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## 锡林河上游雨季降水、河水和地下水转化关系

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摘要:为深入理解草原内陆河流域的水文循环过程及不同水体转化机制,以锡林河流域大气降水、河水和地下水为研究对象,对其氢氧稳定同位素进行了测试及多尺度时空特征分析,探究流域不同水体间的定量转化关系.结果表明:①锡林河流域具有明显的内陆性半干旱气候特征,大气降水是流域河水和地下水的主要补给源,地下水和河水同时经历了不同程度的非平衡蒸发;②河水同位素组成在季节上表现出春秋贫化、夏季富集的特征,在空间上表现为自上游到下游逐渐升高的趋势;浅层和深层地下水δ<sup>18</sup>0在生长季的波动变化基本一致,二者的主要差异发生在生长季末期,即前者趋于稳定而后者呈上升趋势,反映出深层地下水对大气降水和地表水入渗补给具有滞后响应,在空间上二者均由东南向西北逐渐贫化;③基于端元混合模型的估算结果可知,夏季大气降水和浅层地下水对河水的平均补给比例分别为52.69%和47.31%,说明对于内陆河流域,即使在多雨季节,浅层地下水也是河水的重要补给来源,研究旨在为半干旱典型草原内陆河流域的水资源调控和生态环境保护提供理论指导.

关键词:氢氧同位素;不同水体转化关系; 氘盈余; 端元混合模型; 内陆河流域 中图分类号: X143 文献标识码: A 文章编号: 0250-3301(2023)12-6754-13 **DOI**: 10.13227/j. hjkx. 202211172

# Relationship Between Precipitation, River Water, and Groundwater Conversion in the Upper Reaches of Xilin River During the Rainy Season

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Abstract: To deeply understand the hydrological cycle process and the transformation mechanism of different water bodies in the grassland inland river basin, the atmospheric precipitation, river water, and groundwater in the Xilin River Basin were taken as the research objects, the hydrogen and oxygen stable isotopes were analyzed, and the multiscale spatio-temporal characteristics were analyzed to explore the quantitative transformation relationship between different water bodies in the basin. The results showed that; the Xilin River Basin had an obvious inland semi-arid climate, the atmospheric precipitation was the main source of recharge for the river water and groundwater, and the groundwater and river water experienced different degrees of non-equilibrium evaporation at the same time. ② The isotopic composition of the river water showed the characteristics of depletion in spring and autumn and enrichment in summer and showed a trend of increasing from upstream to downstream in space. The variation in  $\delta^{18}$ 0 in shallow and deep groundwater during the growing season was basically the same, and the main difference between the two occurred at the end of the growing season, that is, the former tended to be stable, whereas the latter showed an upward trend, which reflected that the deep groundwater had a lagged response to the infiltration and recharge of atmospheric precipitation and surface water, and both of them were depleted gradually from southeast to northwest in space. ③ Based on the estimation results of the endmember mixing model, the average recharge ratio of atmospheric precipitation and shallow groundwater to river water in summer was 52.69% and 47.31%, respectively, indicating that shallow groundwater was an important recharge source of river water in the inland river basin even during the rainy season. The results of this study provide theoretical guidance for water resource regulation and ecological environment protection in a typical semi-arid grassland inland river basin.

Key words: hydrogen and oxygen isotopes; transformation relationship of different water bodies; deuterium surplus; end member mixing model; inland river basins

随着同位素水文学的发展,氢氧同位素方法已成为当前水科学研究领域揭示流域水循环过程与机制、阐释气候变化所引起的水循环效应等最为重要的研究方法之一<sup>[1,2]</sup>. D 和<sup>18</sup>O 分别是自然界中氢和氧的两种稳定同位素,流域水循环相变过程或多种水体的混合过程中的同位素分馏作用导致不同"水源"的水体同位素存在差异<sup>[1]</sup>,因此氢、氧稳定同位素被认为是研究水循环过程的理想示踪剂<sup>[3]</sup>. 结合端元混合模型、水文模型、水文地质模型以及水化学质量平衡等方法,氢氧稳定同位素技术被广泛应

用于不同水体转化关系与作用机制等研究[1,4~6].

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前人基于氢氧同位素技术于泰国蒙河流域<sup>[7]</sup>、亚洲中低纬度地区<sup>[8]</sup>、青藏高原地区<sup>[9]</sup>和印度西南地区<sup>[10]</sup>探究了多种水体的时空演化特征及驱动机制,为水体转换和水循环过程提供了大量的基础性研究.后基于同位素又衍生了诸多的不同水体转化估算的方法,如李静等<sup>[11]</sup>和许秀丽等<sup>[12]</sup>分别在长江和黄河流域应用 d-excess 聚类分析和同位素质量平衡模型开展不同水体动态转化的研究,张兵等<sup>[13]</sup>、梁丽娥等<sup>[14]</sup>和胡玥等<sup>[15]</sup>以氢氧同位素为主,结合水化学技术,系统分析了不同水体之间的交互关系.

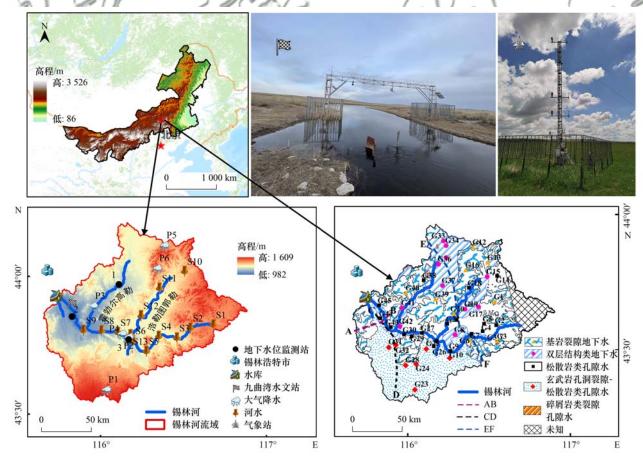
锡林河流域是位于内蒙古高原典型草原区的代表性内陆河流域,但流域受半干旱气候控制,对气候变化极为敏感,近年来极端洪旱事件频发,水资源短缺问题日趋严重.杨璐<sup>[16]</sup>基于大气降水、地表水和地下水同位素特征,对锡林河流域不同水体之间的转化关系做出了定性描述,但流域水文循环过程和不同水体的相互转化关系复杂多变,仍需深入开展相关研究明确流域水体转换的定量关系.因此,本文基于环境同位素技术,分析锡林河流域大气降水、河水和地下水氢氧同位素的时空分布特征,利用同

位素二元线性混合模型,量化大气降水和地下水对河水的贡献,从而揭示锡林河流域不同水体之间的相互关系,以期为深入探究典型草原内陆河流域不同水体的相互转化机制提供理论基础,对于气候变化背景下流域水资源调控和生态环境保护具有重要的现实意义.

#### 1 研究区概况

本文选择锡林河流域上游(116°00′~117°14′ E,43°24′~44°03′N)为研究区(图1). 锡林河发源于内蒙古自治区赤峰市克什克腾旗宝尔图山,流经锡林郭勒盟阿巴嘎旗,在贝力克牧场转向西北流经锡林浩特市,最终注入查干淖尔沼泽自然消失 $^{[17]}$ . 锡林河全长 175 km,河宽1~5 km,其右岸有两条支流汇入,分别为浩勒图郭勒和霍勃尔高勒(多年干涸) $^{[18]}$ . 流域气候类型为温带半干旱大陆性季风气候,年均降水量 278.9 mm,年均蒸发量1862.9 mm ( $\Phi$  20 cm 蒸发皿),年均气温 2.8°C,年均风速 3.4 m·s $^{-1}$ [19].

研究区地下水分布广泛,含水层类型多,依据地下水埋藏条件以及含水层岩性,主要划分为4种类



AB、CD 和 EF 为水文地质剖面带

图 1 研究区位置、采样点分布示意

Fig. 1  $\,$  Study area location and sampling sites distribution diagram

型,即河谷冲积平原区第四系松散岩类孔隙水、河谷冲积平原和低山平原区双层结构类地下水、岩溶台地玄武岩孔洞裂隙-松散岩类孔隙水和低山丘陵区基岩裂隙水,前两类埋藏较浅(<30 m),可归为浅层地下水,后两类埋藏较深(>30 m),可归为深层地下水.地下水径流方向受复杂地形条件影响,但总体与锡林河流向相符,自东向西排泄至河谷地区[16].

#### 2 材料与方法

#### 2.1 样品采集与测试

依据锡林河流域水系特征和水文地质条件(图2),在研究区均匀布置降水采样点6处(P1~P6)、

河水采样点 13 处(S1~S13)和地下水采样点 45 处(G1~G45),于 2017 年 5~10 月进行水样采集,大气降水按次降水事件取样,通过自制集雨器(内含量筒、导水漏斗和乒乓球)人工收集,每次降水结束后立即取回水样置于 100 mL 聚乙烯采样瓶内;地表水沿干流和支流在河道中央水面以下 30 cm 处取样,一个月采样 1 次;地下水样根据监测井及居民用井埋深将其分为浅层和深层,在抽水 15 min 后采集新鲜水样,一个月采样 1 次. 所有水样密封冷藏(4℃)保存,并在 1~2 d 内进行测试分析.实验期共采集有效水样 345 件,其中大气降水样 25 件、地表水样 77 件和地下水样 243 件.

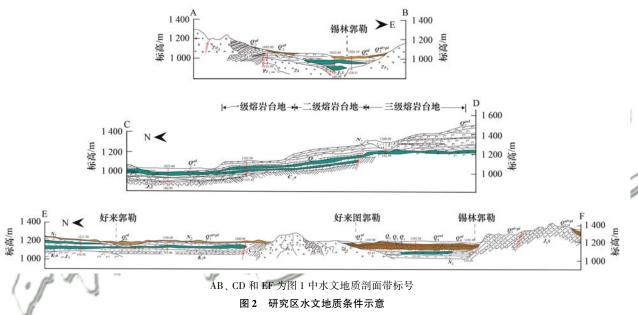


Fig. 2 Hydrogeological conditions in the study area

样品测试采用美国 Los Gatos Reserarch (LGR) 公司生产的液态水稳定同位素分析仪 (LWIA-45-EP),测试样品之前,利用标样 (908-0008-9101: $\delta$ D 为 - 145.82‰, $\delta^{18}$ O 为 - 19.35‰、908-0008-9103: $\delta$ D 为 - 74.88‰, $\delta^{18}$ O 为 - 10.7‰ 和 908-0008-9104: $\delta$ D 为 - 46.88‰, $\delta^{18}$ O 为 - 7.24‰)进行标定,实验过程中用 1 mL 注射器吸取水样,再用前端装有0.2  $\mu$ m 微孔滤头进行过滤,将过滤后的样品移入测试瓶内进行测试,每个样品注入 6 针,去除前两针,取后四针数据进行平均,从而消除记忆效应。 $\delta^{18}$ O和  $\delta$ D 测试精度分别为 ± 0.1‰和 ± 0.3‰,所得结果用相对维也纳标准平均海洋水 (V-SMOW)的千分偏差来表示[20]:

$$\delta(\%e) = \frac{R_{\text{Sample}} - R_{\text{V-SMOW}}}{R_{\text{V-SMOW}}} \times 1000 \tag{1}$$

式中, $R_{\text{Sample}}$ 和  $R_{\text{V-SMOW}}$ 分别为水样中和维也纳标准 平均海洋水中的氧(或氢)与稳定同位素的比值 ( $^{18}O/^{16}O$ 或 D/H).

#### 2.2 数据与方法

#### 2.2.1 数据获取与处理

本研究使用 MODIS 逐月陆地标准合成的地表温度产品(LST,空间分别率为 0.05°×0.05°),获取于 NASA 网站(https://ladsweb. modaps. eosdis. nasa.gov/);气象数据采用研究区自设气象站(116°29′E,43°38′N)连续监测的气温、降水、相对湿度和风速数据.实验数据的统计分析采用 SPSS 26.0 完成,图形绘制采用 R 语言 ggplot2 完成.

#### 2.2.2 线性端元混合模型

基于同位素质量平衡原理,通过对比地表水不同水源的  $\delta D$  和 $\delta^{18}O$ 值,可判断不同时期地表水的补给来源以及转化关系<sup>[12]</sup>,有学者提出三元混合模型和多元混合模型,这些模型在进行n 种水源划分时,需要引入n-1 种示踪剂,但在示踪剂选择上却存在多种限制条件<sup>[4]</sup>,故应用二元线性混合模型估算各端元的混合比值<sup>[5]</sup>,模型如下:

$$\delta_{\rm S} = f_1 \cdot \delta_1 + f_2 \cdot \delta_2 \tag{2}$$

$$\beta_{\rm S} = f_1 \cdot \beta_1 + f_2 \cdot \beta_2 \tag{3}$$

$$f_1 + f_2 = 1 (4)$$

式中, $\delta_{\rm s}$  和  $\beta_{\rm s}$  分别为混合后目标水体中的  $\delta$ D 和  $\delta^{\rm l8}$ O值, $\delta_{\rm l}$  和  $\delta_{\rm l}$  分别为不同补给水源的  $\delta$ D 值, $\beta_{\rm l}$  和  $\beta_{\rm l}$  分别为不同补给水源的 $\delta^{\rm l8}$ O值, $f_{\rm l}$  和  $f_{\rm l}$  分别为不同补给水源的混合比值.

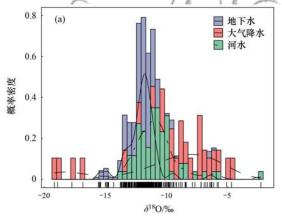
#### 3 结果与讨论

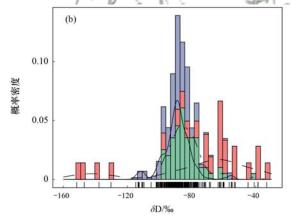
#### 3.1 大气降水、河水和地下水氢氧同位素统计 特征

为了定量评价不同水体同位素特征的差异性, 计算了  $5 \sim 10$  月锡林河流域大气降水、河水和地下 水水样的  $\delta D$  和 $\delta^{18}$ O算术均值、偏度和峰度等统计特 征值(表 1),进一步用高斯分布函数对其进行拟合 (图 3).

根据表1和图3可知:①不同水体中稳定同位素的均值、25%和75%分位数均呈现出地下水<河水<大气降水的规律,证实了地下水由于埋藏较深,不易受到蒸发分馏的影响,使得其稳定同位素较为

贫化且分布更集中,而大气降水受降水过程中二次 蒸发的作用,促使其稳定同位素的富集程度较高.这 与李广等[21]在研究长沙地区不同水体稳定同位素 组成时,发现地下水贫化而降水富集的规律一致; ②大气降水水样稳定同位素值的变化范围和变差系 数均高于其他水体,说明大气降水稳定同位素值的 离散程度比地下水和地表水的更高. 已有学者在洞 庭湖流域[22]、黄土丘陵区[23]和长沙地区[21]也观察 到类似的现象,这是区域大气降水受水汽源地、水 汽输送、相变过程以及水汽的补充和交换等因素综 合作用的结果: ③地下水 $\delta^{18}$ O和  $\delta$ D 的偏度均 > 0 及 峰度均>3,呈正偏厚尾分布,表明地下水同位素偏 贫化且分布较集中;与地下水相比,河水 $\delta^{18}$ O和  $\delta$ D 亦呈正偏厚尾分布,但贫化程度较小且离群值较多; 大气降水 $\delta^{18}$ O和  $\delta$ D 的偏度均 < 0 及峰度均 < 3, 呈负 偏瘦尾分布,表明降水同位素值偏富集且分布较分 散. 不同水体中稳定同位素统计分布特征的差异性 规律,与不同水体的补给来源和其蒸发分馏效应的 衰减规律相关[23]





坐标轴黑色竖线表示河水、地下水及大气降水  $\delta D$  和 $\delta^{18}$ 0出现的频数

#### 图 3 河水、地下水及大气降水 $\delta D$ 和 $\delta^{18}$ O高斯分布

Fig. 3 Gaussian distribution of  $\delta D$  and  $\delta^{18}O$  in river water, groundwater, and atmospheric precipitation

#### 表 1 河水、地下水及大气降水 $\delta D$ 和 $\delta^{18}O$ 描述性统计指标

Table 1 Descriptive statistical indicators of  $\delta D$  and  $\delta^{18}O$  in river water, groundwater, and atmospheric precipitation

| 指标类型 —                       |         |          | 水样类型   |          |
|------------------------------|---------|----------|--------|----------|
|                              |         | 河水       | 地下水    | 大气降水     |
|                              | 均值      | - 10. 74 | -12.01 | -9.98    |
|                              | 25% 分位数 | -12.15   | -13.49 | -11.80   |
| $\delta^{18}\mathrm{O}/\%$ o | 75% 分位数 | -10.41   | -11.45 | -7.06    |
| δ <sup>13</sup> U/%ο         | 变差系数    | -0.18    | -0.11  | -0.42    |
|                              | 偏度      | 2.15     | 0.76   | -1.08    |
|                              | 峰度      | 7.31     | 5.52   | 0.25     |
|                              | 均值      | -83.41   | -89.03 | -74.81   |
|                              | 25% 分位数 | -90.32   | -99.27 | -88.09   |
| SD /64                       | 75% 分位数 | -78.72   | -84.28 | - 52. 97 |
| $\delta \mathrm{D}/\%_o$     | 变差系数    | -0.12    | -0.10  | -0.46    |
|                              | 偏度      | 1.79     | 0.34   | -1.02    |
|                              | 峰度      | 6.81     | 3.10   | 0.36     |

#### 3.2 大气降水 d-excess 值变化特征及其指示意义

大气降水 d-excess 值能够反映降水来源、水汽运移规律以及降水过程中由于动力分馏而偏离平衡分馏的程度<sup>[24]</sup>,受到相对湿度(RH)、风速(WS)、降水量(P)以及温度(T)等气象因素的影响<sup>[25]</sup>.根据 Dansgaard 定义的大气降水 d-excess =  $\delta$ D - 8 ×  $\delta$ <sup>18</sup>O,计算了流域夏季( $\delta$  ~ 8 月)大气降水 d-excess 值(图 4). 研究区夏季大气降水 d-excess 值介于 -1. 42% ~ 16.  $\delta$ 3% 之间,振荡变化较大,表明水汽来源复杂;d-excess 均值为  $\delta$ . 16% ,小于全球大气降水 d-excess 值(10% ),说明流域水汽由低纬度海洋蒸发所形成<sup>[26]</sup>;d-excess 值对季风特性较为敏感<sup>[27]</sup>,研

究区夏季受东南季风影响,大量来自太平洋的水汽气团<sup>[28]</sup>向内陆地区输移,沿途不断冷凝形成降水,由于轻重同位素之间存在分馏速率差异,导致降水气团中重同位素被持续"冲刷",动力分馏作用较强, dexcess 值偏小;进入8月后,大气降水 d-excess 值呈现下降趋势,其主要原因是8月的气温相比7月呈上升趋势,雨期二次蒸发增强导致绝对湿度变大,但同时饱和水汽压随温度的升高而大幅增加,使得相对湿度反而减小,即雨滴在降落过程中由于动力分馏而更易偏离平衡分馏,促使 d-excess 值下降,这与赵明华等<sup>[29]</sup>研究发现 d-excess 值与温度存在负相关、与相对湿度存在正相关的结论基本一致.

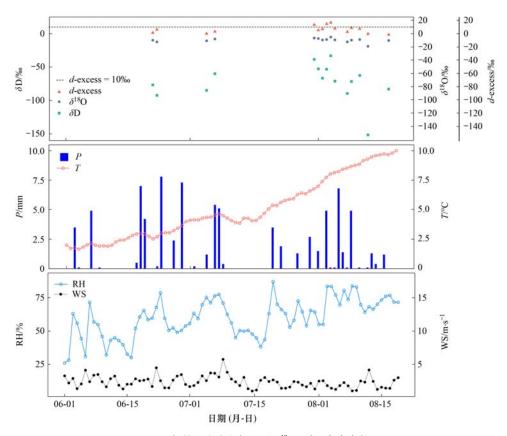


图 4 锡林河流域降水  $\delta D$  和 $\delta^{18}O$ 组成及气象参数

Fig. 4 Composition of  $\delta D$  and  $\delta^{18}O$  in precipitation and meteorological parameters in Xilin River Basin

#### 3.3 河水稳定同位素的时空分布

锡林河干流、支流河水δ<sup>18</sup>O值 5~10 月沿程变化情况如图 5 所示. 空间上,干流的 δD 和δ<sup>18</sup>O组成总体表现出从上游到下游逐渐升高趋势,与锡林河流域地表温度的空间分布(图 6)趋势一致,这主要是由于在降雨下渗或地表径流形成的过程中,均会受到地面蒸发作用而使其同位素分馏,同位素质量相对较轻的水汽会优先蒸发,剩余水体富集重同位素<sup>[30]</sup>,而锡林河干、支流下游区域的地表温度相对于上游普遍偏高,促使地表蒸发作用更强烈,因此越靠近下游地区氢氧同位素越富集.

相比干流,支流的河道狭窄且补给源单一,断流现象时常发生,流量的季节性明显,故支流δ<sup>18</sup>O从上游到下游波动变化不明显(图 5).但干流与支流交汇处 S6 和 S13 的稳定同位素富集明显,这可能是由于在交汇处地势趋于平坦且河道变宽,水流流速减缓,水面蒸发作用更加强烈,导致重同位素富集<sup>[31]</sup>.

时间上,干流氢氧同位素呈先下降(5~6月)后 上升(6~9月)再下降(9~10月)的趋势(图7),支 流氢氧同位素变化幅度较小. 说明锡林河流域河水 存在着明显的季节性差异,即春秋季贫化、夏季富 集的特征,此结果同房丽晶等<sup>[32]</sup>发现内蒙古高原巴

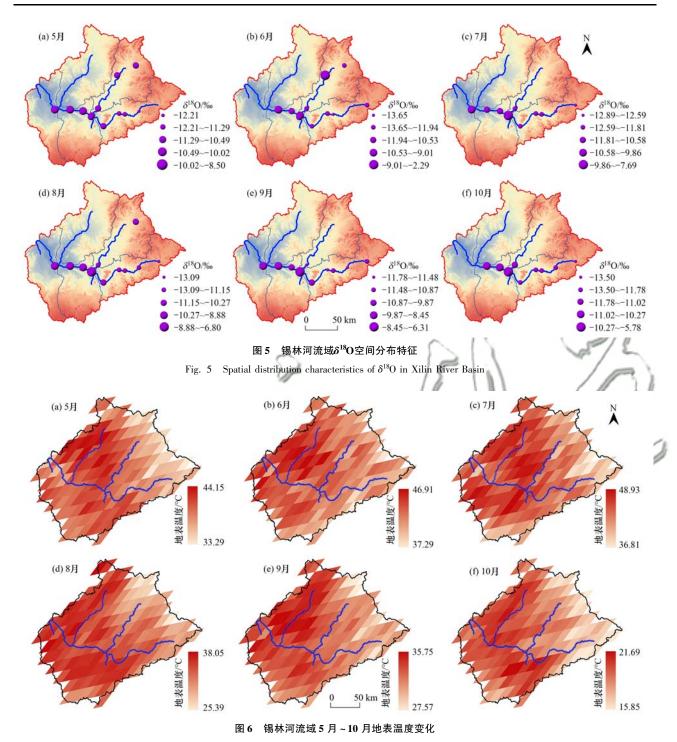


Fig. 6 Surface temperature change in Xilin River Basin from May to October

拉格尔河河水  $\delta D$  和 $\delta^{18}O$ 丰水期高而枯水期低的变化特征一致.

由图 7 可知, 5 月,流域河水 $\delta^{18}$ O的均值为 -9.89%,相比其他月份最低,说明 $\delta^{18}$ O在春末最为贫化,这可能是由于春末(5 月)气温回升,冰雪融水和冻土融水混合补给地表水造成的,这与杨永刚等 [33] 在研究马粪沟流域不同景观带水文过程中发现春末地表水 $\delta^{18}$ O均值趋于贫化的结果一致.6~8月,流域地表水 $\delta^{18}$ O均值仍较低,分别为 -9.17%。、-9.53%和 -9.63%,这是由于夏季(6~8 月)雨水

丰沛,大气降水因降水量效应而偏贫化,从而水接受了较贫化的大气降水补给.9月和10月,流域河水 $\delta^{18}$ O均值分别为 -8.59‰和 -8.53‰,此时已进入秋季,降水减少,气候干燥且风速较大,而空气湿度偏低,动力非平衡分馏作用显著,轻同位素优先蒸发,因此该时段 $\delta^{18}$ O均值有所升高,说明秋季河水的补给来源复杂,河水和地下水的转化关系发生转变.

#### 3.4 地下水稳定同位素的时空分布

根据不同类型地下水同位素值随时间的变化过程(图8),发现可按照浅层和深层地下水进行分类

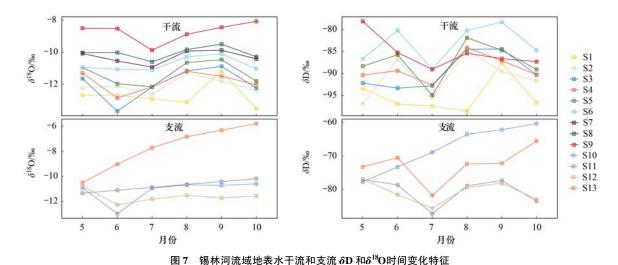


Fig. 7 Temporal variation characteristics of  $\delta D$  and  $\delta^{18}O$  in main stream and tributaries of Xilin River Basin

讨论.

由图 8 可知,相比河水,地下水同位素的波动变化并不剧烈; 浅层地下水较深层地下水 6 l 0 值在 5 ~7 月呈先下降后上升的趋势更加明显,其极小值出现在 6 月,这是由于研究区地处半干旱区,浅层地下水受融冰融雪径流补给影响要远大于深层地下水,而这些水源具有显著偏负的氢、氧同位素值[43],而深层地下水 6 l 8 O 低值则出现在 7 ~ 8 月,说明深层地下水受融冰融雪径流补给要明显滞后于浅

层地下水1~2个月;9月进入干旱季节后,浅层地下水δ<sup>18</sup>O值处于相对稳定阶段,深层地下水呈上升趋势,如前分析结果显示,同位素值呈地下水<河水<大气降水的规律,而融冰融雪径流与大气降水和地表水的时效性相比,特别是地表水人渗补给地下水具有明显的滞后效应<sup>[21]</sup>,深层地下水受大气降水和地表水的混合作用使其δ<sup>18</sup>O值呈上升趋势;同时,根据图9所示,3个自建地下水位监测井水位的动态变化可知,9月之后,地下水水位的抬升也证实了

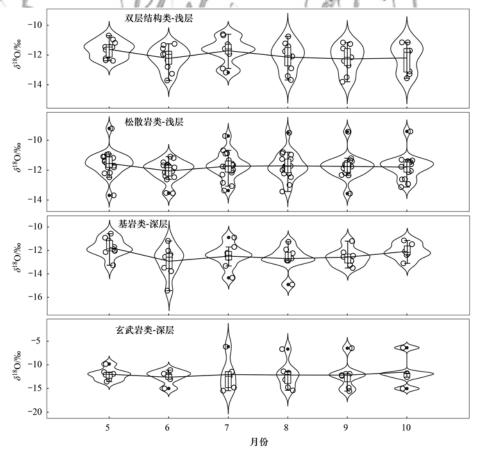


图 8 浅层地下水和深层地下水δ<sup>18</sup>O时间变化特征

Fig. 8 Temporal variation characteristics of  $\delta^{18}O$  in shallow groundwater and deep groundwater

大气降水和地表水的滞后补给,且水温的升高使得地下水 $\delta^{18}$ O值因温度效应影响而呈上升趋势.此外,地下水明显偏负的 $\delta^{18}$ O值可能与水体和周边介质发生同位素交换及生物(植物)作用有关[34].

由研究区浅层地下水和深层地下水δ<sup>18</sup>O的空间分布可知(图 10),二者表现出较为一致的空间分布特征,即δ<sup>18</sup>O值总体从东南向西北逐渐贫化,浅层地下水δ<sup>18</sup>O的这一趋势(变差系数为 - 0.073)比深层地下水的(变差系数为 - 0.057)更明显,这与浅层地下水埋深较浅,受温度、大气降水和地表水体混合渗漏补给影响有关,王雨山等<sup>[35]</sup>在雄安新区白洋淀的研究指出,地表水渗漏对浅层地下水垂向的影响深度为 20 m,以及推测了地下水位埋深较浅受蒸发影响导致同位素富集.

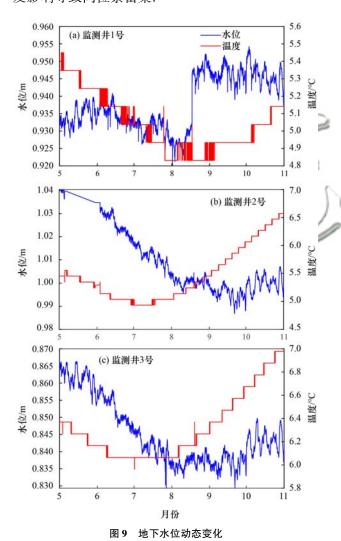


Fig. 9 Dynamic change in groundwater level

**3.5** 大气降水、河水和地下水氢氧同位素特征及相互转化关系

3.5.1 大气降水、河水与地下水 $\delta^{18}$ O和  $\delta D$  关系及 其指示意义

大气降水线可以较好地反映一个地区的自然地

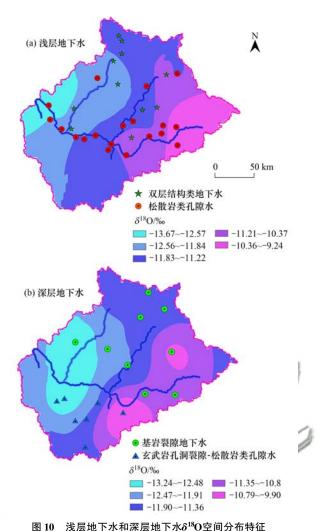


图 10 戊层地下水和床层地下水0 U至间分布特征

Fig. 10 Spatial distribution characteristics of  $\delta^{18}$ O in shallow groundwater and deep groundwater

理和气象条件<sup>[26]</sup>,Craig<sup>[36]</sup>基于大气降水中的 $\delta$ <sup>18</sup>O和  $\delta$ D 的关系,提出了全球大气降水线(GMWL: $\delta$ D =  $8\delta$ <sup>18</sup>O + 10),将其与其他水体 $\delta$ <sup>18</sup>O和  $\delta$ D 的组成进行比对,可解释区域地下水、河水的来源并阐明不同水体的相互转化关系<sup>[6]</sup>.

利用最小二乘法拟合研究区夏季(6~8月)当地大气降水线方程:

$$\delta D = 8.01 \, \delta^{18} O + 5.1$$
  
 $(n = 25, R^2 = 0.977)$  (5)

式中,斜率 8.01 表示δ<sup>18</sup>O 和 δD 的分馏速率,截距 5.1 表示 δD 对平衡状态的偏离程度<sup>[37]</sup>. 研究区当地大气降水线斜率接近于全球大气降水线斜率,主要原因是本次研究大气降水取样时间集中在雨季,降水量效应显著,从而掩盖了温度对δ<sup>18</sup>O 和 δD 的影响,分馏速率减缓,使得δ<sup>18</sup>O 和 δD 值贫化;当地大气降水线截距偏低,这与该流域属半干旱气候、蒸发量远大于降水量、在夏季主要受东南季风和局地蒸发的影响<sup>[28]</sup>、动力非平衡分馏效应起主导作用<sup>[25]</sup>

等有关. 相比 Wu 等<sup>[38]</sup>利用锡林河流域 2007 年 6 月至 2008 年 9 月降水数据得到降水线方程  $\delta D$  = 7. 89  $\delta^{18}O$  + 9. 5 (n = 11,  $R^2$  = 0. 97),本研究拟合的大气降水线斜率略微偏高且截距明显偏低,这主要是因为本研究所建方程以夏季的采样数据为基础,水文循环和水体交换的速率较年尺度更高.

由 5~10 月实测数据拟合逐月当地大气降水线、河水蒸发线、浅层地下水蒸发线和深层地下水蒸发线如图 11,对比发现地下水和河水蒸发线均位于当地大气降水线和全球大气降水线下方,表明地下水和河水受降水补给的同时,经历了不同程度的非平衡蒸发<sup>[32]</sup>;另一方面,与大气降水线相比,地下水蒸发线和河水蒸发线斜率较相近,表明流域地下水与河水的水力联系较强<sup>[39]</sup>;此外,各水体δ<sup>18</sup>O和 δD 关系线斜率在 5~10 月均表现为深层地下水>浅层地下水>河水,说明河水受蒸发影响最大,深层地下水基本不受蒸发影响.

观察各水体δ<sup>18</sup>O和 δD 关系随时间的变化(图 11),可以发现:①河水蒸发线的季节性差异较明显,其斜率和截距在生长季均表现为先减小后增大再减小的趋势,这主要是因为 5 月气温较 6 月偏低

且河水受融冰融雪径流补给,其蒸发分馏作用较弱, 6月随气温的升高,河水蒸发分馏作用增强,7~8月 进入雨季,受大气降水补给导致δ¹®O和 δD 贫化,9~ 10 月进入枯水季,河水受大气降水的补给减少,蒸 发分馏作用较强:② 浅层地下水蒸发线也表现出 季节性的变化特征,其斜率和截距均表现为先增大 后减小再增大的趋势,这是因为5~6月气温升高, 蒸发分馏作用增强,而7~8月浅层地下水受大气降 水入渗以及冻土融水的混合补给,其蒸发分馏减弱, 9~10 月进入枯水季,浅层地下水补给减少,蒸发分 馏作用增强;③ 深层地下水蒸发线的变化相比于 浅层地下水蒸发线存在明显的滞后,因为埋深较大 的地下水受大气降水和河水下渗补给的周期长,蒸 发分馏作用较弱,总体而言深层地下水蒸发线在6 ~10月相对稳定,在5月其斜率和截距较大,这是 对春冬季大气降水补给的滞后响应,前人基于氢氧 同位素的深层地下水的相关研究也有类似的发 现[30~42],7 月以后,深层地下水蒸发线均在浅层地 下水蒸发线下方,可能是因为夏秋季深层地下水会 受到浅层地下水的频繁补给; ④浅层地下水蒸发线 和深层地下水蒸发线在5月均位于全球水线和当地

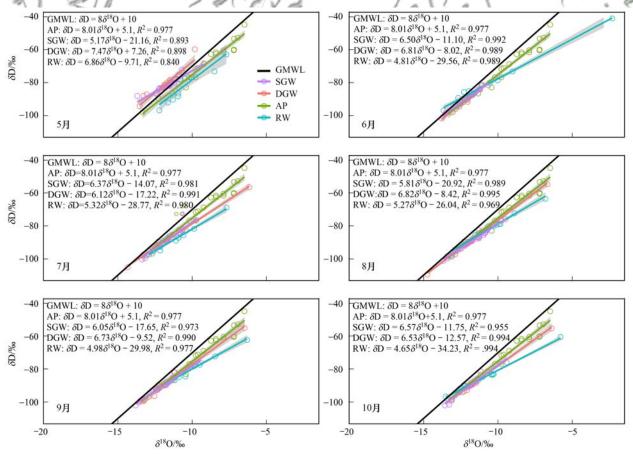


图 11 河水(RW)、浅层地下水(SGW)、深层地下水(DGW)和大气降水(AP)的δ<sup>18</sup>O、δD 关系 Fig. 11 Relationship between δ<sup>18</sup>O and δD of river water(RW), shallow groundwater(SGW),

deep groundwater(  $\ensuremath{\mathrm{DGW}})$  , and atmospheric precipitation(  $\ensuremath{\mathrm{AP}})$ 

大气水线上方,而在其余月份均位于全球大气降水线右下方(图 11),说明 5 月地下水一部分由春冬季较夏秋季更贫化的大气降水补给,另一部分伴随其他水源补给,且 5 月的浅层地下水蒸发线与深层地下水蒸发线的相交映证了浅层地下水和深层地下水的相互补给,且地下水补给源复杂;⑤ 各月河水蒸发线均位于浅层和深层地下水蒸发线下方,表明大气降水和地下水均对该流域的河水有补给作用<sup>[43]</sup>,这同杨淇越等<sup>[44]</sup>在锡林河流域相关研究中的发现一致,即在径流季节,除大气降水对地表水的主要贡献外,地下水也是地表水的主要补给来源.

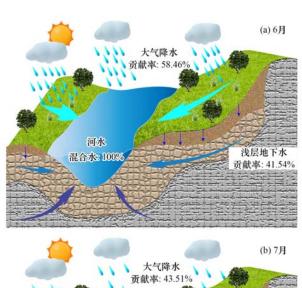
#### 3.5.2 大气降水-河水-地下水转化关系

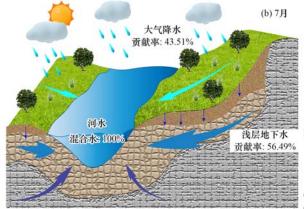
流域水文循环过程促使不同水体在不同时间进行着复杂的补给与交换<sup>[25]</sup>. 研究区属中温带半干旱大陆性季风气候地区,降水年内分配不均,6~8 月降水量占年总降水量的50%以上(图4)<sup>[19]</sup>. 由锡林河流域九曲湾水文站(2019年自建水文站)2020~2021年的流量变化可知,锡林河自6月开始流量迅速增大,7月和8月达到峰值,6~8月径流量占全年总径流量34.41%. 由此综合考虑流域水文气象条件,选取6、7和8月探讨不同水体间的转化规律,应用同位素质量平衡模型,估算浅层地下水与大气降水对河水的转化比例(表2).

6月,大气降水、浅层地下水、河水的 &D 和&<sup>18</sup>O 的均值分别为 - 76. 78‰和 - 9. 76‰、 - 89. 99‰和 - 12. 12‰、 - 81. 70‰和 - 10. 84‰,可见氢氧同位素值总体大小表现为:大气降水 > 河水 > 浅层地下水,其中大气降水和河水的氢氧同位素值较为接近,表明大气降水为河水的主要补给来源;基于二元同位素质量平衡模型估算得 6 月河水受大气降水和浅层地下水补给比例分别为 58. 46% 和 41. 54%,大气降水对河水的补给占主导[表 2 和图 12(a)].

7月,大气降水、浅层地下水、河水的  $\delta$ D 和 $\delta$ <sup>18</sup>O 均值分别为 - 75.91‰和 - 9.75‰、 - 88.63‰和 - 11.73‰、 - 84.88‰和 - 10.59‰,各类水体氢氧同位素变化规律依然是大气降水最富集,浅层地下水最贫化,但各类水体  $\delta$ D 和 $\delta$ <sup>18</sup>O比 6 月略有增大,可能与 7 月温度升高(图 3),水体受蒸发分馏作用增强有关; 7 月大气降水和浅层地下水对河水的补给比例分别为 43.51%和 56.49%,浅层地下水成为河水的主要补给源[表 2 和图 12(b)].

8月,大气降水、浅层地下水、河水 δD 和δ<sup>18</sup>O均值分别为 - 71.54‰和 - 9.80‰、 - 89.78‰和 - 11.87‰、 - 80.99‰和 - 10.44‰,各类水体氢氧同位素的变化规律同前期一致,而河水受大气降水和浅层地下水的补给比例分别为 56.11%和





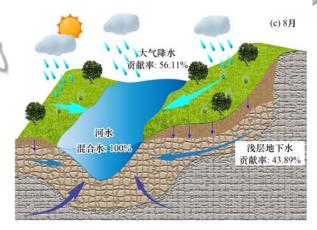


图 12 锡林河流域大气降水-河水-浅层地下水转化示意

Fig. 12 Atmospheric precipitation-river water-shallow groundwater conversion in the Xilin River Basin

43.89%,说明了大气降水在8月补给仍然较强[表2和图12(c)].

综上所述,锡林河流域河水在6~8月的主要补给来源包括大气降水和浅层地下水,补给比例分别为43.51%~58.46%和41.54%~56.49%,平均补给比例分别为52.69%和47.31%.经过实地勘探发现,锡林河河道内有多处地下水泉眼沿途溢流(如卧龙泉、锡林河源头),证实了深层地下水对河水也有较强的补给作用. 孙从建等[45] 在塔里木盆地西南部内陆河流域通过径流分割组分研究发现,冰雪融水、地下水及降水对于年径流的贡献率分别为

17%、40% 和 43%, 地下水与降水贡献率仅相差 3%. 尹立河等[46] 通过梳理西北内陆河流域地下水循环特征研究成果, 得出西北内陆河流域的地下水

与地表水频繁转化,含水层-河流系统具有密切的水力联系.说明对于内陆河流域,地下水和地表水之间的补排关系更密切且转化机制更复杂.

#### 表 2 河水、浅层地下水及大气降水混合比值计算

| Table 2 | Calculation | of mixing ratio | of river water | , shallow groundwater | and atmospheric | precipitation |
|---------|-------------|-----------------|----------------|-----------------------|-----------------|---------------|
|         |             |                 |                |                       |                 |               |

| 月份 | 计算项    | 计算项构成 | $\delta \mathrm{D}/\%$ o | $\delta^{18} \mathrm{O}/\%$ | 按δ <sup>18</sup> O混合比值<br>/% | 按 δD 混合比值<br>/% | 均值混合比值<br>/‰ |
|----|--------|-------|--------------------------|-----------------------------|------------------------------|-----------------|--------------|
|    | 端元     | 大气降水  | -76.78                   | -9.76                       | 54. 20                       | 62. 72          | 58. 46       |
| 6  | 케이기다   | 浅层地下水 | - 89. 99                 | -12.12                      | 45. 80                       | 37. 28          | 41. 54       |
|    | 混合水    | 河水    | -81.70                   | - 10. 84                    | 100.00                       | 100.00          | 100.00       |
|    | 端元     | 大气降水  | -75. 91                  | -9.75                       | 57. 55                       | 29. 46          | 43. 51       |
| 7  | 케이기다   | 浅层地下水 | -88.63                   | -11.73                      | 42. 45                       | 70. 54          | 56. 49       |
|    | 混合水    | 河水    | -84.88                   | - 10. 59                    | 100.00                       | 100.00          | 100.00       |
|    | 端元     | 大气降水  | -71.54                   | -9.80                       | 66. 55                       | 45. 68          | 56. 11       |
| 8  | 711170 | 浅层地下水 | -89.78                   | -11.87                      | 33. 45                       | 54. 32          | 43. 89       |
|    | 混合水    | 河水    | - 80. 99                 | - 10. 44                    | 100.00                       | 100.00          | 100.00       |

#### 4 结论

- (1)锡林河流域具有明显的内陆半干旱气候特征,水汽来源复杂;地下水和河水主要受大气降水补给,并经历了不同程度的非平衡蒸发;河水受蒸发影响最大,深层地下水基本不受蒸发影响.
- (2)河水 δD 和δ¹δO组成,表现出春秋贫化、夏季富集的季节性特征,说明锡林河流域在春季和秋季分别以融雪径流和降雨径流为主;空间上总体呈现出从上游到下游逐渐升高的趋势,干支流交汇处稳定同位素富集明显.
- (3)浅层地下水较深层地下水δ<sup>18</sup>O在 5~7 月表现出先降后升趋势,但在生长季末期(9~10 月)深层地下水δ<sup>18</sup>O明显增大,说明大气降水和河水对深层地下水入渗补给的滞后响应;地下水同位素在空间上总体呈现出由东南向西北逐渐贫化,且浅层地下水的这一空间变化更明显.
- (4)夏季河水的大气降水和浅层地下水补给比例分别为52.69%和47.31%,同时深层地下水也具有重要的补给作用,证实了内陆河流域河水和地下水之间密切的转化关系.

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| Water Environment Characteristics and Water Quality Assessment of Typical Lakes in Inner Mongolia  Relationship Between Precipitation, River Water, and Groundwater Conversion in the Upper Reaches of Xilin River During the Rai  Hydrochemical Characteristics and Control Factors of Groundwater in the Northwest Salt Lake Basin  Groundwater Pollution Risk Assessment in Plain Area of Barkol-Yiwu Basin  Bisulfite Promoted Minute Fe <sup>2+</sup> -Activated Peroxydisulfate for Paracetamol Degradation  Degradation of Ciprofloxacin by Activating Peroxymonosulfate with Sludge Biochar  Adsorption of Iopamidol by NaHCO <sub>3</sub> -activated Buckwheat Biochar  Preparation of Bamboo-based N, P Co-doped Activated Carbon and Its Lanthanum Ion Adsorption Performance  Analysis of Vegetation Change and Influencing Factors in Southwest Alpine Canyon Area  Effect of Biochar on Agricultural Soil Aggregates and Organic Carbon; A Meta-analysis  Carbon Cycling Processes in Croplands and Their Quantification Methods  Effects of Biochar Application on Soil Organic Carbon Component in Eucalyptus Plantations After Five Years in Northern Guangxi Evolution Characteristics of Soil Active Organic Carbon and Carbon Pool Management Index Under Vegetation Restoration in Karst  Pollution Characteristics and Ecological Risk Assessment of Typical Antibiotics in Environmental Media in China   | iny Season   | (3754)<br>(3767)<br>(3790)<br>(3801)<br>(3811)<br>(3823)<br>(3833)<br>(3833)<br>(3857)<br>(3857)<br>(3869)   |
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| Water Environment Characteristics and Water Quality Assessment of Typical Lakes in Inner Mongolia  Relationship Between Precipitation, River Water, and Groundwater Conversion in the Upper Reaches of Xilin River During the Rai  Hydrochemical Characteristics and Control Factors of Groundwater in the Northwest Salt Lake Basin  Groundwater Pollution Risk Assessment in Plain Area of Barkol-Yiwu Basin  Bisulfite Promoted Minute Fe <sup>2+</sup> -Activated Peroxydisulfate for Paracetamol Degradation  Degradation of Ciprofloxacin by Activating Peroxymonosulfate with Sludge Biochar  Adsorption of Iopamidol by NaHCO <sub>3</sub> -activated Buckwheat Biochar  Preparation of Bamboo-based N, P Co-doped Activated Carbon and Its Lanthanum Ion Adsorption Performance  Analysis of Vegetation Change and Influencing Factors in Southwest Alpine Canyon Area  Effect of Biochar on Agricultural Soil Aggregates and Organic Carbon; A Meta-analysis  Carbon Cycling Processes in Croplands and Their Quantification Methods  Effects of Biochar Application on Soil Organic Carbon Component in Eucalyptus Plantations After Five Years in Northern Guangxi  Evolution Characteristics of Soil Active Organic Carbon and Carbon Pool Management Index Under Vegetation Restoration in Karst  Pollution Characteristics and Ecological Risk Assessment of Typical Antibiotics in Environmental Media in China  Spatial Prediction Modeling for Soil pH Based on Multiscale Geographical Weighted Regression (MGWR) and Its Influencing Facto  Characteristics and Source Analysis of Heavy Metal Pollution in Farmland Around a Coal-fired Power Plant   | iny Season  SUN Jin, WANG Yi-xuan, YANG Lu, et al. (6)  YAN Yan, GAO Rui-zhong, LIU Ting-xi, et al. (6)  LIU Yu, ZENG Yan-yan, ZHOU Jin-long, et al. (6)  MO Xi-ting, NIE Shu-hua, YAN Cai-xia, et al. (6)  ZHENG Da-yang, ZOU Jia-li, XU Hao, et al. (6)  WANG Gui-long, LIU Yan-yan, JING Li-ming, et al. (6)  WANG Gui-long, LIU Yan-yan, JIANG Rong-yuan, et al. (6)  MENG Yan, SHEN Ya-wen, MENG Wei-wei, et al. (6)  MENG Yan, SHEN Yu-wi, CAO Yang, et al. (6)  MOU Zhi-yi, SHEN Yu-yi, CAO Yang, et al. (6)  Area  CAI Hua, SHU Ying-ge, WANG Chang-min, et al. (6)  CHEN Li-hong, CAO Ying, LI Qiang, et al. (6)  SHAO Ming-song, CHEN Xuan-qiang, XU Shao-jie, et al. (6)  ZHANG Jun, LI Xu, LIU Lei-yu, et al. (6)  ZHANG Yu-rong, LUO Shuai, CHEN Yuan, et al. (6)   | (1754)<br>(1767)<br>(1778)<br>(1790)<br>(1801)<br>(1823)<br>(1823)<br>(1823)<br>(1823)<br>(1823)<br>(1823)<br>(1823)<br>(1823)<br>(1847)<br>(1857)<br>(1869)<br>(1894)<br>(1992)<br>(1992)<br>(1993)                               |
| Water Environment Characteristics and Water Quality Assessment of Typical Lakes in Inner Mongolia  Relationship Between Precipitation, River Water, and Groundwater Conversion in the Upper Reaches of Xilin River During the Rai  Hydrochemical Characteristics and Control Factors of Groundwater in the Northwest Salt Lake Basin  Groundwater Pollution Risk Assessment in Plain Area of Barkol-Yiwu Basin  Bisulfite Promoted Minute Fe <sup>2+</sup> -Activated Peroxydisulfate for Paracetamol Degradation  Degradation of Ciprofloxacin by Activating Peroxymonosulfate with Sludge Biochar  Adsorption of Iopamidol by NaHCO <sub>3</sub> -activated Buckwheat Biochar  Preparation of Bamboo-based N, P Co-doped Activated Carbon and Its Lanthanum Ion Adsorption Performance  Analysis of Vegetation Change and Influencing Factors in Southwest Alpine Canyon Area  Effect of Biochar on Agricultural Soil Aggregates and Organic Carbon; A Meta-analysis  Carbon Cycling Processes in Croplands and Their Quantification Methods  Effects of Biochar Application on Soil Organic Carbon Component in Eucalyptus Plantations After Five Years in Northern Guangxi  Evolution Characteristics of Soil Active Organic Carbon and Carbon Pool Management Index Under Vegetation Restoration in Karst  Pollution Characteristics and Ecological Risk Assessment of Typical Antibiotics in Environmental Media in China  Spatial Prediction Modeling for Soil pH Based on Multiscale Geographical Weighted Regression (MGWR) and Its Influencing Facto  Characteristics and Source Analysis of Heavy Metal Pollution in Farmland Around a Coal-fired Power Plant  Characteristics of Soil Pollution and Source Analysis of Typical Pollutants in the Petrochemical Site  | iny Season  SUN Jin, WANG Yi-xuan, YANG Lu, et al. (6)  YAN Yan, GAO Rui-zhong, LIU Ting-xi, et al. (6)  MO Xi-ting, NIE Shu-hua, YAN Cai-xia, et al. (6)  ZHENG Da-yang, ZOU Jia-li, XU Hao, et al. (6)  WANG Gui-long, LIU Yan-yan, JING Li-ming, et al. (6)  WANG Gui-long, LIU Yan-yan, JIANG Rong-yuan, et al. (6)  WANG Gui-long, LIU Yan-yan, JIANG Rong-yuan, et al. (6)  MENG Yan, SHEN Ya-wen, MENG Wei-wei, et al. (6)  MENG Yan, SHEN Yu-yi, CAO Yang, et al. (6)  MOU Zhi-yi, SHEN Yu-yi, CAO Yang, et al. (6)  Area  CAI Hua, SHU Ying-ge, WANG Chang-min, et al. (6)  CHEN Li-hong, CAO Ying, LI Qiang, et al. (6)  TS  ZHAO Ming-song, CHEN Xuan-qiang, XU Shao-jie, et al. (6)  ZHANG Yu-rong, LIO Shuai, CHEN Yuan, et al. (6)  Areas  CHEN Rui, CHENG Jian-hua, TANG Xiang-yu (6)   | (754)<br>(778)<br>(778)<br>(790)<br>(881)<br>(8823)<br>(8833)<br>(8847)<br>(8869)<br>(8869)<br>(894)<br>(9999)<br>(9993)<br>(9947)   |
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