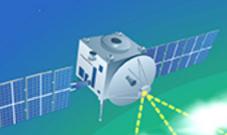


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ENVIRONMENTAL SCIENCE

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PM_{2.5}和O₃污染协同防控区的遥感精细划定与分析 李沈鑫,邹滨,张凤英,刘宁,薛琛昊,刘婧



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- 主办 中国科学院生态环境研究中心
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汾河流域地表水水化学同位素特征及其影响因素

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摘要: 汾河是山西的母亲河. 由于水资源过度开发及社会经济发展影响,生态环境恶化. 经过一系列治理保护措施,水质得到改善. 利用统计学、Piper 三线图和 Gibbs 模型等方法,分析了汾河流域地表水水化学和氢氧同位素特征及其来源,揭示了汾河流域地表水水质演化过程. 结果表明,汾河干流地表水中主要水化学组分沿着径流路径含量逐渐增加; 汾河上游地表水水化学类型主要以 $\mathrm{HCO}_3 \cdot \mathrm{SO}_4 \cdot \mathrm{Cl-Ca} \cdot \mathrm{Na} \cdot \mathrm{Mg}$ 为主,中游和下游地表水水化学类型主要以 $\mathrm{SO}_4 \cdot \mathrm{HCO}_3 \cdot \mathrm{Cl-Ca} \cdot \mathrm{Na} \cdot \mathrm{Mg}$ 为主. 汾河流域地表水水化学组成主要受岩石风化作用和蒸发结晶作用影响,而降雨影响较小. Na^+ 和 K^+ 主要来源于蒸发盐岩的溶解以及周边黄土中含 Na 矿物溶解,水体中 Ca^{2+} 、 Mg^{2+} 和 HCO_3 主要来源于碳酸盐岩的溶解, SO_4^2 除来源于石膏的溶解,还可能来源于汾河周边黄土层中硫化矿物的溶解. 干流地表水 $\mathrm{\delta D}$ 和 $\mathrm{\delta^{18}O}$ 平均值分别为 $\mathrm{-62.60}$ %和 $\mathrm{-8.42}$ %。氢氧同位素特征进一步表明其主要受蒸发作用的影响. 流域内支流及岩溶水水化学组分差异较大. 结果可为汾河流域生态修复保护及生态文明建设提供依据.

关键词:汾河流域;水化学;离子来源;氢氧同位素;地表水

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Hydrochemical and Isotopic Characteristics in the Surface Water of the Fenhe River Basin and Influence Factors

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Abstract: The Fenhe River Basin is the mother river of Shanxi Province. Due to the over-exploitation of water resources and the impact of social and economic development, the ecological environment has deteriorated. After a series of treatment and protection measures, the water quality has since been improved. Mathematical statistics, Piper diagrams, Gibbs model, hydrogen and oxygen isotopes, and other methods were used to analyze the characteristics and sources of hydrochemistry in the surface water of the Fenhe River basin, which revealed the evolution process of surface water quality of the Fenhe River basin. The results showed that the content of the main hydrochemical components in the main stream surface water of Fenhe River basin increased gradually along the runoff path. The hydrochemical types of surface water of Fenhe River basin were mainly HCO₃·SO₄·Cl-Ca·Na·Mg and SO₄·HCO₃·Cl-Ca·Na·Mg. There were great differences in hydrochemical components of tributaries and karst water in the basin. The hydrochemical types of surface water of karst water were mainly SO₄·HCO₃-Ca·Mg. The hydrochemical composition of surface water in Fenhe River basin was mainly affected by rock weathering and evaporation crystallization, whereas rainfall had little effect. Na + and K + mainly came from the dissolution of evaporated salt rocks with Na in the surrounding loess. Ca²⁺, Mg²⁺, and HCO₃-mainly came from the dissolution of carbonate rocks. SO₄²⁻ may have also come from the dissolution of sulfide minerals in the loess layer around Fenhe River in addition to the dissolution of gypsum. The values of δD and δ¹⁸O of Fenhe River surface water were gradually enriched from upstream to downstream. The characteristics of hydrogen and oxygen isotopes further showed that the surface water was mainly affected by evaporation. The results of this study can provide evidence for ecological restoration and protection and ecological civilization construction in the Fenhe River basin.

Key words: Fenhe River Basin; hydrochemistry; ion sources; hydrogen and oxygen isotopes; surface water

水体在径流、渗流过程中与其周围环境发生着一系列的水岩作用,在水化学特征上表现为水化学指标持续发生着一系列变化^[1,2].通过水体中的主要离子组分可以划分水化学类型、分析其离子来源、演化机制和影响因素等^[3~6].水化学结合同位素技术可以有效识别不同水体的补给来源及水质演化过程^[7~11].

汾河属黄河一级支流,为山西省最大的河流, 也是山西省人民的母亲河,为当地经济和社会发 展提供了不可替代的支撑性水源保障. 然而随着 经济和社会的快速发展,多年来流域水资源开发 利用程度长期居高不下,造成流域水环境日趋恶化等一系列生态环境问题^[12~20],一定程度上制约了山西省经济社会的生态健康发展.近年来,国家和山西省政府出台了一系列关于汾河流域生态修复的政策措施和工程^[21~23],汾河地表水水质正逐步改善.为系统分析流域水污染特征和水化学演

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变机制,为汾河流域生态综合整治提供技术支撑,本研究选取汾河流域典型断面,采集水化学和同位素样品,利用离子分析及同位素方法,探讨汾河水化学及氢氧同位素特征,分析其主要影响因素和驱动力,以期为汾河流域生态环境综合治理及汾河水环境评价提供数据支撑.

1 材料与方法

1.1 研究区概况

汾河是黄河流域第二大支流,发源于山西省宁 武县管涔山,在万荣县庙前村汇入黄河,全长716 km,流经山西6个地市、34个县级城市,流域总面 积39 721 km²,约占山西省面积的四分之一,多年平 均径流量 21 × 108 m3[23]. 沿途汇入众多支流,其中 流域面积大于1 000 km² 的有 9 条,大于 500 km² 的 支流有 16 条;支流中流域面积最大的支流为潇河, 其次为文峪河、昌源河和浍河,流域面积均大于 2000 km². 流域内自上游到下游发育有雷鸣寺、晋 祠、兰村、洪山、郭庄、霍泉、龙子祠和古堆这8个 岩溶大泉,构成汾河的重要补给源[23,24]. 汾河流域 西靠吕梁山、东临太行山,地势北高南低,由北向南 纵贯省境中南部,各大支流水系发源于两大山系之 间,地形地貌总体上表现为南北长、东西狭,呈不规 则宽带状分布在省境中部地区. 地处中纬度大陆性 季风带,多年平均降水量 472 mm^[25]. 流域内地层发 育较为齐全,区域岩溶含水岩层主要为寒武-奥陶系 碳酸盐岩地层.

1.2 样品采集与分析方法

为系统分析汾河流域地表水水化学及同位素特征,沿汾河干流、各主要支流进入干流前、流域典型岩溶大泉和岩溶水水井采集样品.于2020年5月枯水期采集47组样品(图1),其中干流样品31组,支流样品8组,岩溶泉4组,岩溶井样品4组(其中S2井深5m,S3井深780m,S4井深100m,S8井深240m).因为各支流流域面积大和情况复杂,本次评价难以全面覆盖,仅以入汾口代表子流域的水化学和地质环境特征.

样品采集按照《水质 采样技术指导》(HJ 494-2009)进行取样. 水化学组分由 PHS-3C pH 计、可见光光度计和离子色谱仪等测定. 氢氧同位素利用稳定同位素质谱仪 MAT253 测定, 氢同位素采用 Pt水平衡法测定, 氧同位素采用 CO₂-H₂O水平衡法测定, 水化学组分和氢氧同位素均由国土资源部岩溶地质资源环境监督检测中心测试. 硫同位素利用气体同位素质谱仪测定,由国土资源部中南矿产资源监督检测中心测试.

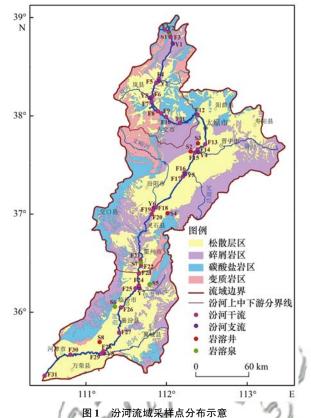


Fig. 1 Sketch map of the Fenhe River Basin and the sampling sites

2 结果与讨论

2.1 水化学特征

统计分析汾河干流地表水主要水化学组分(表1和图2),表明干流地表水整体上呈弱碱性,pH范围为7.41~8.10,平均值7.76,上游pH明显高于中下游pH[图2(a)],且由于两侧支流汇入混合后,pH降低;支流pH范围为7.49~7.97,平均值为7.72.从上游到下游,干流地表水中主要水化学组分均呈现从上游到下游逐渐增加的趋势[图2(a)~2(d)],说明沿着径流路径溶解矿物增多、河道蒸发浓缩作用加强. 汾河干流溶解性总固体(TDS)范围为214.00~966.00 mg·L⁻¹,平均值为590.06 mg·L⁻¹,从上游到下游TDS逐渐增加[图2(c)],说明地表水在径流过程中TDS升高主要是硫酸盐和碳酸盐岩矿物的溶解,盐岩的溶解量较小^[7]. 汾河干流地表水氨氮均达到地表水 V 类水标准,这与同期汾河国考断面分析结果相一致^[26,27].

汾河流域各主要支流 pH 大都呈弱碱性, pH 范围为 7.49 ~ 7.97(表 1). 各支流变异系数大小为: $SO_4^{2^-} > K^+ > Na^+ > Cl^- > Mg^{2^+} > TDS > Ca^{2^+} > HCO_3^- > pH, 变异系数较大, 说明各支流的水化学成分差异较大, 反映了不同的支流其水文地球化学背景及社会经济发展和工农业布局等差异较大. 岩溶$

大泉排泄水是汾河流域重要的补给水源,其水化学特征与流域总体水化学特征基本一致(表1).

表 1 汾河流域主要水化学组分统计结果1)

Table 1	Statistics of	major ione	in Fanha	River basin
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水体类型	指标	рН	Ca ^{2 +}	Mg ^{2 +}	K +	Na +	HCO ₃	Cl -	SO ₄ -	TDS
	最大值	8. 10	161.03	64. 93	13. 53	104. 10	250. 33	146. 69	548. 27	966. 00
	最小值	7.41	53. 11	12. 46	1. 34	4. 13	161. 14	5. 63	35. 86	214. 00
汾河干流	平均值	7. 76	84. 56	29. 96	6. 93	75. 96	219.81	90.02	179.63	590.06
	标准差	0.19	24. 07	9. 30	0.77	28. 52	22. 55	38. 87	90. 95	180. 48
	变异系数	0.03	0. 28	0.31	0.57	0.38	0.10	0.43	0.51	0.31
	最大值	7. 97	110. 42	67. 90	17. 41	185. 28	379. 15	161.90	421. 50	1 166. 00
	最小值	7.49	35. 74	15. 19	1.45	12. 65	145. 34	9. 58	33. 79	242. 00
汾河支流	平均值	7. 72	67. 35	28. 23	8. 61	86. 56	242. 24	82. 45	153. 27	559. 75
	标准差	0.18	24. 57	17. 91	6. 82	67. 78	72. 55	60.71	128. 58	317. 65
	变异系数	0.02	0.36	0. 63	0.79	0. 78	0.30	0.74	0.84	0. 57
	最大值	8. 06	294. 42	72. 76	5. 02	86. 82	261. 19	85. 75	822. 37	1 364. 00
	最小值	7. 02	63. 24	14. 84	1.04	3. 76	208. 53	3. 26	56. 80	274. 00
岩溶水	平均值	7. 74	135. 39	37. 58	2. 36	28. 01	235. 52	27. 24	304. 73	658. 25
	标准差	0.32	74. 61	17. 53	1.30	27. 77	17. 24	29.67	242. 86	336. 62
	变异系数	0.04	0.55	0. 47	0.55	0. 99	0.07	1.09	0.80	0. 51

1)pH 无量纲,其余组分单位为mg·L⁻¹

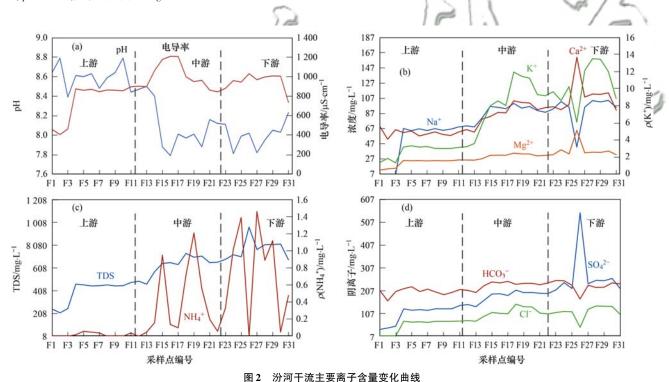


Fig. 2 Variation in main ions contents in main stream of Fenhe River

2.2 水化学类型

Piper 三线图是研究地表水的水化学类型特征的重要工具^[3,28,29]. 从汾河干流 Piper 图可以看出[图3(a)和3(b)],汾河上游碳酸盐岩裸露山区地表水水化学类型主要以 HCO₃-Ca·Mg 和 HCO₃·SO₄-Ca·Mg 为主(F1、F2 和 F3 点),且同期无大量有效降雨和地表产流,反映了枯水期汾河水主要来源于浅层地下水基流渗出补给,这与上游地质情况一致.由于引黄水、东碾河及岚河水的汇入,汾河水中 Cl⁻和 Na⁺比例逐渐增多,水化学类型向 HCO₃·SO₄·Cl-

Ca·Na·Mg 型演化,这是由于汾河周边地层从碳酸盐岩地层过渡到松散岩地层,而且支流中携带了大量的泥沙造成的. 汾河中游水化学类型总体上由 HCO₃·SO₄·Cl-Ca·Na·Mg 型逐步过渡为 SO₄·HCO₃·Cl-Ca·Na·Mg 型;中间由于潇河的汇入,水中 Na 离子比例增加,水化学类型由 HCO₃·SO₄·Cl-Ca·Na·Mg 型逐步过渡为 HCO₃·SO₄·Cl-Na·Ca·Mg 型.由于昌源河及文峪河的汇入,水中 HCO₃。离子比例增加,水化学类型发生变化,由于晋祠泉域、郭庄泉域和洪山泉域岩溶水对地表水的补给以及溶解的黄土层矿物的影响,导

致水中 SO_4^2 比例增加. 汾河下游水化学类型主要以 $SO_4 \cdot HCO_3 \cdot Cl$ - $Ca \cdot Na \cdot Mg$ 型为主. 由于浍河水的汇 人,导致水中 Cl 比例增加,混合浍河水后的水化学 类型为 $SO_4 \cdot Cl \cdot HCO_3$ - $Ca \cdot Na \cdot Mg$ 型.

总体上,汾河水在径流过程中,水中阴离子从以HCO₃⁻ 为主,逐渐过渡为以SO₄² 和HCO₃⁻ 为主,逐渐过渡为以SO₄² 和HCO₃⁻ 为主;优势阳离子从Ca²⁺和Mg²⁺为主逐步过渡到以Ca²⁺、Na⁺和Mg²⁺为主,Na⁺高于Mg²⁺,反映了从上游到下游,汾河流域地形地貌由山区逐渐过渡到盆地、区域地层由碳酸盐岩地层逐步过渡到松散层等;地表水在径流过程中一方面由于其他水流混入发生混合作用,另一方面蒸发作用加强,发生一系列水岩作用导致水中Na⁺含量增多,钙镁矿物沉淀^[30],以及

与流域内黄土层矿物渗入、工农业用水排放有关. 需要说明的是 F26 点周边有垃圾堆放及城市污水排放,导致地表水化学类型为 SO₄-Ca·Mg,其水化学特征显著不同于其他点,且经过一段时间的扩散混合后,至下游断面,地下水类型转变为 SO₄·HCO₃·Cl-Ca·Na.

由汾河流域各支流的水化学类型 Piper 图可知 [图 3(c)],上游支流阴离子以 HCO₃ 为主,阳离子以 Ca²⁺为主,反映了上游其主要来源于浅层地下水基流渗出补给;中下游的各支流由于平原区占比较大,Na⁺、Cl⁻和 SO₄ 比例逐渐加大,逐渐转为主要离子成分,反映了农田灌溉用水、浅层黄土层渗滤水的混入、蒸发浓缩作用等逐渐加强以及岩溶水对

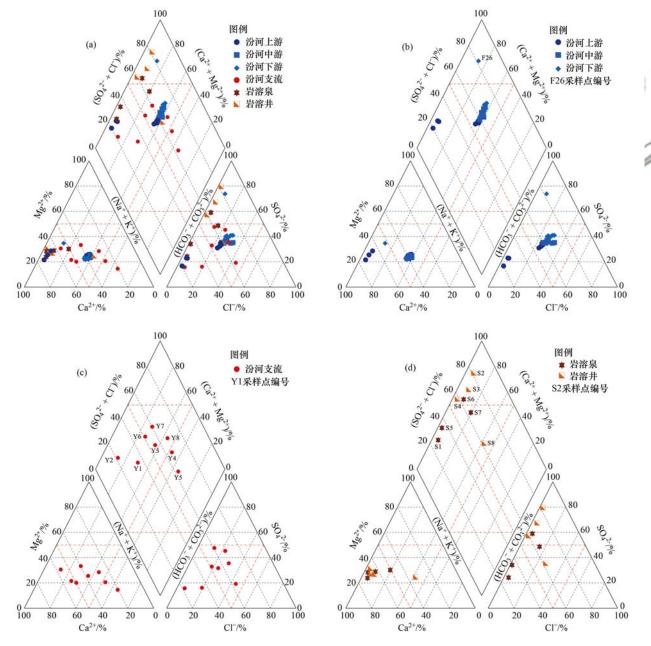


图 3 汾河流域主要离子 Piper 三线图

Fig. 3 Piper diagram showing major ion composition of the Fenhe River basin

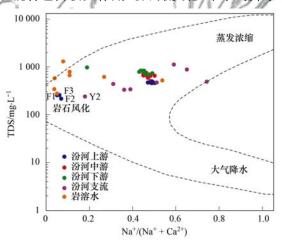
地表水的补给等.

同期采集的 4 组典型岩溶大泉和 4 个典型岩溶 井水样品[图 3(d)],除雷鸣寺泉地下水化学类型 为 HCO₃-Ca 型外,其他 7 组样品中 SO₄²⁻ 含量均占 较大比例,平泉(S2)水化学类型为 SO₄-Ca·Mg 型, 其余除霍泉(S5)及古堆泉(S8)附近样点外,其他 6 组水化学类型均为 SO₄·HCO₃-Ca·Mg 型,其所在泉 域均为特殊的"水煤共存系统",岩溶水受到不同程 度的污染,水中 SO₄²⁻ 比例较大^[31,32].

2.3 离子来源及影响因素

2.3.1 Gibbs 分析

水的水文地球化学过程研究常采用 Gibbs 图来进行^[5,33~35].由图 4 可知,研究区各水样点除汾河上游 3 个样点及下游 1 个样点外,其余样点主要位于Gibbs 图中上部分,处于岩石风化和蒸发结晶带中间,而且都远离大气降水带,说明本次采集的水样主要来源于地下水和工农业弃水,大气降水所占比例不大,浅层地下水和岩溶大泉排泄的基流构成了枯水期汾河径流的主要部分.在沿程径流过程中,蒸发浓缩和混合也占较大作用. 汾河流域多年平均降水



量为 472. 4 mm,水面蒸发量1 000~2 000 mm^[36],年内 5 月为枯水期末,降雨少,而 5 月气温快速回升、风速较大,蒸发旺盛,地表水受降雨影响小而受蒸发影响大^[23],这与 Gibbs 图中各样点分布区相一致.

汾河上游岔上水文站(F1)和雷鸣寺泉汾河上 下游断面(F2 和 F3) 这 3 个样点的 TDS 均值为 238 mg·L⁻¹,Cl⁻/(Cl⁻ + HCO₃)(当量浓度比,下同)在 0.04~0.06之间, Na+/(Na++Ca²⁺)在0.05~0.07 之间,东碾河(Y2)点其水源主要来源于山区地下水, 地下水溶滤作用较强,其离子来源主要受岩石风化影 响. 汾河地表水水样点除以上 3 个点外,其余点 TDS 为 450~818 mg·L⁻¹, Cl⁻/(Cl⁻ + HCO₃⁻)为 0.34~ 0.54, Na⁺/(Na⁺ + Ca²⁺) 为 0.43 ~ 0.5, 在 Gibbs 图中 位于岩石风化区域蒸发结晶区之前,说明汾河地表水 主要来源于以溶滤作用为主的地下水,在地面径流过 程中又受蒸发浓缩作用影响,且愈往下游,蒸发浓缩 过程愈强. 而各支流点主要受蒸发浓缩作用和岩石风 化作用影响. 岩溶水则多以岩石风化作用为主,这与 岩溶水排泄区接受上游地下水的补给,且埋深较大, 蒸发浓缩作用不明显.

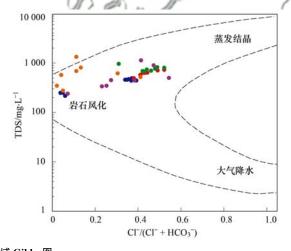


图 4 汾河流域 Gibbs 图
Fig. 4 Gibbs plots of the Fenhe River Basin

2.3.2 离子来源分析

枯水期汾河流量主要来源于地下水,利用 Ca²⁺/Na⁺与 Mg²⁺/Na⁺的关系可以判断水中各离子是来源于蒸发盐岩、碳酸盐岩或硅酸盐岩矿物的溶解^[37,38]. 汾河上游 3 个点 F1、F2、F3 及支流 Y2 以及岩溶水点主要受碳酸盐岩矿物溶解的影响,其所处地区地层岩性以碳酸盐岩为主. F26 点受人为污染外,其他汾河流域地表水的 Ca²⁺/Na⁺为 0. 34 ~ 2. 26, Mg²⁺/Na⁺为 0. 25 ~ 1. 32, 汾河流域地表水样点主要分布于蒸发盐岩和碳酸盐岩区中间(图 5), 说明汾河地表水主要受蒸发盐岩和碳酸盐岩等矿物溶解的影响, 这与其地质背景中的碳酸盐岩、膏岩分布一致^[39].

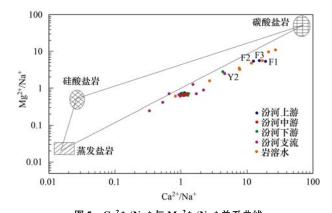


图 5 Ca²⁺/Na⁺与 Mg²⁺/Na⁺关系曲线

Fig. 5 Relationship of the Ca2+/Na+ and Mg2+/Na+

天然水体中 Na⁺和 K⁺主要来源于蒸发盐岩和硅酸盐岩矿物^[40]. (Na⁺ + K⁺)/Cl⁻为 1 表示水体中 Na⁺与 K⁺主要来源于蒸发盐岩的溶解^[2,41]. 汾河流域地表水大部分样点都在(Na⁺ + K⁺)/Cl⁻的比值1:1线上(图 6),说明汾河流域地表水 Na⁺和 K⁺除受蒸发盐岩溶解影响外,还受周边黄土中的含 Na矿物溶解的影响^[5,42]. 而汾河上游点(F1、F2、F3 和Y2)Na⁺和 K⁺主要来源于蒸发盐岩的溶解(图 6). 而潇河(Y4)、昌源河(Y5)和浍河(Y8)水 Na⁺、K⁺和 Cl⁻含量明显高于其他地表水点,主要是由于周边工矿企业及农业污染造成的,且此区段水土流失较为严重,受黄土中的含 Na 矿物溶解影响也较大.

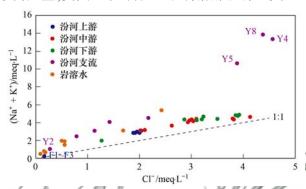


图 **6 (Na⁺ + K⁺)**与 Cl⁻ 关系曲线 Fig. **6** Relationship of (Na⁺ + K⁺) and Cl⁻

利用($Ca^{2+} + Mg^{2+}$)/HCO₃⁻ 与($Ca^{2+} + Mg^{2+}$)/(HCO₃⁻ +SO₄²⁻)可以反映水体中碳酸盐岩和蒸发盐岩(石膏)等的溶解特征^[7,43]. 若($Ca^{2+} + Mg^{2+}$)/HCO₃⁻ 为1,则说明水体中 Ca^{2+} 和 Mg^{2+} 主要来源于碳酸盐岩的溶解;若($Ca^{2+} + Mg^{2+}$)/(HCO₃⁻ +SO₄²⁻)为1,则说明水体中 Ca^{2+} 和 Mg^{2+} 还来源于蒸发盐岩(石膏)的溶解, SO_4^{2-} 还来源于石膏的溶解.

汾河河水样点均位于(Ca²++Mg²+)/HCO₃=1:1比值线以上(图7),说明河水中 Ca²+和 Mg²+除了来源于碳酸盐岩的溶解,还来源于其他矿物的溶解.如图8所示,岩溶水和大部分地表水样点位于(Ca²++Mg²+)/(HCO₃+SO₄²-)=1:1比值线附近,说明其 Ca²+、Mg²+和 SO₄²-主要来源于碳酸盐岩和石膏的溶解,其他点则位于(Ca²++Mg²+)/(HCO₃+SO₄²-)=1:1比值线以下,结合研究区地球化学背景,说明 SO₄²-除来源于石膏的溶解,还可能来源于汾河周边黄土层中硫化矿物的溶解[⁴⁴,⁴⁵].汾河上游主要地层岩性为中奥陶统含石膏的泥质白云岩,易溶的石膏使汾河水中硫酸根离子大大增加,汾河中下游进入临汾盆地,河谷宽阔,水流速度缓慢,加之流域内为黄土丘陵地貌,河流冲刷两岸黄土层矿物均可造成河水中水化学成分发生变化,而支流潇河、

昌源河及浍河明显在 $(Ca^{2+} + Mg^{2+})/(HCO_3^- + SO_4^{2-}) = 1:1$ 比值线以下 (图 8),说明其 SO_4^{2-} 还受到煤矿矿坑水污染 $[^{44,46}]$. 泉水样点均位于 $(Ca^{2+} + Mg^{2+})/(HCO_3^- + SO_4^{2-}) = 1:1$ 比值线上,泉水中 SO_4^{2-} 主要来源于碳酸盐岩地层中石膏的溶解.

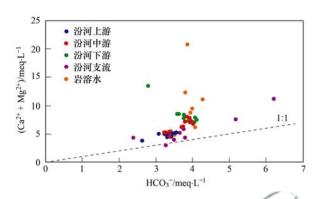


图 7 (Ca²⁺ + Mg²⁺) 与 HCO₃⁻ 关系曲线

Fig. 7 Relationship of (Ca²⁺ + Mg²⁺) and HCO₃

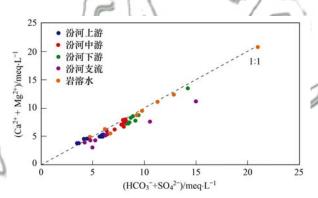


图 8 (Ca²⁺ + Mg²⁺)与(HCO₃ + SO₄ -)关系曲线

Fig. 8 Relationship of $(Ca^{2+} + Mg^{2+})$ and $(HCO_3^- + SO_4^{2-})$

2.4 氢氧同位素特征

受污染的 F26 点其 δ D 和 δ ¹⁸O 分别为 – 20.3‰ 和 - 0.06‰, 明显区别于其他点. 其余汾河干流地表 水 δD 介于 - 75.10‰ ~ - 51.10‰, 平均值为 -62. 60‰; δ¹δ0 介于 -10. 81‰ ~ -7. 09‰, 平均值 为 - 8.42‰(图 9). 除上游 F1 点岔上水文站点和 F3 雷鸣寺泉下游点外,其余水样点均位于当地大气 降水线(LMWL)下和全球大气降水线(GWML) 下[36,47],且远离当地大气降水线,表明干流地表水 主要受到了蒸发作用的影响,受大气降水影响较 小 $^{[36]}$. 从上游到下游汾河地表水 δ D 和 δ^{18} O 值逐渐 富集,这是由于从上游 F1 点到下游 F31 点,地形高 差 1432 m,下游河面较宽,地势平缓,地表水径流缓 慢,蒸发浓缩作用逐渐加强.同样的,沿汾河从上游 到下游的各支流的 δ D 和 δ ¹⁸O 值逐渐富集,蒸发浓 缩作用加强. 汾河支流地表水 δD 介于 - 68. 60% ~ -27.80‰,平均值为-54.98‰; δ¹⁸0介于-9.29‰ ~ - 3.85‰, 平均值为 - 7.28‰. 支流中昌源河

(Y5)的 δ D 和 δ ¹⁸O 值最高,说明昌源河受蒸发浓缩作用更强. 岩溶水 δ D 介于 - 78. 60% ~ - 58. 20% , 平均值为 - 71. 05%; δ ¹⁸ O 介于 - 10. 89% ~ -7. 64%, 平均值为 - 9. 72%. 雷鸣寺泉(S1)和广胜寺泉(S5)点均位于 GMWL线上,这两点均位于碳酸盐岩裸露区,主要接受降水补给. 其余点均位于GMWL线的右侧,受蒸发作用影响较大.

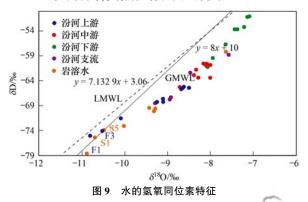
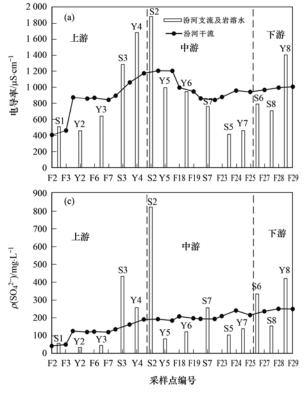


Fig. 9 Plot of δD and $\delta^{18}O$ values in the Fenhe River basin



2.5 支流及泉水汇入对汾河干流的影响

汾河中游各支流水化学含量最高(图 10),其次 是下游和上游. 中游主要是由于河道清水流量减少, 河水稀释能力低,而太原及临汾城市周边生活污染 及工农业用水排入导致水质较差. 各支流汇入干流 后,均会影响干流河水的水质,大部分支流河水水化 学含量低于干流的. 潇河(Y4)、昌源河(Y5)和浍河 (Y8)支流流域内工农业经济发达,大量灌溉用水及 工业生活污水排入河流,造成河流水质较差. 三支流 水汇入汾河后,导致汾河电导率、TDS、SO₄ 和 Cl 含量升高,经弥散、混合等,离子浓度升幅小,说明 干流在较大水量情况下,具有一定的纳污能力.另一 方面也反映了支流流量相对较小,水动力条件较弱, 水体自净能力差,水环境容量较小. 岩溶水是汾河清 水径流的重要组成部分,由于岩溶水(如郭庄泉 S7 点)的汇入,汾河水各含量有所降低,对河水水质有 净化稀释作用.

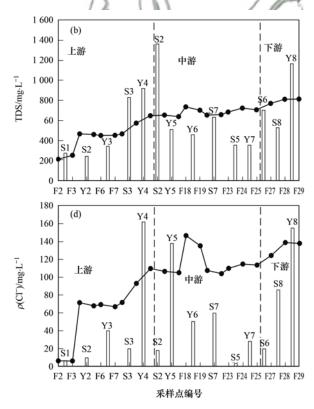


图 10 汾河地表水及岩溶水主要组分关系

Fig. 10 Relationship diagram of main components of surface water and karst water in the Fenhe River basin

3 结论

(1)汾河干流地表水整体上呈弱碱性,汾河干流主要水化学组分均呈现从上游到下游逐渐增加的趋势; 阴离子以 HCO_3^- 和 SO_4^{2-} 为主,逐渐过渡为以 SO_4^{2-} 和 HCO_3^- 为主,阳离子从以 Ca^{2+} 和 Mg^{2+} 逐步过渡到以 Ca^{2+} 、 Na^+ 和 Mg^{2+} 为主; 支流阳离子以

 Ca^{2+} 和 Na^{+} 为主, 阴离子以 HCO_{3}^{-} 、 Cl^{-} 和 SO_{4}^{2-} 为主, 岩溶水阳离子以 Ca^{2+} 和 Mg^{2+} 为主, 阴离子以 SO_{4}^{2-} 和 HCO_{3}^{-} 为主.

(2)枯水期汾河地表水主要来源于碳酸盐岩岩溶水、浅层地下水和工矿企业废水.水化学组成主要受岩石风化作用和蒸发结晶作用影响,大气降水对其影响较小. Na⁺和 K⁺主要来源于蒸发盐岩的溶解以

及周边黄土中的含 Na 矿物溶解,水体中 Ca^{2+} 、 Mg^{2+} 和 HCO_3 主要来源于碳酸盐岩的溶解, SO_4^{2-} 除来源于石膏的溶解,还可能来源于汾河周边黄土层中硫化矿物的溶解,部分地区可能还来源于煤矿矿坑水.

(3)汾河干流地表水 δD 和 δ¹⁸O 平均值分别为 -62.60‰和 -8.42‰,从上游到下游其 δD 和 δ¹⁸O 值逐渐富集,蒸发浓缩作用逐渐加强.支流昌源河受蒸发浓缩作用更强.碳酸盐岩裸露区的岩溶水点主要受降水影响,而其余点受蒸发作用影响较大.

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