饱和水平有重要作用.沿江平原区土壤经长期利用 后必须注重保持养分平衡,以维持土壤生产能力并促进SOC积累,避免土壤固碳能力下降.

关键词:有机碳;无机碳;颗粒态有机碳;矿物结合态有机碳;土壤时间序列;土地利用

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Soil Carbon Pool Allocation Dynamics During Soil Development in the Lower Yangtze River Alluvial Plain

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Abstract: The allocation dynamics of soil carbon pools during soil development and land use are the key to revealing the carbon cycle process. To clarify the distribution of the soil carbon pool and its change trend, a soil reclamation chronosequence (0 a, 60 a, 160 a, 280 a, 1 000 a, and 1 500 a reclamation) was established in a typical alluvial plain in the Lower Yangtze River, and the content and density of soil organic carbon (SOC), soil inorganic carbon (SIC), particulate organic carbon (POC), and mineral-associated organic carbon (MAOC), along with carbon sequestration potential (CSP) indicators of topsoil under different land use types were measured and analyzed. The results showed that after approximately 1 500 a reclamation, the SOC content developed from the Yangtze River alluvial deposits generally increased by 4.9% after the initial decline, whereas the SIC content decreased to 0, 2% from 25.8% of the total carbon content due to its rapid leaching. The MAOC content was normally higher than that of POC, and MAOC was contributing 48.0% -79.7% of the SOC accumulation. In this region, the soil organic carbon density (SOCD) accounted for 57.4%-100% of the total carbon density, the soil carbon sequestration levels (CSL) ranged from 18.6% to 56.1%, and CSP under paddy-dryland rotation increased by 20.8% compared to that under dryland. The C/N ratio and total nitrogen content are key factors in explaining soil carbon accumulation processes, and the reclamation year plays an important role in evaluating soil carbon sequestration levels. After long-term utilization, the cultivated soil in the Yangtze River floodplain must be carefully managed through balanced fertilization to maintain soil productivity, promote the accumulation of SOC, and avoid the decline in soil carbon sequestration capacity.

Key words: organic carbon; inorganic carbon; particulate organic carbon; mineral-associated organic carbon; soil chronosequence; land use

土壤碳库由有机碳库和无机碳库构成,明晰土 壤碳库分配及其动态变化对于揭示土壤碳循环过程 和碳源-汇功能的转化具有重要意义^[1].近年来的相 关研究常将SOC在操作上分为两个或更多个具有不 同稳定机制的部分,如由粒径定义的颗粒态有机碳 (POC,>53 μm)和矿物结合态有机碳(MAOC,<53 μm),因其能够在成因、周转和功能方面将有机碳分 为两种具有明显差异的组分,成为解释有机碳库与 土壤结构关系的较理想分离方式^[2,3].各个土壤碳组 分的积累过程和转化速率不同,因而其含量和分配 比例成为评价土壤固碳潜力的关键指标^[1,4].

当前,由人为驱动的土地利用和土地覆盖变化 已导致大量的土壤碳损失^[5],为了改善土壤碳库对环 境变化的响应预测,须加强土壤碳组分动态及其影 响因子研究^[6,7].一方面,土壤发育过程中物质成分和 环境条件的变化往往通过影响土壤养分有效性、有 机物料来源及微生物活性等因素控制土壤碳的循 环,相同母质和气候条件下,土壤理化性质及各碳组 分的变化速率和轨迹也可能不尽相同^[8~10].Yang等^[11] 有关草原恢复年代序列的研究强调微生物对 SOC 稳 定的主要贡献.Tao等^[12]有关国内农田土壤的研究表 明,水分和氮肥施用是无机碳密度损失的关键驱动 因素.Cotrufo等^[13]和Hemingway等^[14]的研究揭示与矿 物结合是 SOC 的重要稳定机制,其间相互作用限制

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了微生物对其他可分解基质的接触,因此增加MAOC 可能是提高土壤固碳能力并维持土壤长期碳储的关 键[2]. 另一方面,人类活动通过一系列的土地利用在 很大程度上影响了土壤水分和养分状况、微生物及 有机物料的输入输出,从而决定了土壤碳的周转、分 配和固碳潜力[15,16]. 蔡岸冬等[4]研究表明国内林地土 壤 SOC 和 MAOC 含量显著高于草地土壤, 刘振杰 等印研究表明生物炭添加可以增加免耕和翻耕土壤 SOC含量,但对SIC含量没有显著影响.可见SOC和 SIC的分解平衡受到多种条件的调节和制约,需要改 进土壤管理方法,促进土壤碳固存[18,19].

长江中下游沿江平原广泛分布着由河流冲积物 母质发育的土壤,存在多种碳输入源[20],同时由于地 形平坦、水热条件好,农业利用强度很高,土壤生态 系统深受人类活动影响,但目前对这一地区土壤碳 分布及其在较长时间尺度上的变化还缺乏系统的研 究.据此,本研究基于长江下游沿江平原地区建立的 土壤时间序列,运用时空代换方法,量化不同土地利 用下 SOC、SIC 以及 POC、MAOC 的动态演变,以期 揭示人类活动影响下土壤碳的分配、周转和存储 过程.

1 材料与方法

研究区概况 1.1

研究区位于安徽省中南部的无为市(东经 117°28′~118°21′,北纬30°56′~31°30′),地处长江 下游北岸(图1). 在长江河道演变过程中,研究区因 处于凸岸而持续接受长江沉积物的堆积,形成不断 淤长的冲积平原.在随后千余年的开垦利用过程中, 当地不断修建江堤以保护圩区内的居民和土地.在 早期江堤建成后,由于江岸和冲积平原继续南移扩 张,又陆续修建新江堤并形成新的圩区.各圩区的形 成时间和开垦利用历史由南向北逐渐增加.根据《无 为县志》和《无为大堤志》记载的建堤时间和实地调 查走访,确定本研究区由南向北可划分为距今分别 约为0、60、160、280、1000和1500a的6个圩区(图 1),构成围垦历史和土壤发育由新到老的时间序列, 其中1区为代表土壤母质的河漫滩现代冲积物.

研究区属亚热带季风气候区,四季分明,雨量充 沛,无霜期长,适宜农业生产,土地利用集约化程度 高.区内土壤均为长江冲积物发育的潮土和水稻 土.2区即围垦时间最短的圩区,均为旱作,3区开始 有水旱轮作(旱作以小麦、油菜和蔬菜种植为主),围 垦时间较长的4~6区以水旱轮作为主,大多为稻/麦 (油菜)轮作,有少量旱作,少见成片的林地和草地.



1.2.1 样品采集与土壤理化性质测定

图1中1区为现江堤外的河漫滩,样品采自光 滩、草地和林地,2~4区包括人工林地和农业用地, 样品采自林地、旱地和水旱轮作地,5和6区以农业 用地为主,基本没有成片的林地,样品采自旱地和水 旱轮作的耕地. 各区每种土地利用土壤至少5个重 复,采样深度为0~20 cm 土层,共采集表层土壤/沉积 物样品90个.

土壤pH采用电位法测定;土壤容重(BD)采用环 刀-烘干法测定;土壤黏粒、粉粒和砂粒含量采用比 重计法测定;SOC含量采用重铬酸钾-外加热法测定; SIC含量采用气量法测定;土壤全氮(TN)含量采用凯 氏定氮法测定;土壤常量元素 Al、Ca和 Fe采用 X射 线荧光光谱法(XRF)测定^[21].

1.2.2 土壤有机碳组分分离

参考 Cambardella 等^[22]的方法对 SOC 进行物理分 组:取过2mm筛的干土20g,置于250mL三角瓶,加 入 60 mL 5 g·L⁻¹的六偏磷酸钠溶液,于 25℃ 180 r·min⁻¹振荡18 h. 将土壤溶液过53 μm筛,用去离子 水反复淋洗直至水流清澈,分别用烧杯收集筛上筛 下部分,大于53 µm部分测得的SOC为POC,小于53 µm部分测得的 SOC 为 MAOC. 测定 SOC 含量前将所 有待测样品经过60℃烘干至恒重,研磨过0.149 mm 筛.实验中每种土壤类型3个重复,碳含量的计算见 式(1)和式(2).

$$POC = SOC_{\underline{m}\underline{n}\underline{k}\underline{a}} \times m_{\underline{m}\underline{n}\underline{k}\underline{a}} / m \tag{1}$$

$$MAOC = SOC_{\vec{w} \text{ bischem}} \times m_{\vec{w} \text{ bischem}} / m \qquad (2)$$

式中,m 题版本为分离后颗粒态有机物质量(g),m 面面的 为分离后矿物结合态有机物质量(g),m为分离前总 土壤质量(g),SOC max 为颗粒态有机物中有机碳含量 (g·kg⁻¹),SOC_{矿物结合态}为矿物结合态有机物中有机碳含 量(g•kg⁻¹).

1.2.3 相关指标计算

MAOC/POC 值是反映土壤有机质稳定程度的重 要指标,更高的 MAOC/POC 值代表更稳定的 SOC 结构^[2,3].

表层土壤有机碳密度(SOCD)、土壤无机碳密度 (SICD)、颗粒态有机碳密度(POCD)和矿物结合态有 机碳密度(MAOCD)的计算见式(3)~(6).

$$SOCD = SOC \times BD \times (1-Gr) \times Th$$
 (3)

 $SICD = SIC \times BD \times (1-Gr) \times Th$ (4)

$$POCD = POC \times BD \times (1-Gr) \times Th$$
 (5)

 $MAOCD = MAOC \times BD \times (1-Gr) \times Th$ (6)

[24]的研究

将粉粒和

式中,Th为土层厚度(m),Gr为土壤砾石体积分 数(%).

基于 Hassink^[23]和徐嘉晖等

主,只有少部分土壤呈酸性,土壤质地以黏壤土 土和砂壤土为主,少部分为粉黏壤土、粉壤土、黏土 和壤砂土.从变异系数可看出,除土壤 pH 和容重表 现为弱变异性,其余属性的变异系数变化范围为 11.1%~47.4%,均表现为中等强度变异

23.1	Table 1 Statistical characteristics of soil physical and chemical properties										
土壤属性	最小值	最大值	范围	均值	标准差	变异系数	偏度	峰度			
рН	5.42	8.41	2.99	7.47	0.67	9.02	-1.50	1.85			
容重	0.90	1.62	0.72	1.37	0.12	8.88	-1.09	2.38			
黏粒	7.92	44.15	36.23	24.31	7.87	32.40	-0.13	-0.68			
粉粒	8.60	55.42	46.82	40.01	9.02	22.56	-1.13	1.34			
砂粒	13.92	83.48	69.56	35.68	15.32	42.94	1.10	0.56			
TN	0.55	2.36	1.80	1.46	0.37	25.56	-0.13	-0.12			
Al_2O_3	10.12	16.82	6.69	13.80	1.53	11.11	-0.50	-0.47			
CaO	0.92	4.81	3.90	2.44	1.16	47.41	0.49	-0.91			
$\mathrm{TFe_2O_3}$	3.85	7.17	3.32	5.68	0.73	12.91	-0.21	-0.37			
C/N	5.92	19.55	13.63	11.32	2.19	19.34	1.02	2.52			

表1 土壤理化性质统计特征1)

1)pH和C/N无量纲,容重单位为g·cm⁻³,黏粒、粉粒和砂粒的单位为%,TN单位为g·kg⁻¹,Al,O₄、CaO和TFe,O₄的单位为%,变异系数单位为%

2.2 SOC 和 SIC 含量特征

土壤发育和土地利用方式变化促进了 SOC 积累 和SIC的淋失(图2),SOC含量积累最大值出现在围 垦中后期(280~1000 a),SIC含量损失最快速阶段出 现在围垦早期(0~60a). 光滩 SOC 含量最低,随着围 垦时间增加,林地SOC含量呈现先减少、后增加的趋 势,280 a时相较 0 a提高了 1.70%,旱地和水旱轮作 地SOC含量呈现先增加、后减少的趋势,1000 a的旱 地 SOC 含量相较 60 a 时提高了 55.8%, 280 a 的水旱 轮作地 SOC 含量相较 160 a 时提高了 36.3%. 围垦

1500 a后 SOC 含量又略有下降.

未围垦区域的 SIC 含量最高 (2.90~7.37 g·kg⁻¹),随着围垦时间增加,林地、旱地和水旱轮作 地SIC含量均呈现减少趋势,60a时的林地和旱地相 较0a的光滩分别降低了68.1%和61.5%,围垦1500 a后多数样品中的 $\omega(SIC)$ 都在0.20g·kg⁻¹以下.

2.3 SOC分配及其动态

POC 和 MAOC 含量的变化趋势与 SOC 相似(图 3). 未围垦区域的草地和林地ω(POC)(6. 19 g·kg⁻¹和 7. 52 g·kg⁻¹) 和 ω (MAOC) (10. 3 g·kg⁻¹ 和 9. 65 g·kg⁻¹)

黏粒结合态碳库定义为土壤碳库稳定容量[13],土壤 固碳潜力(CSP)和土壤碳饱和水平(CSL)的计算见式 $(7) \sim (9).$

$$A = \omega(34\pm) + \omega(8\pm) \tag{7}$$

$$CSP = [(4.09+0.37 \times A) - MAOC] \times$$

$$BD \times (1-Gr) \times Th \times 0.01$$
 (8)

$$CSL = MAOC/(4.09+0.37 \times A)$$
(9)

式中,A为土壤黏粒与粉粒质量分数之和(%). 1.3 统计方法

应用 IBM SPSS Statistcs 26 进行数据处理和双因 素方差分析,采用LSD和Dunnett T3检验进行指标差 异显著性的多重比较(P < 0.05),应用Canoco 5进行 冗余分析,应用Origin 2021进行绘图.

2 结果与分析

2.1 土壤理化性质统计特征

研究区土壤理化性质的统计结果见表1,区内表 层土壤 pH 值范围为 5.42~8.41,以偏碱性土壤为



不同大写字母表示不同围垦年限间含量差异显著(P < 0.05),不同小写字母表示不同土地利用间含量差异显著(P < 0.05) 图 2 不同围垦年限各利用土壤 SOC 和 SIC 含量分布

Fig. 2 Distribution of SOC and SIC content under various land uses in different reclamation times

均高于光滩(3.92 g·kg⁻¹和4.37 g·kg⁻¹). 围垦后, 同一 时期各利用 POC 和 MAOC 含量没有显著差异, 但土 壤 POC 和 MAOC 随围垦时间延长整体呈增加趋势, 林地 POC 和 MAOC 在围垦 280 a 比围垦 60 a 时分别提 高了 36.1% 和 136%, 旱地 POC 和 MAOC 在围垦 1 000 a 比围垦 60 a 时分别提高了 77.3% 和 134%,水 旱轮作地 POC 和 MAOC 在围垦 1 000 a 比围垦 280 a 时分别提高了 99.5% 和 151%. 各利用 POC 含量随围 垦时间的变化均未达显著差异水平,而 MAOC 含量 因土壤发育和土地利用过程有较大改变.





Fig. 3 Distribution of POC and MAOC content under various land uses in different reclamation times

在不同围垦阶段,POC和MAOC对SOC贡献率变 化范围分别为20.3%~52.0%和48.0%~79.7%(图 4).林地、旱地和水旱轮作地POC对SOC贡献率变化 范围分别在35.2%~49.8%、25.9%~40.8%和 40.3%~49.1%之间,MAOC对SOC贡献率变化范围 分别为50.2%~64.8%、59.2%~74.2%和50.9%~ 59.8%.未围垦区域MAOC/POC表现为:光滩<林 地<草地,经过一段时间的围垦利用,林地、旱地和 水旱轮作地土壤MAOC/POC呈现波动增大的趋势, 且在同时期均表现为:水旱轮作<林地<旱地,说明 旱地SOC稳定性最高(图5).

对 POC、MAOC 与 SOC 进行线性拟合,发现 POC 和 MAOC 与 SOC 均存在显著正相关关系(图 6).在 3 种土地利用类型中,POC 随 SOC 含量增加的变化率表

现为:旱地 < 林地 < 水旱轮作,MAOC变化率表现为: 水旱轮作 < 林地 < 旱地,说明在增加一个单位 SOC 时,旱地表现出较低的 POC 增加速率和较高的 MAOC 增加速率.不同利用方式下 MAOC 变化率均大于 POC 变化率,说明在增加一个单位 SOC 时,将有较多 的 SOC 以 MAOC 的形式存在,MAOC 对 SOC 的贡献占 主导地位.

2.4 土壤碳密度与固碳潜力

土壤碳含量和容重共同决定了表层土壤的碳 密度与固碳潜力.研究区内SOCD和SICD范围分 别为1.76~8.00 kg·m⁻²和0.00~2.07 kg·m⁻²,CSL 为18.6%~56.1%,其中水旱轮作地POCD、 MAOCD、SOCD、CSP和CSL均值均高于其他利用 类型,相较光滩分别提高了129%、167%、149%、







20.8%和73.2%(表2).经过长期围垦利用,旱地 CSL最大值出现在围垦1500a,而水旱轮作CSL最 大值出现在1000a.

对土壤围垦年限和土地利用方式进行主效应分析,发现围垦年限×土地利用交互作用仅对SOC、SIC和CSP有显著影响(P<0.05),MAOC、MAOCD、SIC、SICD、CSP和CSL主要受围垦年限影响

 $(F_{\pm u v \eta \Pi} < F_{\Pi \not E + \overline{u}})$, POC、POCD、SOC和SOCD主要受 土地利用影响 $(F_{\Pi \not E + \overline{u}} < F_{\pm u v \eta \Pi})$ (表 3).

通过冗余分析探究土壤碳相关指标和土壤属 性指标之间的关系,发现C/N和TN含量是解释土 壤碳积累过程的最关键因素,总体解释度分别达 到53.5%和32.1%(图7).围垦过程中,较高的 C/N、TN、Al₂O₃和TFe₂O₃含量促进了SOC、POC和 MAOC的含量积累和密度提高,较低的pH和CaO 含量则伴随着较低的SIC含量和密度,围垦年限对 评价CSL有重要作用.

3 讨论

3.1 POC、MAOC及其分配情况

POC被认为是活性有机碳库的重要组成部分和 衡量标准,由主要来自植物的结构性高分子化合物 组成,在土壤中的周转速度较快,对于表层土壤中植 物残体积累和根系分布的变化极为敏感^[25],具有较 高POC/SOC值的土壤有机质活性高,容易被矿化.一 般而言,周转时间较短的组分可能对土壤有机质增 长更为敏感,长期管理将会促进土壤中易分解的活 跃颗粒态有机物在团聚体内的物理闭蓄以及土壤矿 物的化学吸附^[26],这是养分有效性和有机碳质量提 高的表现.研究区POC含量没有明显的变化规律,



图 6 研究区土壤 POC、MAOC 与 SOC 的变化趋势 Fig. 6 Trends of POC, MAOC, and SOC in study area

表 2 各利用类型土壤碳密度和固碳功能指标¹⁾

Table 2 Soil carbon density and carbon sequestration functional indicators under various land uses									
SOCD/kg	$\mathbf{r} \cdot \mathbf{m}^{-2}$	SICD/kg·m ⁻²	POCD/kg·m ⁻²	MAOCD/kg·m ⁻²	$CSP/kg \cdot m^{-2}$				

利用	回重 年限/a	$SOCD/kg \cdot m^{-2}$	$SICD/kg \cdot m^{-2}$	$POCD/kg \cdot m^{-2}$	$MAOCD/kg \cdot m^{-2}$	$CSP/kg \cdot m^{-2}$	CSL/%
光滩	0	$2.30\pm0.31\mathrm{c}$	$1.86 \pm 0.36a$	$1.09\pm0.15\mathrm{b}$	1.21 ± 0.16b	$4.16\pm0.44\mathrm{b}$	22.61 ± 2.26b
草地	0	$4.04\pm0.54\mathrm{b}$	$1.24\pm0.12\mathrm{b}$	$1.52\pm0.24\mathrm{b}$	2.53 ± 0.31a	$4.47\pm0.19\mathrm{b}$	35.99 ± 1.99a
	0	$4.88\pm0.28\mathrm{ABa}$	1.80 ± 0.07 Aa	$2.14\pm0.27\mathrm{Aa}$	2.74 ± 0.14 Ca	$5.17\pm0.36\mathrm{A}$	$34.64 \pm 1.77 A$
** +++	60	$3.51 \pm 1.13 \mathrm{B}$	0.57 ± 0.16BC	$1.73\pm0.49\mathrm{A}$	1.77 ± 0.65 C	$3.32\pm0.71\mathrm{B}$	$34.89 \pm 13.06 A$
孙卫	160	$5.00\pm0.25\mathrm{ABa}$	$0.70\pm0.44\mathrm{Ba}$	$2.06\pm0.33\mathrm{Aa}$	2.94 ± 0.31Ba	$6.02\pm0.53\mathrm{A}$	$32.85 \pm 3.77 \mathrm{A}$
	280	6.31 ± 1.11Aa	0.14 ± 0.19 Ca	$2.28\pm0.95\mathrm{Aa}$	$4.03 \pm 0.49 \mathrm{Aab}$	$5.34 \pm 0.59 \mathrm{A}$	$43.05 \pm 5.33 \mathrm{A}$
_	60	$2.19\pm0.56\mathrm{B}$	$0.69 \pm 0.14 \mathrm{A}$	$0.81 \pm 0.25 \mathrm{A}$	$1.39\pm0.33\mathrm{B}$	2.94 ± 0.51 B	31.88 ± 2.12A
1	160	$3.35\pm0.35\mathrm{Bb}$	0.56 ± 0.19 ABa	1.36 ± 0.14 Aa	$1.98 \pm 0.24 \mathrm{Bb}$	$5.37\pm0.19\mathrm{A}$	26.94 ± 2.23A
旱地	280	$3.72\pm0.53\mathrm{ABb}$	0.44 ± 0.13Ba	1.01 ± 0.17Aa	$2.72\pm0.67\mathrm{ABb}$	$4.68\pm0.91\mathrm{A}$	$36.97 \pm 10.09 \text{A}$
11	1 000	$5.05\pm0.47\mathrm{A}$	$0.00 \pm 0.00C$	$1.55\pm0.28\mathrm{A}$	3.49 ± 0.41A	$5.70\pm0.67\mathrm{A}$	$38.08 \pm 5.28 \mathrm{A}$
23	1 500	$4.68 \pm 1.51 \mathrm{AB}$	0.01 ± 0.01 C	$1.22\pm0.45\mathrm{A}$	$3.46 \pm 1.07 \mathrm{A}$	$5.16 \pm 1.15 \mathrm{A}$	$40.24 \pm 12.86 \text{A}$
Cel	160	$3.77\pm0.86\mathrm{Bb}$	0.55 ± 0.21 Aa	$1.85\pm0.46\mathrm{Aa}$	$1.92\pm0.44\mathrm{Cb}$	$6.31 \pm 0.16 \mathrm{A}$	$23.21 \pm 4.56\mathrm{C}$
水旱	280	$5.17 \pm 0.56 \mathrm{Bab}$	0.17 ± 0.22Ba	$2.18\pm0.46\mathrm{Aa}$	$2.99 \pm 0.17 \mathrm{Bb}$	$5.07\pm0.57\mathrm{B}$	$37.20 \pm 2.76 \mathrm{B}$
轮作	1 000	7.29 ± 1.18A	$0.05 \pm 0.06B$	$3.16\pm0.52\mathrm{A}$	$4.13\pm0.75\mathrm{A}$	$4.10\pm0.65\mathrm{B}$	$50.10 \pm 7.72 \mathrm{A}$
	1 500	$6.75 \pm 1.59 \mathrm{AB}$	$0.02\pm0.03\mathrm{B}$	$2.80 \pm 1.10 \mathrm{A}$	$3.95\pm0.50\mathrm{A}$	$4.62\pm0.68\mathrm{B}$	$46.16\pm6.68\mathrm{AB}$

1)不同大写字母表示不同围垦年限间数值差异显著(P<0.05),不同小写字母表示不同土地利用间数值差异显著(P<0.05)

表3 围垦时间和土地利用的主效应分析

Table 3	Main	effects	analysis	of re	eclamation	time	and	land	1186
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项目		POC	POCD	MAOC	MAOCD	SOC	SOCD	SIC	SICD	CSP	CSL
围垦年限	F	3.29	2.70	19.88	20.23	8.16	7.33	32.26	32.39	15.26	7.08
	P	0.017	0.039	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
土地利用	F	13.51	13.24	8.60	8.65	10.56	9.40	6.52	11.19	1.65	3.24
	P	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.189	0.025
围垦年限×土地利用	F	1.58	1.11	1.67	0.95	3.63	2.27	2.77	2.13	3.06	1.13
	P	0.195	0.384	0.173	0.463	0.006	0.057	0.024	0.072	0.024	0.368

POC分配比例参照文献[4],应属于中等水平.本研究中POC部分损失较大,这与Lugato等^[3]研究的结果相似,可能是区内频繁耕作条件下,作物根系的较弱水平和植物残余添加的相对快速周转时间使得颗粒态有机物难以在团聚体中受到物理性封闭保护,导致土壤POC累积量较低^[13].

MAOC被认为是长期稳定的有机碳库组分,是有

机物形成的最终较稳定产物,周转时间从几年到几 千年不等,与微生物群落结构和组成变化密切相 关^[27].具有较高MAOC/SOC值的土壤有机质活性低, 不易被生物所利用.MAOC是与土壤中与粉黏粒紧 密结合的有机碳组分,其含量有随着土壤粒径减小 而增加的趋势.本研究中MAOC分配比例较高,说明 土壤有机碳库以MAOC为主,原因可能是区内土壤

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质地以壤土、砂壤土和黏壤土为主,黏粒对有机质的 保护能力强^[26],且土壤养分含量高,有利于植物生 长,同时围垦过程中微生物代谢分泌物所增加的 MAOC含量始终未达到饱和水平,促进了 MAOC 的稳 定和积累.从数值看来,本研究中MAOC/SOC值中位 数表现为:光滩(52.6%) <水旱轮作(55.2%) <林地 (55.3%) < 草地(61.8%) < 旱地(68.1%),均低于世 界土壤平均水平(0.86~0.89)^[28]. 死亡的植物及颗 粒态有机物容易被微生物获取和分解,它们在土壤 系统中停留的时间越长,就越可能经历物理化学转 化,被稳定在团聚体中并通过配位交换、阳离子键桥 和络合作用等关联形式与矿物表面相结合[29,30],在这 一稳定过程中,有机质可能在低生物利用率和长周 转时间条件下更强烈地吸附在土壤矿物上[31],提高 自身抗性,由POC转化为MAOC,这可进一步解释本 研究中MAOC/SOC相较POC/SOC更大的增长幅度. 3.2 围垦后土壤碳组分的动态变化

SOC在土壤中的长久固存受到土壤发育过程的 影响,土壤围垦时间是影响区内MAOC动态的最主 要因素(表3).未围垦的土壤母质受到外界的扰动 小,接受了长江中上游带来的较多有机质,同时河流 过境期间生物圈POC转移和埋藏较多^[32],MAOC分配 比例相对较小.围垦60年间(0~60 a)SOC含量的下 降可能与较高的SOC初始值以及土地利用方式改变 有关,耕作措施加速了原有SOC的消耗,而耕作初期 形成的作物根茬等还不能补偿SOC的消耗,导致其 含量有不同程度的下降.60~1000 a围垦时间的土 壤,SOC含量随着开垦持续时间的延长而增加,一方 面, 土壤成土过程伴随着有机质的积累以及黏粒含量的增加, 作为无机胶结物质的黏土矿物和铁铝氧化物含量的增加均能提升土壤对SOC的吸附, 促进MAOC积累.另一方面, 区内土壤条件发生变化, 土壤地上和地下生物量增多, 营养元素含量增加, 提高了POC和MAOC含量, 期间SOC含量的提高又有利于构建生态酶体系促进SOC稳定性提高^[33], 因此SOC含量积累最大值在280~1000 a出现.围垦年限更长(1000~1500 a)的土壤, POC和MAOC含量均有一定程度的下降, 相较于SOC快速累积的前中期, 此时土壤出现容重增加、有机质减少等退化现象, 同时SOC稳定机制随土壤酸度增强可能发生改变^[34], 碳含量随之下降.

土壤发育时间同样是影响区内 SIC 动态的最主要因素(表 3). 未围垦区 SIC 主要来自上游冲积物中富含无机碳的原生、次生碳酸盐及生物碎屑^[35]. 围垦初期(0~60 a)SIC 含量与土壤 pH、氧化钙同步快速下降,随着围垦年限的增加,SIC 淋失的速度随其含量降低而放缓,围垦 1 000 a 左右土壤中仍有一定量的 SIC 存在,其存在形态可能是较稳定的淡水钙质生物的碎屑. 围垦 1 000 a 以上时土壤中 SIC 含量已经趋近于零.

3.3 土地利用方式对土壤碳分布的影响

SOC含量是有机质分解与生成动态平衡的结果^[16].未围垦区主要是由光滩、草地和林地构成的湿地,枯水季节出露水面,丰水季节常被淹没.光滩应该是最新的沉积,该区的草地和林地覆盖了稍早的长江冲积物,经过一段时间的植被生长,草地和林地的SOC总量有了一定程度的提高^[36],POC和MAOC含量均高于光滩.围垦区林地SOC在土壤发育时间序列上呈现持续的上升趋势,而旱地、水旱轮作地SOC含量的波动变化可能与受到人为干扰有关,农业管理中施肥、灌溉、翻耕等措施在稳定土壤结构、提高农田土壤有机质含量方面起到了积极作用^[37-39],尤其水旱轮作土壤的淹水条件促进了有机质积累^[40],秸秆还田措施又相对促进了土壤POC积累,表现出更高的SOC含量.

研究区土壤由富含碳酸钙的长江冲积物发育而成,因此未围垦区 SIC 含量明显高于其他土地利用类型.随着围垦年限的增加,碳酸盐逐渐淋溶损失,矿质肥料的施用还会降低土壤 pH值,加速碱性阳离子迁移和 SIC 溶解.围垦千年以上时,表层土壤中 SIC 基本淋失殆尽(图2).

本研究区的生物气候条件下,表层土壤中SIC逐 渐淋失的趋势不可逆转,土壤中碳的固存依赖于SOC 数量和质量的提升.区内土壤碳饱和水平远离饱和 度,并低于我国水稻土63.5%的碳饱和平均水平[41], 说明在当前的管理利用下,SOC和MAOC仍有继续积 累的潜力.然而,本文研究表明围垦早期SOC、POC 和 MAOC 总体上都呈现逐渐上升的趋势. 在本区水 旱轮作这种最主要的利用方式下,围垦1000 a后土 壤碳饱和水平达到50.1%,但到1500a时又有所下 降. 区内成土母质为富含各种速效养分的长江冲积 物,围垦后具有较高的肥力和较高的生物量,有利于 SOC的形成和积累. 在过去防洪能力较弱的情况下, 沿江地区不断受洪水的侵袭,土地间歇性地被淹没, 土壤因为持续接受长江输送的新鲜养分而保持较高 肥力.随着长江大堤的加高加固,防洪能力不断提高 的同时,长江对于沿江平原土壤养分的持续供应能 力也因而终止.在这种情况下,如果土地利用强度过 高且人为补充的养分不足,则会导致土壤肥力下 降[42,43],其后果除了土地产出生物量的降低,还可能 导致 SOC 的丧失和土壤碳饱和水平的降低. 当地相 关部门在无为市的调查结果也表明,在测土配方施 肥区,土壤SOC含量普遍得到提高,而在肥料投入不 足的情况下,土壤速效养分和SOC含量出现明显降 低现象[44,45].因此,为保持和提高区内农田土壤的固 碳潜力,需要通过定期的土壤监测,根据养分的实际 消耗给予补充,避免因土壤养分亏损而导致的生物 量下降和 SOC 水平降低.

(1)长江下游平原区土壤有机碳在经历逐渐上升的阶段后,长期耕种的农田土壤其含量又有所下降,围垦1000 a和1500 a后含量均值相较未围垦区上升约28.9%和4.9%.无机碳初始含量占总碳量的20.7%左右,围垦后逐渐下降,1000 a后仅占2.0%以下.

(2)不同利用方式下,矿物结合态有机碳含量均高于颗粒态有机碳,MAOC/POC值表现为:水旱轮作<林地<旱地,总有机碳积累主要通过较稳定的矿物结合态有机碳的增加而实现,矿物结合态有机碳的增加而实现,矿物结合态有机碳的贡献率达48.0%~79.7%.

(3) 区内土壤碳饱和水平较低,仅18.6%~ 56.1%,当前土壤固碳仍有较大潜力,各利用方式下固碳潜力表现为:光滩 < 草地 < 旱地 < 林地 < 水旱轮作.

(4)沿江平原区土壤经长期利用后易出现肥力 质量下降,从而导致生物量减少、有机碳含量和土壤 固碳能力下降.

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结论

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