新施 静 夏 (HUANJING KEXUE)

ENVIRONMENTAL SCIENCE

第 45 卷 第 1 期 2024 年 1 月 15 日

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珠江河口地表水锰氧化物对磷的"载-卸"作用

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摘要:地表径流输送的磷在河口水环境营养结构平衡和初级生产力调节中起到了关键作用.通过测定不同季节珠江下游及河口 地表水阴阳离子、总磷及其形态、颗粒态铁(PFe)、颗粒态锰(PMn)和颗粒态铝(PAl)等理化性质,探究咸-淡水交互区内磷的时 空分布特征,并识别控制磷迁移转化的关键因素.结果表明,ρ(TP)(28.88~233.68 μg·L⁻¹)受到沉降和稀释作用随盐度升高呈 下降趋势,且占比大小为:溶解态无机磷(DIP,37.3%)>颗粒态无机磷(PIP,22.7%)>溶解态有机磷(DOP,21.0%)>颗粒态有 机磷(POP,19.0%).PIP与PFe、PMn和PAl有显著(P<0.05)的共迁移相关关系,水体盐度的升高促进了悬浮颗粒态无机磷的解 吸,且该过程主要发生在咸-淡水界面附近.无机磷的固-液分配系数(K_a)与盐度间的显著正相关关系(P<0.001)表明,高盐度水 域 PIP 主要以更稳定的结合形态存在.K_a与 PMn的显著正相关关系(P<0.01)证明了 Mn氧化物对磷的"载-卸"作用:从淡水环境 中吸附携载无机磷,迁移至高盐度水体释放.

关键词: 咸-淡水交互区; 盐度梯度; 磷; 赋存形态; 迁移转化

中图分类号: X142 文献标识码: A 文章编号: 0250-3301(2024)01-0173-08 DOI: 10. 13227/j. hjkx. 202301010

"Load-Unload" Effect of Manganese Oxides on Phosphorus in Surface Water of the Pearl River Estuary

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Abstract: Phosphorus (P) conveyed by surface runoff plays an essential role in regulating nutrient balance and primary production in estuarine waters. In this study, basic physiochemical properties, total phosphorus (TP, including speciation), particulate iron (PFe), particulate manganese (PMn), and particulate aluminum (PAI) of the surface water in the Pearl River Estuary (PRE) in different seasons were determined to investigate the spatiotemporal distribution characteristics of P and to identify the crucial factor controlling P migration and transformation in the freshwater-saltwater interaction zone. TP concentrations (28, 88-233, 68 µg; L⁻¹) decreased with increasing salinity gradient owing to deposition and dilution. The proportions of P speciation followed a decreasing order as dissolved inorganic phosphorus (DIP, 37, 3%) > particulate inorganic phosphorus (PIP, 22, 7%) > dissolved organic phosphorus (DOP, 21, 0%) > particulate organic phosphorus (POP, 19, 0%). PIP was positively related to PFe, PMn, and PAI (P < 0.05), confirming their concurrent migration behaviors. In addition, the increase m salinity promoted the desorption of phosphate on the suspended particulate matters, which mainly took place near the freshwater interface. A significant positive correlation (P < 0.001) between the solid-liquid phase partitioning coefficient (K_a) of phosphate and salinity indicated that PIP was present mainly in more stable forms in the brackish water. Most importantly, a better relationship between K_d and PMn (P < 0.01) supported our scientific hypothesis of the "load-unload" effect of Mn oxides on P: particulate-carrying phosphates transported from the freshwater scient is for the brackish water.

Key words: freshwater-saltwater interaction zone; salinity gradient; phosphorus; speciation; migration and transformation

磷是生物体遗传代谢、能量传输和细胞膜合成 等过程所必需的生源要素,在调节水体初级生产力 方面具有关键作用^[1-3].过量的磷进入水体会引发水 体富营养化进而导致水华或赤潮等现象^[4],然而水体 的磷浓度过低亦可能致使鱼类减产或破坏浮游生物 的群落结构^[5-7].维持适宜的磷浓度对于水生态系统 的健康和稳定性具有重要意义.然而,珠江三角洲快 速的城市化进程与人口膨胀增加了珠江河口的水环 境负担^[8],导致赤潮频发^[9,10],而磷是藻类生长的主要 限制元素^[11,12].

河流输送的磷是河口主要的磷来源方式,其中 铁(Fe)在磷迁移过程中起到"传输带"作用:无机磷被 悬浮颗粒物上的Fe吸附或结合,运移至河口并沉降 于沉积物中,在缺氧条件下的Fe氧化物还原溶解过 程中被释放^[13,14]. 高强度 SO₄²⁻还原环境产生的二价硫 化物 $(\sum S^{2-})$ 可以与 Fe 反应,形成非溶解性 FeS_{*}并降 低其对无机磷的吸附能力^[15-17]. 由于环境中 Fe 的含 量往往远高于锰(Mn),以往的研究更多关注 Fe 循环 对水环境中磷循环的影响^[18-24];然而,Mn氧化物具有 低 Zeta 电位、高比表面积和高表面电荷数,对水体中 无机磷迁移转化行为的影响不容忽视^[25-27]. 此外,铝 (Al)的氢氧化物对磷也有良好的吸附亲和力^[28,29]. 华

收稿日期: 2023-01-03;修订日期: 2023-03-10

基金项目: 国家自然科学基金项目(42077376,41961144027);中国 科学院青年创新促进会项目(2022352);中国科学院退化 生态系统植被恢复与管理院重点实验室开放基金项目 (VRMDE2202)

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南地区富含Fe、Mn和Al的土壤在降水冲刷作用下进 入水环境可以影响磷的迁移转化行为[30,31]. 珠三角城 市河流中发现了磷和Mn之间紧密的地球化学循环 过程[15],然而珠江河口却主要是铁-磷耦合的迁移机 制^[13,32]. 这种差异可能源于磷在河流输送的过程中 发生的形态转化.关于河口水体磷形态的研究已多 有报道^[33~36],但是鲜有研究区分Fe、Mn和Al在磷迁 移过程中的作用.由于海水中Mn氧化物对无机磷的 吸附亲和力低于Fe氧化物和Al氢氧化物^[29,37],本文 提出咸-淡水交互区常量金属 Mn对无机磷的输送过 程起到关键"载-卸"作用的科学假设:Mn氧化物吸附 富集淡水环境的无机磷在高盐度水体中释放.基于 此,本文以珠江河口咸-淡水交互区为研究区,探索地 表水磷及其形态的时空变化规律,揭示沿盐度梯度 常量金属(Fe、Mn和Al)对磷迁移转化的影响,研究 结果可以提升河口地区磷循环的认知,以期为珠江 河口削减磷负荷提供科学依据.

1 材料与方法

1.1 研究区概况与样品采集

珠 江 流 域 (112°14′~115°53′E, 21°31′~ 26°49′N)由西江、北江、东江和珠江三角洲诸河组 成,流域面积453 690 km²,多年平均年径流量3.38× 10¹¹ m³、属于热带和热带季风气候,降水主要集中在 4~10月.各水系交汇于珠江三角洲平原,形成散射 状水网入海,其中东部4个入海口虎门、蕉门、洪奇 沥和横门的水流汇入伶仃洋,潮汐属于不正规半日 潮(图1).



Fig. 1 Sampling sites for the surface water in the Pearl River Estuary

1.2 样品采集与现场测定

如图1所示,沿盐度梯度在珠江口干流与伶仃洋 水域设置5个采样点位:①中山大学码头(ZD)位于珠 江前航道;②海鸥岛(HOD)是冲积而成的内河岛屿, 岛上以种养业为主;③舢板洲(SBZ)位于广州南沙港 附近;④淇澳岛(QAD)北部是红树林保护区;⑤桂山 岛(GSD)主要发展旅游业,也是重要的网箱养殖基 地.本研究于2021年4月初旱季与7月中旬雨季,在 高高潮、低低潮以及高低潮或低高潮时段采集地表 水样品(在不同潮位时段采集以增加样品的代表 性),用0.45 μm孔径醋酸纤维滤膜(提前烘干、称 重)分离溶解态物质和悬浮颗粒物(SPM).采用便携 式多功能分析仪(HQ-40d;Hach,Loveland,USA)现场 测定 pH、盐度、温度和溶解氧(DO).采用酸碱滴定 试剂盒(Mquant Alkalinity Test;Merk,USA)现场测定 重碳酸根(HCO₃⁻).

1.3 实验分析方法

地表水溶解阴离子 CI⁻和 SO₄²⁻浓度用离子色谱仪 (ICS-900; Dionex, USA)测定.Na⁺、K⁺、Ca²⁺、Mg²⁺等溶 解阳离子浓度用电感耦合等离子体原子发射光谱仪 (ICP-AES; IRIS, USA)测定.溶解态无机磷(dissolved inorganic phosphorus, DIP)使用磷钼蓝法测定^[38],溶解 态总磷(dissolved total phosphorus, DTP)经过硫酸钾消 解后用磷钼蓝法测定,溶解态有机磷(dissolved organic phosphorus, DOP)为 DTP和 DIP之差.

醋酸纤维滤膜上的悬浮颗粒物样品在60℃下烘 干至恒重后计算其质量.用陶瓷刀将滤膜对半分开, 一半滤膜用1 mol·L⁻¹盐酸浸提24 h测定颗粒态无机 磷(particulate inorganic phosphorus, PIP),另一半滤膜 用 H₂SO₄-HClO₄ 消解后测定颗粒态总磷(particulate total phosphorus, PTP)、颗粒态铝(particulate aluminum,PAI)、颗粒态铁(particulate iron, PFe)和颗 粒态锰(particulate manganese, PMn)等.上述浸提液 与消解液以2,4-二硝基苯酚为指示剂用4 mol·L⁻¹ NaOH调节至pH≈3后,用磷钼蓝法测定其中磷含量. 颗粒态有机磷(particulate organic phosphorus, POP)为 PTP和PIP之差,总磷(total phosphorus, TP)则为DTP 和PTP之和.

1.4 数据分析方法

本研究涉及的阴阳离子数据经过离子电荷平衡 计算验证.地表水总溶解性固体(total dissolved solid, TDS)为所有溶解阴阳离子质量浓度总和.采用 IBM SPSS Statistics 22 进行数据统计分析,如果参数符合 正态分布,选用 Pearson 相关及线性相关分析参数间 的相关关系;否则,采用非参数 Spearman 进行相关分 析.采用非参数用 Origin 2021 描绘 Piper 三线图和 Gibbs图,拟合水质理化参数间的二元回归方程.

2 结果与讨论

2.1 地表水基本理化性质

如表1所示,研究区地表水pH范围为7.19~ 8.03,平均值为7.59,呈弱碱性.盐度范围为0.14‰ ~21.12‰,属于淡水(<0.5‰)或微咸水(0.5‰~ 30‰),其中HOD点位在旱季和雨季分别属于微咸水 与淡水;此外,SBZ、QAD和GSD点位在雨季时受地 表径流的稀释作用较为明显. ρ (DO)范围为4.25~ 8.60 mg·L⁻¹,旱季时的水体 DO浓度总体上高于雨 季;其中,ZD点位的水体 DO浓度显著低于其他点位 (P < 0.05),可能与城市排放的污染物在分解过程中 消耗了大量溶解氧有关. ρ (SPM)范围为5.27~ 79.23 μ g·L⁻¹,雨季时受降水冲刷作用影响高于 旱季.

表1 珠江河口地表水基本理化性质¹⁾

季节	采样点	$_{\rm pH}$	温度/℃	盐度/‰	$ ho(\mathrm{DO})/\mathrm{mg}\cdot\mathrm{L}^{-1}$	$ ho(\mathrm{SPM})/\mathrm{mg}\cdot\mathrm{L}^{-1}$
旱季	ZD	8.03 ± 0.10a	$24.18 \pm 0.98a$	$0.18\pm0.00\mathrm{e}$	$4.25\pm0.78\mathrm{b}$	$9.59 \pm 4.60 \mathrm{ab}$
	HOD	$7.47 \pm 0.09 \mathrm{b}$	$23.88 \pm 1.15a$	$3.35\pm0.39\mathrm{d}$	$7.59\pm0.67a$	$56.11 \pm 43.01a$
	SBZ	$7.60 \pm 0.15 \mathrm{b}$	$25.03 \pm 0.67a$	$8.37\pm0.91\mathrm{c}$	$8.18 \pm 0.14 \mathrm{a}$	$10.83 \pm 12.68 \mathrm{ab}$
	QAD	$7.59 \pm 0.64 \mathrm{b}$	$25.00\pm0.63a$	$11.52\pm0.20\mathrm{b}$	$8.60 \pm 0.53a$	18.93 ± 11.16 ab
	GSD	$7.52\pm0.08\mathrm{b}$	$23.10\pm0.66a$	$21.12 \pm 0.03a$	8.51 ± 0.22a	$5.27 \pm 1.90 \mathrm{b}$
雨季	ZD	$7.51 \pm 0.15 \mathrm{ab}$	$31.00 \pm 1.55a$	$0.14\pm0.01\mathrm{e}$	$4.68 \pm 1.15 \mathrm{e}$	22.88 ± 2.24ab
	HOD	$7.52 \pm 0.21 \mathrm{ab}$	$32.27 \pm 1.94a$	$0.23\pm0.00\mathrm{b}$	$5.82 \pm 0.19 \mathrm{b}$	79.23 ± 48.60a
	SBZ	$7.19\pm0.11\mathrm{b}$	$30.30 \pm 2.07a$	3.64 ± 1.21ab	$6.62 \pm 0.50 \mathrm{ab}$	$26.31 \pm 8.01 \mathrm{ab}$
	QAD	$7.37\pm0.19\mathrm{ab}$	$29.80 \pm 2.05a$	4.92 ± 1.61ab	$6.37 \pm 0.12 \mathrm{ab}$	28.47 ± 19.01 ab
	GSD	$7.93 \pm 0.12a$	29.27 ± 1.12a	17.96 ± 1.11a	$7.35 \pm 0.34a$	$11.50 \pm 6.34 \mathrm{b}$

1)不同小写字母表示不同点位存在显著性差异(P<0.05)

2.2 水化学特征

如图 2 所示,研究区地表水 ρ (TDS) 范围为 270 ~ 37 500 mg·L⁻¹. GSD、QAD、SBZ 和 HOD(旱季) 点 位的 ρ (Cl⁻)/ ρ (Cl⁻+HCO₃⁻)和 ρ (Na⁺)/ ρ (Na⁺+Ca²⁺)均接 近 1,说明这些点位的地表水主要受蒸发浓缩作用 控制,受海水的影响;而 ZD 和 HOD(雨季)点位为陆 地流域水岩作用控制,可被视为淡水水体^[39]. HOD 点位在雨季和旱季时分属不同水化学类型,说明该 点位于咸-淡水交互作用的界面附近,理化性质取 决于径流与潮汐作用的相对强度.该点位的水流方 向可能随潮汐涨落出现相反的变化,而这种复杂多 变的水动力条件不利于悬浮颗粒物的沉降,导致地 表水中高浓度的悬浮颗粒物(表1).综合基本理化 性质及水化学特征将采样点划分为3类:淡水区 (ZD)、咸-淡水界面(HOD)、微咸水区(SBZ、QAD、 GSD).



图 2 珠江河口地表水 Gibbs 图 Fig. 2 Gibbs map of the surface water in the Pearl River Estuary

2.3 珠江河口地表水磷的时空分布特征

珠江河口地表水 $\rho(TP)$ 范围为28.88~233.68 µg·L⁻¹,从淡水区向微咸水区呈现降低趋势(表2).随 水流迁移距离的增加,咸水逐渐稀释了磷浓度;此 外,河口区域的水流趋缓,有利于颗粒态磷的沉降作 用.旱季时,地表水 $\rho(TP)$ (均值:95.93µg·L⁻¹)略低 于雨季(均值:116.18µg·L⁻¹).一方面,雨季的降水增 强了磷的面源负荷,可能是总磷浓度升高的主要原 因;另一方面,径流量增大使咸-淡水界面向河口区推 移,扩大了陆源物质输送的影响范围.如表2所示, 水体中主要的磷形态为DIP(10.98~54.41µg·L⁻¹), 占TP的37.3%,从淡水区到微咸水区呈现显著下降 的趋势(P < 0.05);旱季时,ZD和HOD点位的DIP浓 度高于雨季,而微咸水区点位却低于雨季,这主要是因为旱季采样期间正值浮游植物生长期,大量无机磷被吸收利用^[11,33]. ρ (DOP)范围为9.04~50.80 μ g·L⁻¹,占TP的21.0%;雨季时,DOP从淡水区到咸水区呈现显著的降低趋势(P < 0.05),而旱季无显著空间差异(P > 0.05).水体中的DOP可能与降水导致的面源污染有关,雨季时加剧了这一过程. ρ (PIP)和 ρ (POP)范围分别为2.54~74.67 μ g·L⁻¹和4.39~61.24 μ g·L⁻¹,分别占TP的22.7%和19.0%,雨季时PIP和POP普遍大于旱季,随盐度的升高大致呈现降低的趋势.而雨季时,GSD点位的PIP和POP浓度较QAD点位更高,说明该点位可能存在除径流输入以外的颗粒态磷来源.

表2 珠江河口地表水总磷及磷形态浓度 ¹⁾ /μg·L ⁻¹	
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 $Table \ 2 \quad Concentrations \ of \ TP \ and \ phosphorus \ speciation \ in \ the \ surface \ water \ in \ the \ Pearl \ River \ Estuary/\mug \cdot L^{-1}$

季节	磷形态	ZD	HOD	SBZ	QAD	GSD
	ТР	156.56 ± 23.53a	$139.70 \pm 50.84a$	60.99 ± 12.39b	$70.39 \pm 11.15\mathrm{b}$	$28.88 \pm 4.16 \mathrm{c}$
	DIP	$54.41 \pm 5.22a$	$50.64 \pm 3.32a$	28.12 ± 3.65 b	$24.20\pm3.77\mathrm{b}$	$10.98 \pm 3.85 \mathrm{e}$
旱季	DOP	$18.90 \pm 9.13 \mathrm{a}$	$16.08 \pm 3.84a$	14.92 ± 3.41a	$19.22 \pm 4.07a$	10.97 ± 2.68a
	PIP	$49.97 \pm 15.15a$	$44.45 \pm 29.78a$	9.02 ± 6.86b	$13.02 \pm 2.48 \mathrm{b}$	$2.54 \pm 0.67c$
	POP	33.28 ± 6.14a	$28.53 \pm 15.01 \mathrm{ab}$	$8.92 \pm 4.82 \mathrm{ab}$	$13.95 \pm 1.63 \mathrm{b}$	$4.39 \pm 0.81c$
	TP /	233.68 ± 45.60a	$133.72 \pm 31.16b$	$79.07 \pm 19.08 \mathrm{b}$	$73.08\pm20.49\mathrm{b}$	$61.33 \pm 14.02\mathrm{b}$
雨季	DIP	$46.97 \pm 15.47a$	$33.13 \pm 2.80 \mathrm{ab}$	40.71 ± 2.53ab	35.15 ± 4.27 ab	16.29 ± 5.63e
	DOP	50.80 ± 13.42a	$36.89 \pm 2.87 \mathrm{b}$	$17.54 \pm 11.11 \mathrm{bc}$	$9.04 \pm 1.06 \mathrm{bc}$	9.25 ± 5.73c
	PIP	74.67 ± 17.96a	32.70 ± 20.26 ab	12.09 ± 2.42 ab	$14.97 \pm 8.73 \mathrm{ab}$	$20.99 \pm 6.70 \mathrm{c}$
15	POP	61.24 ± 13.76a	$30.99 \pm 11.23b$	$8.72 \pm 5.70 \mathrm{b}$	$13.91 \pm 8.31\mathrm{b}$	$14.79 \pm 5.30 \mathrm{b}$
1) 不同小馆字码表示不同点位方在显著研教员(D>0.05)						

珠江口水体磷主要以溶解态(41.9%~75.5%) 的形式进行迁移,不同点位的DTP和PTP的占比有明 显差异,表明各形态的磷在不同水动力、水环境条件 下的迁移转化过程迥异(图3).从淡水区点位ZD到 微咸水区点位SBZ,水动力条件减弱,促进了PTP沉 降,导致其比例降低;而在微咸水区(除雨季GSD 点),DTP的比例随盐度没有明显的变化.GSD点位雨 季的PTP比例显著大于旱季(P<0.05),与SPM浓度 的季节变化特征一致(表1).这可能与桂山岛的人类 活动污染排放有关,而降水事件则进一步加剧了GSD 局部的陆源磷负荷.

由于各采样点受到的海水稀释作用差异,磷与 盐度的关系曲线可以更准确地反映其在咸-淡水交互 区的迁移行为(图4).如图4(a)所示,TP浓度沿盐度 梯度呈现负对数下降趋势,在盐度小于5‰的区域大 幅减少,说明TP除受到咸水的稀释作用外还发生了 明显的沉降过程.在径流、潮汐和咸淡水混合等因 素的作用下,可能出现滞流或往复流现象,使悬浮颗 粒物在原位大量沉降,并在口门区附近形成拦门





沙^[40].例如海鸥沙,就是近代珠江三角洲河网区快速 沉积形成沙坝,而后进一步发育扩大^[41].TP浓度沿盐 度梯度分布的季节差异不明显.地表水中不同形态 的无机磷均随盐度升高而下降(图4);溶解态磷(DIP 和 DOP)主要受到咸水和淡水混合后的稀释作用,而 颗粒态磷(PIP和POP)受到稀释作用的同时在沿程上 发生沉降.

2.4 常量金属元素对无机磷迁移过程的影响

如表 3 所示,旱季时,磷及其各形态均与 PFe、 PMn和PAl呈显著正相关(P < 0.05);而雨季时,仅发 现 TP、PIP和 POP与 PMn之间的显著正相关关系 (P < 0.05).这说明磷的迁移、沉降过程与 PFe、PMn 和 PAl密切相关.由于悬浮颗粒物主要来源于面源 负荷,因此受雨季降水增强的影响,出现磷与常量金 属相关关系的部分解耦^[42].此外,PMn与磷及其形态 之间的相关系数较 PFe和 PAl更高(表 3),表明 PMn 在磷的迁移过程中可能发挥了更重要的作用.

表3 不同季节磷(形态)与颗粒态铁、颗粒态锰 和颗粒态铝相关性分析¹⁾

Table 3 Correlation analysis between phosphorus (including speciation), PFe, PMn, and PAl in different seasons

		-		
季节	磷(形态)	PFe	PMn	PAl
	TP	0.721^{**}	0.889^{**}	0.736**
	DIP	0.475^*	0.669^{**}	0.504^*
旱季	DOP	0.548^*	0.598^{**}	0.542^*
	PIP	0.738**	0.905^{**}	0.754**
	POP	0.730**	0.907^{**}	0.738**
	TP	0.046	0.618^{*}	0.032
	DIP	0.118	0.232	0.136
雨季	DOP	-0.350	0.268	-0.371
11	PIP	0.214	0.839**	0.186
3	POP	0.089	0.754**	0.061

1)*和**分别表示在0.05和0.01水平上的显著相关

如图 5 所示, PIP 与 PFe、PMn 和 PA1之间在淡水 区、咸-淡水界面和微咸水区都存在显著正线性相关 关系(P < 0.05),说明悬浮颗粒物上的 Fe、Mn 和 A1 与无机磷在咸-淡水交互区的吸附-解吸过程有关,证 实了 Fe、Mn 和 A1携载磷的理论假设.不同水域 PFe、PMn和PAl与PIP 拟合曲线斜率均呈现相同大 小顺序:淡水区>咸-淡水交互区>微咸水区(图5). 这表示在盐度越高的水域,等量的PFe、PMn和PAl 携载的无机磷越少,主要原因为高浓度阴离子的竞 争吸附作用促进了吸附态无机磷的解吸.而且从淡 水区到咸-淡水交互区的拟合曲线斜率变幅(2.4~ 5.7倍)远大于从咸-淡水交互区到微咸水区(1.3~ 1.6倍),说明PIP的解吸过程主要发生在颗粒物通过 咸-淡水界面的过程.当悬浮颗粒物通过咸-淡水界 面进入盐度更高的水域时,无机磷的释放量相对较 少,可能是因为此时PIP主要以更稳定的Fe、Mn或Al 结合态的形式存在.

固-液分配系数表征地表水中PIP与DIP之间的 配分关系,可以反映悬浮颗粒物上无机磷的吸附-解 吸特征:

$$K_{\rm d} = \frac{\rho(\rm PIP)}{\rho(\rm DIP) \times \rho(\rm SPM)}$$

式中, K_a 为无机磷的固-液分配系数(L·mg⁻¹), ρ (PIP) 为珠江口地表水 PIP的质量浓度(μ g·L⁻¹); ρ (DIP)为 珠江口地表水 DIP的质量浓度(μ g·L⁻¹); ρ (SPM)是珠 江口地表水 SPM的质量浓度(μ g·L⁻¹); ρ (SPM)是珠 江口地表水 SPM的质量浓度(ng·L⁻¹).如图6所示, 旱季时淡水点位ZD的 K_d (0.101 L·mg⁻¹)高于雨季(均 值:0.074 L·mg⁻¹). 悬浮颗粒物能起到"缓冲"作用: 当水体 DIP被降水稀释时,悬浮颗粒物能够释放 PIP, 维持固-液相间 DIP的动态平衡^[43,44]. 咸-淡水界面点 位 HOD 的 K_d (0.007 ~ 0.030 L·mg⁻¹)均低于 ZD (0.053 ~ 0.130 L·mg⁻¹),也证实了 PIP在通过咸-淡水 界面时被大量释放.在微咸水水域(盐度 > 0.5‰), K_d 与盐度存在显著正相关关系(P < 0.001),而 DIP随 盐度的升高而减小(图4),说明微咸水区内 PIP的稳



图4 珠江口地表水磷及其形态与盐度的关系

Fig. 4 Relationship between P and salinity of the surface water in the Pearl River Estuary





定性随盐度而升高,以更稳定的Fe、Mn或Al结合态存在.

Fe、Mn和Al氧化物或氢氧化物能通过物理吸附、络合反映或化学结合的形式与磷一起迁移,然而珠江河口PFe、PMn和PAl对PIP吸附-释放特征的影响程度迥异.如图7所示,ZD点到QAD点的PMn与 K_a 相关关系(R=0.93, P < 0.001)较PFe(R=0.53, P < 0.01)和PAl(R=0.46, P < 0.05)更显著,说明无机磷的吸附-释放过程与Mn有更紧密的联系.天然Mn氧化物通过静电吸附或络合反应的机制吸附无机磷,但是这种反应是完全可逆的^[18,26,45].尤其在咸水环境





中,钙和镁离子可以增强阴离子和锰氧化物表面之 间的相互作用,一定程度上降低锰氧化物对无机磷 的吸附^[37].相对而言,Fe氧化物表面的吸附点位在 pH>7的海水中对无机磷的吸附亲和力都高于Mn氧 化物,而且Al的氢氧化物对无机磷的吸附选择性远 高于 Cl⁻和 SO₄⁻等其他阴离子,因此 Fe 和 Al 携载的 PIP在高盐度环境中解吸释放的可能性更小.那么 通过咸-淡水界面后,PIP沿盐度升高梯度的释放 要来源于 Mn氧化物,证实了 Mn在无机磷迁移过程 中起到了关键的"卸-载"作用. PIP与 PFe和 PAI之间 的显著正相关关系(P<0.05)说明Fe和Al是无机磷 的良好载体(图5),但是在咸水环境依然保持相对稳 定,不具备"卸"磷的作用. 华南地区富锰的土壤条件 放大了水环境 Mn氧化物对磷迁移的影响,相似的情 况也在福建九龙江河口发现^[30,31].在GSD点,K,仅与 PMn具有显著相关关系(P<0.01),且拟合曲线偏离 其他点位(图7),这说明PIP除了珠江主干流输送,也 有一部分来源于桂山岛等周围岛屿.桂山岛与珠三 角污染源性质的差异可能是该点位 Ka与 PFe 和 PAI 解耦的主要原因.

3 结论

河口区的磷主要来源于陆地径流带来的面源负荷,总磷的质量浓度沿着盐度梯度呈下降趋势,各形态占比大小顺序依次为:DIP > DOP > PIP > POP. 悬浮颗粒物中的 Fe、Mn和Al的氧化物或氢氧化物吸附并携载无机磷向河口迁移.微咸水环境中PMn是无机磷吸附-解吸的关键控制因素,受阴离子竞争吸附作用而解吸.这证明了Mn氧化物在咸-淡水交互区无机磷迁移过程中起着重要的"载-卸"作用:将淡水环境中吸附的无机磷携载至高盐度水域释放.咸水



图 7 珠江口地表水无机磷的固-液分配系数与颗粒态的 Fe、 Mn和 Al质量分数相关性

Fig. 7 Relationships between solid-liquid phase partitioning coefficient (K_d) of inorganic phosphorus and the mass percentage concentrations of PFe, PMn, and PAl in the surface water in the Pearl River Estuary

区内,PIP的稳定性沿盐度梯度增大,转为更稳定的 Fe、Mn或Al结合态磷.上述结论验证了本文的理论 假设,并强调了氧化环境中Mn氧化物对于无机磷环 境行为的重要意义.

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