

WINDLY WATER

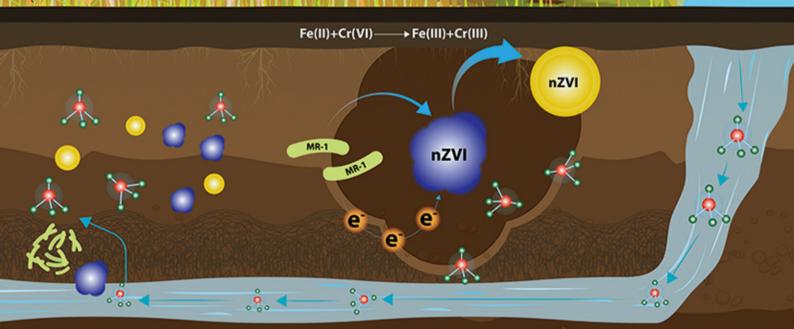
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电活性微生物激活生物质炭/零价铁协同钝化Cr(VI)及机制

廖聪坚,赵晓蕾,刘凯,钟松雄,李芳柏,方利平,叶挺进,石虎砚



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降水变化对荒漠草原土壤呼吸的影响

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摘要:全球气候变化使得降水格局发生显著改变.土壤呼吸作为土壤碳库向大气释放 CO₂ 的重要途径,其对降水变化的响应可能会影响陆地生态系统碳循环进程,并对全球气候变化产生反馈作用,但目前关于土壤呼吸对降水变化的响应没有一致的结论.以黄土高原西部荒漠草原为对象,通过野外降水控制实验减水 40%(-40%)、减水 20%(-20%)、自然降水、增水 20%(20%)和增水 40%(40%),探究降水变化对土壤呼吸动态的影响及其与土壤含水量、土壤温度、地上生物量、土壤有机碳、微生物量碳和碳氮比(有机碳总氮比)等因素的关系. 结果表明,在 3 a 期间不同降水处理下土壤呼吸的日变化呈现较一致的单峰和双峰模式. 土壤呼吸随降水量的增加均呈增加趋势,且相较对照,土壤呼吸在降水控制实验第二年(偏湿年份)和第三年(偏干年份)表现出显著差异,表明降水变化对土壤呼吸产生了遗留效应. 同时,相比对照,偏湿年土壤呼吸在。40%处理下显著最低,在 40%处理下显著最高,土壤呼吸对减水处理的负响应强于对增水处理的正响应;偏干年土壤呼吸在增水处理下显著高于对照,且对增水处理的正响应明显强于减水处理.此外,土壤含水量、地上生物量、土壤有机碳和碳氮比的增加而增加,随、生物量碳的增加而减少,其中土壤含水量对土壤呼吸的解释率最高,这表明土壤含水量是控制荒漠草原区土壤呼吸的主要环境因子.无论在偏湿或偏干年份,降水变化下,植物生物量输入幅度均低于土壤呼吸输出幅度,表明降水变化对土壤呼吸的影响可能对生态系统碳循环过程产生不同的影响,进而为区域碳预算评估提供参考.

关键词:降水处理; 土壤呼吸(SR); 地上生物量; 土壤碳库; 荒漠草原

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Influence of Precipitation Change on Soil Respiration in Desert Grassland

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Abstract: Global climate change has significantly changed precipitation patterns. Soil respiration (SR), as an important pathway through which CO₂ is released from the soil carbon pool into the atmosphere, may affect the carbon cycle process of terrestrial ecosystems and have a feedback effect on global climate change in response to precipitation change. However, at present there is limited understanding of how SR is affected by precipitation change. Field precipitation control experiments were conducted (with -40%, -20%, natural, 20%, and 40% precipitation) on desert grassland in the west of the Loess Plateau, to investigate the influence of precipitation change on SR dynamics and its relationship with soil water content, soil temperature, aboveground biomass, soil organic carbon, microbial biomass carbon, carbon-nitrogen ratio, and other factors. The results show that the diurnal variations of SR under different precipitation treatments were consistent in unimodal and bimodal models over three years. SR showed an increasing trend with added precipitation, relative to the control, and significant differences were observed between the second year (wetter) and the third year (drier) of the precipitation-manipulation experiment, indicating that precipitation changes had a legacy effect on SR. At the same time, SR was lowest under the -40% treatment and highest under the 40% treatment during the wetter year. The negative response of SR to precipitation exclusion treatments was stronger than the positive response to precipitation addition treatments. SR in drier years was significantly higher under precipitation addition treatments than the control, and the positive response of SR to increased precipitation treatment was significantly stronger than that under decreased precipitation treatment. In addition, soil water content, aboveground biomass, soil organic carbon, and carbon-nitrogen ratio were the environmental factors that obviously affected SR and increased with additional precipitation. SR increased with increases in soil water content, aboveground biomass, soil organic carbon, and carbon-nitrogen ratio, but decreased with increases in microbial biomass carbon. Among these factors, soil water content had the highest interpretation rate for SR, indicating that soil water content was the main environmental factor controlling SR in desert grassland. In both wetter and drier years, the amplitude of plant biomass input was lower than the amplitude of SR output under

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precipitation change, indicating that precipitation change may be unfavorable to soil carbon sequestration, especially in drier years, when precipitation change has a stronger influence on carbon pool output. Therefore, precipitation changes on SR in desert grassland in various dry and wet years may have different influences on the carbon cycle process of ecosystems, thus providing a reference for regional carbon budget assessment.

Key words: precipitation treatment; soil respiration (SR); aboveground biomass; soil carbon pool; desert grassland

全球气候变化模型预测 21 世纪下半叶降水格局将发生重大改变^[1],这将对陆地生态系统碳循环过程产生显著影响^[2,3]. 土壤呼吸(soil respiration, SR)作为全球碳循环中最关键的组分,其微小变化都会显著影响大气 CO₂ 浓度,进而影响区域和全球的碳平衡和气候变化^[4]. 在以水分为主要限制因子的半干旱区, SR 主要受土壤水分和土壤温度的控制^[5,6],所以降水变化将通过改变土壤微气象条件(如土壤水分和土壤温度)影响土壤 CO₂ 通量,进而增加对区域碳预算评估的不确定性. 因此,揭示 SR 对降水变化的响应将在很大程度上提升人们对全球和区域碳平衡的了解.

目前,国内外学者在不同生态系统类型下开展 了大量控制降水实验,研究了 SR 对降水变化的响 应,包括 SR 的日变化及其年变化[7~12]. 从 SR 对降 水变化响应的日变化方面来看,目前研究结果并不 一致,例如,在高寒草甸生长季, SR 的日变化呈双 峰模式,最大值出现在 13:00~14:00 和 16:00^[7]; 而在中国内蒙古西部的荒漠草原区,不同降水条件 下的 SR 日变化随时间呈单峰模式,峰值在 09:00 13:00^[8]. 从年变化来看, 在不同干湿条件下的 SR 对降水变化的响应规律并不相同. 例如,在西双版纳 热带雨林中,由于降水丰富,土壤含水量对土壤呼吸 的影响远小于土壤温度[9];在高寒湿地,相比于其 他环境因素,土壤水分是影响土壤呼吸的主要因素, 并与土壤呼吸呈极显著正相关[10];在加利福尼亚 一年生草原中,干季 SR 对降水增加的响应显著,而 在湿季对降水变化无明显响应[11];在中国南部3 种常见的亚热带森林区(包括马尾松林,针阔混交 林和季风常绿阔叶林),旱季 SR 随降水量的增加而 增加,而湿季 SR 随降水的减少而增加[12].可见,在 干旱条件下 SR 对降水增加的响应比湿润条件下更 明显[13,14]. 尽管在多个研究中已经报道了多地点和 多年控制降水实验中 SR 对降水变化的响应呈非线 性模式[15,16],但关于 SR 对降水变化的响应规律并 不一致,以及降水变化下主导土壤呼吸的环境因素 亦不明确. 因此,仍然需要从多梯度降水实验中获得 大量证据来表征这一响应模式[17],因为验证这一关 系对在干旱半干旱地区模拟气候变化下生态系统碳 循环的稳定性至关重要.

本研究采用野外控制实验的方法,模拟降水变

化对荒漠草原 SR 的日变化和年变化的影响,研究降水梯度下 SR 的响应规律及其与环境因素(土壤温度、土壤湿度、地上生物量、土壤有机碳、微生物量碳和碳氮比等)的关系,以期为全球气候变化背景下荒漠生态系统中碳库变化及区域碳预算评估提供科学依据.

1 材料与方法

1.1 研究区概况

本实验地点位于中国科学院西北生态环境资源 研究院皋兰生态与农业综合试验站(36°13"N, 103°47″E; 海拔为1 780 m). 气候类型为半干旱大 陆性气候. 年平均降水量为 263 mm, 其中 70% 的降 水发生在5~9月. 年平均气温为8.4℃,月平均最 高气温为 20.7℃ (7月),月平均最低气温为 -9.1℃(1月).平均土壤有机碳和总氮的含量分别为 7.5 g·kg⁻¹和1.0 g·kg⁻¹, pH 为 8.52. 该地区土壤 由风积黄土母质发育而成,具有均匀的粉砂壤土质 地,在联合国粮农组织/联合国教科文组织中被分类 为普通钙积土. 该地以典型的荒漠草原植被为主,包 括小半灌木的灌木:亚菊(Ajania fruticulosa)、红砂 (Reaumuria songarica)和多年生草本短花针茅 (Stipa breviflora). 常见的伴生多年生草本植物有: 骆驼蓬 (Peganum harmala)、阿尔泰狗哇花 (Heteropappus altaicus)、蝎虎驼蹄瓣(Zygophyllum mucronatum)、茵陈蒿(Artemisia capillaris)和糙隐子 草(Cleistogenes squarrosa). 一年生草本植物丰度较 低,常见的物种包括刺沙蓬(Salsola ruthenica)、狗尾 草(Setaria viridis)、锋芒草(Tragusracemosus)和地锦 (Euphorbia humifusa)等.

1.2 实验设计

本实验采用完全随机区组设计,根据研究区近50 a 的降水变率(-41.1%~39.2%)设置5个降水梯度,分别为在自然降水量的基础上减水40%(-40%)、减水20%(-20%)、增水20%(20%)、增水40%(40%)和对照(自然降水).在皋兰站植被分布较均匀的地段,每个降水处理重复3次,共设置15个2.5m×2.5m的固定样方,每个样方四周用防锈铁皮埋深30cm,露出地表15cm,防止样方内外表层水分径向流动.降水频率和时间没有改变.在2013~2015年每年5~9月实施降水处理^[18].

减水处理:利用遮雨棚截留降雨,减少每次降雨量的20%和40%.遮雨棚由固定位置的金属框架组成,顶部装有透明的 V 型丙烯酸条带(夹角为60°,宽为10 cm),遮雨板通过遮挡样方面积的20%和40%形成减水区^[19].遮雨棚的平均高度为0.50m.与塑料或聚氯乙烯相比,丙烯酸的透光率较高,对光合有效辐射的拦截量也比其他遮挡设计要少.由于空气流动不受限制,温室效应可被消除.并在遮雨棚框架较低端安装集雨槽,将截留的降水收集到密闭容器中.

增水处理:在每次降水事件发生后 8 h 内,通过 将减水处理样方中被拦截的降水手动均匀地添加到 增水处理的样方中,以达到增水 20% 和 40% 处理.

对照处理:无遮雨棚和增水处理,其他环境条件 与减水和增水处理保持一致.

1.3 植物生物量

在生物量高峰时期(9月初)使用收获法对地上生物量进行了评估. 在每个大样方(2.5 m×2.5 m)中排除固定植物调查的1 m×1 m的小样方外,随机选择两个0.5 m×0.5 m的样方测生物量,且每年测生物量的位置选前年没干扰的位置,所有植物接物种剪切到土壤表面. 在65℃的烘箱中烘干48 h 后称量干重以测定各物种生物量(g·m⁻²). 群落的地上生物量(g·m⁻²)是根据所有物种地上生物量的总和估算的. 另外,从固定植物调查的1 m×1 m 的小样方中收集凋落物并用烘干法进行凋落物测定[19].

1.4 土壤数据

0~20 cm 深的土壤含水量在 2013、2014 和 2015 年选择在大样方(2.5 m×2.5 m)中固定调查 植物性状样方的周围无植物生长的空隙用土钻法进行测量.5 cm 深的土壤温度通过与 LI-6400-09 土壤室(Li-Cor, Inc., Lincoln, NE, USA) 连接的热电偶探头测定.

土壤样品分别于 2013、2014 和 2015 年每年 8 月在 15 个样方中和测土壤含水量的样品同时采集, 土壤样品采集量以保证分析使用即可, 采集完土壤样品后进行土壤回填, 并做好标记, 以减少对样方中植物和土壤的影响. 在每个样方中用土钻(直径 5 cm)分别在深度 0~5、5~10 和 10~20 cm 取 3 个土样. 手动去除可见的根和石头后, 先将土壤通过 2 mm 的筛, 再将通过 0.25 mm 筛后的样品自然风干用于化学分析.

采用氯仿熏蒸法测定土壤微生物量碳和氮[20]. 土壤有机碳的测定采用 $K_2Cr_2O_7-H_2SO_4$ 氧化法;总氮采用微量凯氏定氮法测定;速效磷采用 Bray 法进行测定;速效钾采用乙酸铵提取-火焰光度法进

行测定;土壤 pH 值由玻璃电极测定,电导率由电导率仪在土水比率为1:2.5 的情况下测定^[21].

1.5 SR 的测量

为了测定 SR,在每个样方(2.5 m×2.5 m)两个对角线位置,将两个内径为11 cm和高为5 cm的聚氯乙烯项圈插入土壤3 cm.至少在测量前1 d要人工将土壤项圈内的所有活植物取出.使用 LI-6400-09 的土壤室(Li-Cor, Inc., Lincoln, NE, USA)测定 SR 的值.每个项圈中观察时间为2~3 min.将每个样方中的两个项圈的值取平均作为一个重复. SR 在2013、2014和2015年的6~9月每月中下旬选择无雨的2d,在08:00~18:00每2h测定一次.最后,根据测量日前一周的天气情况选择2013年的7月28日和8月21日、2014年的6月27日和7月15日、2015年的7月21日和8月28日的土壤呼吸数据作为本研究的数据.

1.6 数据分析

极端降雨年份被定义为统计学上罕见发生的年份^[22],先前已有研究将小于历史年降雨量第 10 百分位与大于第 90 百分位定义为极端干旱和极端湿润,第 45 ~55 百分位定义为正常年份^[23,24],利用皋兰气象站过去 50 a(1957 ~ 2012 年)生长季降雨量的概率密度函数(probability density functions)和实验期间生长季降雨量,确定 2013、2014 和 2015 年分别属于正常年份、偏湿年份和偏干年份(图 1).

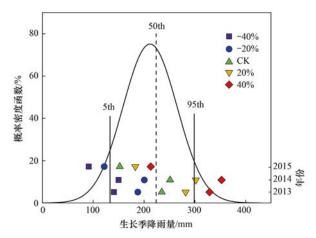


图 1 55 a(1957~2012)年生长季降雨量的概率密度函数

Fig. 1 Probability density function of precipitation during the growing season, over 55 years (1957-2012)

单因素方差分析用来确定降水处理对土壤温度、土壤湿度、地上生物量、土壤有机碳、微生物量碳和碳氮比等和年 SR 的影响,差异显著(P<0.05)则采用LSD 检验进行事后多重比较. 最后,采用回归分析确定降水和环境因子(土壤温度、土壤湿度、地上生物量、土壤有机碳、微生物量碳和碳氮比等)与 SR 之间的关系. 所有统计分析均采用 SPSS 25.0 进行.

2 结果与分析

2.1 降水变化对 SR 的影响

2.1.1 SR 的日变化

由图 2 可知,在 2013 年,各降水处理下的 SR 日变化均呈双峰模式,峰值均出现在 10:00 和 15:00,且 7 月和 8 月较高的峰值分别出现在 15:00 和 10:00;在 2014 年,各降水处理下的 SR 日变化均呈单峰模式,6 月和 7 月的峰值分别出现在 10:00~13:00 和 13:00~15:00;在 2015 年,SR 的日变化在 7 月的 08:00 出现最大值,而 8 月的峰值基本上也出现在 08:00.

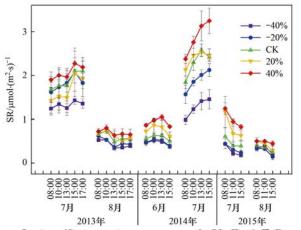
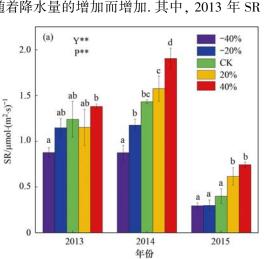


图 2 3 a 不同降水处理下生长季土壤呼吸的日变化(n=3)
Fig. 2 Diurnal change in soil respiration during the growing season of three years, under different precipitation treatments (n=3)

2.1.2 降水处理下 SR 的年际变化

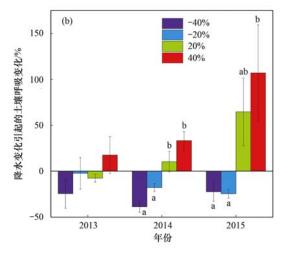
由图 3(a) 可知, SR 的年际变化受降水处理(P < 0.01) 和年际变化(P < 0.01) 的显著影响,且在 3 a 中均随着降水量的增加而增加.其中, 2013 年 SR



只在 40% 和 - 40% 处理之间表现出显著差异(P < 0.05); 而在 2014 年, SR 在 40% 和 - 40% 处理下与对照相比都存在显著差异(P < 0.01), 尤其在 2015 年,增水处理下的 SR 均显著高于对照(P < 0.05). 另外, 从图 3(b)中可得出,在 2013 年 SR 在增水与减水处理之间没有表现出显著差异(P > 0.05); 而在 2014 年相比于增水处理 SR 对减水处理的负响应更显著(SR 在减水处理下的减幅比在增水处理下的增幅高 13.65%,P < 0.05); 在 2015 年相比于减水处理 SR 对增水处理的正响应更显著(SR 在增水处理下的增幅比在减水处理下的减幅高 88.71%,P < 0.05),这反映出在干、湿年份降水处理对 SR 的影响是不一致的.

2.2 降水处理下环境因素的变化

土壤温度在3 a 期间都没有受到降水处理的显 著影响(P>0.05), 见表 1, 而土壤含水量在 2013 年没有受到降水处理的影响(P > 0.05),但在 2014 年和2015年却受到了降水处理的显著影响(P< 0.05),且在40%处理下的土壤含水量都明显高于 减水处理下的土壤含水量(P<0.05), 见表 2. 此 外,降水处理对速效磷(AP)、速效钾(AK)、碳氮比 (C/N)、有机碳(SOC)和微生物量碳(MBC)的影响 不显著,对 pH、总氮(TN)、电导率(EC)、地上生物 量(AGB)、凋落物(L)和微生物量氮(MBN)的影响 显著(P<0.05). 其中, pH 在 20% 处理下显著高于 -40% 处理(P < 0.05); TN 在对照处理下显著高于 -20% (P < 0.01); EC 在 -40% 处理下显著高于 20% (P < 0.05); AGB 在增水处理和对照处理下分 别显著高于减水处理和 -40% 处理(P < 0.05,P <0.01),且在20%处理下呈更高趋势; L 在40%处理



(a)不同降水处理下土壤呼吸的年变化,**表示 P < 0.01,(b)3 a不同降水处理下降水引起的土壤呼吸变化; 竖线代表均值的标准误差,n = 6,不同小写字母表示处理间差异显著(P < 0.05)

图 3 不同降水处理下和降水引起的土壤呼吸年变化

Fig. 3 Annual variation of precipitation-induced soil respiration under different precipitation treatments

下显著高于 – 40% (*P* < 0. 01); MBN 在 – 40% 处理 下显著低于其他处理(*P* < 0. 05), 见表 3.

2.3 SR 和环境因素之间的关系

由图 4 可以看出,降水量与土壤含水量、土壤含水量与呼吸之间存在着显著的正相关关系,但降水量与土壤温度、土壤温度和呼吸之间无显著相关关系. SOC、AGB、C/N 与降水量呈显著正相关关系[图

5(a)、5(c)和5(g)],而 MBC 与降水量呈显著负相 关关系[图5(e)];但 L、MBN、TN、AP、AK、EC 和 pH 与降水量和 SR 无显著相关关系(图6).其 中,降水量对 AGB 解释率最高(30%),其次是 C/N (24%)、MBC(17%)和 SOC(10%);而 C/N 对 SR 解释率最高(27%),其次是 SOC(20%)、AGB (16%)和 MBC(15%)(图5).

表 1 降水处理下 5 cm 深土壤温度的单因素方差分析 1 (均值 ± 标准误差,n=3)

Table 1 One-way ANOVA of soil temperature, at a depth of 5 cm, under different precipitation treatments (Mean ± s. e., n = 3)

年份			降水处理		
平切	-40%	-20%	CK	20%	40%
2013	24. 99 ± 0. 48a	24. 48 ± 0. 38a	$25.04 \pm 0.47a$	$24.68 \pm 0.93a$	24. 87 ± 0. 84a
2014	$26.38 \pm 0.55a$	25. $70 \pm 0.39a$	$26.33 \pm 0.38a$	25. 97 \pm 0. 40a	$26.65 \pm 0.50a$
2015	$25.42 \pm 0.31a$	$24.96 \pm 0.54a$	24. $80 \pm 0.36a$	24. 67 \pm 0. 18a	$25.41 \pm 0.64a$

¹⁾不同小写字母代表降水处理间的差异显著 (P < 0.05),下同

表 2 降水处理下 20 cm 深的土壤含水量的单因素方差分析(均值 \pm 标准误差,n=3)

Table 2 One-way ANOVA of soil water content, at a depth of 20 cm, under different precipitation treatments (Mean ± s. e., n=

年份			降水处理	=	1
平彻	-40%	-20%	CK	20%	40%
2013	$4.04 \pm 0.68a$	4. 96 ± 0. 13a	4. 87 ± 0. 69a	5. 04 ± 0. 81a	5. 78 ± 0. 62a
2014	$6.66 \pm 0.13a$	6.29 ± 0.36 ab	7. 29 ± 0.35 abc	7.53 ± 0.60 be	$8.25 \pm 0.06c$
2015	2. 78 ± 0. 15a	2. 73 ± 0. 24a	3.60 ± 0.19 ab	$3.46 \pm 0.52ab$	4.18 ± 0.45 b

表 3 降水处理下土壤和植物因素的变化(均值 ±标准误差,n=9)

Table 3 Changes in soil and plant factors under different precipitation treatments (Mean \pm s. e., n = 9)

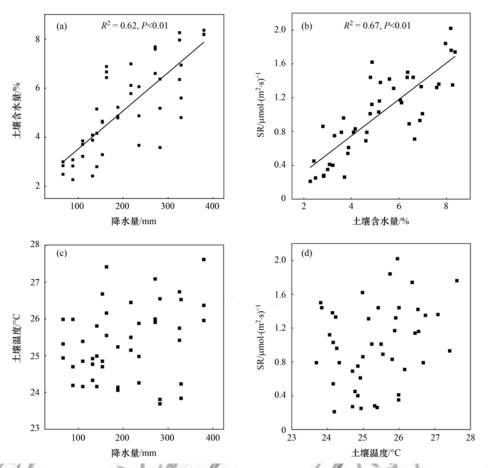
1. 梅和拉伽耳塞		1/1/1/1	降水处理	1	
土壤和植物因素	-40%	-20%	CK	20%	40%
pH	8.38 ± 0.043a	8.43 ± 0.03ab	8.43 ± 0.03ab	8.50 ± 0.03 b	8.48 ± 0.04ab
总氮/g·kg ⁻¹	$1.76 \pm 0.04 \mathrm{bc}$	$1.51 \pm 0.03a$	$1.79 \pm 0.07 \mathrm{c}$	$1.57 \pm 0.11ab$	$1.68\pm0.05\mathrm{abc}$
土壤有机碳/g·kg ⁻¹	13.50 ± 0.34	12.75 ± 0.41	14.14 ± 0.65	12.62 ± 0.99	12.82 ± 0.63
碳氮比	7.71 ± 0.31	8.49 ± 0.29	7.94 ± 0.37	8.09 ± 0.36	7.63 ± 0.31
速效磷/mg·kg ⁻¹	6.05 ± 0.41	6.80 ± 0.25	6.57 ± 0.49	5.61 ± 0.52	5.86 ± 0.51
速效钾/mg·kg ⁻¹	166.94 ± 8.83	165.28 ± 7.69	156.94 ± 2.35	155.28 ± 5.08	167.78 ± 11.00
电导率/μS·cm ⁻¹	$608.81 \pm 146.35\mathrm{b}$	$509.11 \pm 81.73 \mathrm{ab}$	$415.44 \pm 41.08 ab$	$320.36 \pm 15.23a$	391.08 ± 53.18 ab
微生物量氮/mg·kg -1	$10.65 \pm 1.39a$	$16.66 \pm 2.02b$	$20.48 \pm 2.23 \mathrm{b}$	$15.08 \pm 1.59 ab$	19.60 ± 2.06 b
微生物量碳/mg·kg -1	170.00 ± 23.89	212.53 ± 28.84	235.27 ± 29.34	178.45 ± 26.63	209.88 ± 38.46
地上生物量/g	$43.91 \pm 7.29a$	78.54 ± 8.86 b	$103.35 \pm 9.79 {\rm bc}$	$125.99 \pm 11.57c$	$115.07 \pm 11.36c$
凋落物/g	$1.82 \pm 0.16a$	2.50 ± 0.33 ab	$2.20 \pm 0.25 ab$	$2.49 \pm 0.49 ab$	$3.33 \pm 0.50 \mathrm{b}$

3 讨论

3.1 降水变化对 SR 日变化特征的影响及其控制 因素

SR 的日变化呈单峰或双峰的模式,SR 的日变化(图2)与气温日变化(图7)基本一致.在2013年8月,SR 呈双峰模式,其中较高的峰值出现在10:00,较低的峰值出现在15:00,而气温最高值出现在13:00.这可能是因为在8月8~21日之间没有发生降水事件(图8),土壤中的水分不足,而早上10:00温度不高且有露水补充土壤中的水分,所以呼吸出现了最高值,而随着温度继续增加,土壤中的水分蒸发加强,影响地下可溶性有机质的扩散和分

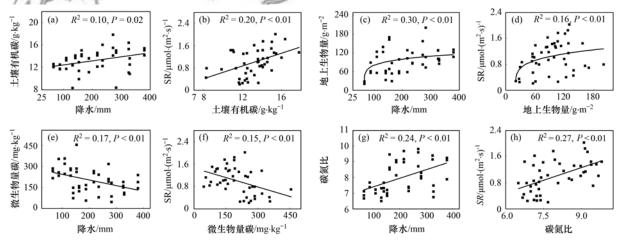
解、减缓微生物和根系的生命活动,从而减弱土壤中的根呼吸和异养呼吸^[25],而当 15:00 时,温度逐渐降低,土壤中水分蒸发减弱,减缓了对地下可溶性有机质扩散和分解的影响,同时在一定程度上促进了微生物和根系活性,所以 SR 有了一定程度的增强^[25].在 2015 年 7 月,SR 最高值出现在 08:00,且随着时间变化呈下降趋势(图 2),而气温呈上升趋势(图 7). SR 最高值出现在 08:00 的原因如下,在测量 SR 日的 3 d前(即 7 月 18 日前)降水事件频繁(图 8),较多的降雨增加了土壤水分,且西北地区昼夜温差大,使早晨露水较多,短时间(1 h)内的替代效应对土壤呼吸产生了激发^[26];另外,2015 年是偏于年份,降水整体较少,所以土壤呼吸对土壤含水



(a)和(c)分别表示土壤含水量和土壤温度随降水量的变化;(b)和(d)分别表示土壤呼吸随土壤含水量和土壤温度的变化

图 4 土壤含水量和土壤温度随着降水变化对土壤呼吸的影响

Fig. 4 Effects of soil water content and soil temperature on soil respiration with precipitation



(a)、(c)、(e)和(g)分别表示生长季土壤有机碳、地上生物量、微生物量碳和碳氮比与降水的关系;(b)、(d)、(f))和(h)分别表示生长季土壤有机碳、地上生物量、微生物量碳和碳氮比与土壤呼吸的关系;n=9

图 5 环境因素随着降水变化对土壤呼吸产生显著影响

Fig. 5 Environmental factors that had a significant effect on soil respiration with changes in precipitation

量的变化较敏感,因此,在早晨土壤含水量充足的情况下,土壤可溶性有机物的有效性和流动性增强,有机物的增加会促进微生物活性并提供给微生物繁殖足够的底物,进而增加土壤呼吸^[27,28].但随着温度的增加,蒸发加强,对植物和微生物产生一定的水分

胁迫,可能抵消了增温对 SR 带来的促进作用,从而间接导致了 SR 的减弱^[29,30];也可能与降雨和温度对微生物活性的直接作用有关,如短期内较大的降雨会直接影响微生物的渗透势,限制土壤微生物的活性和数量^[31],因此影响了有机物的分解速率和微

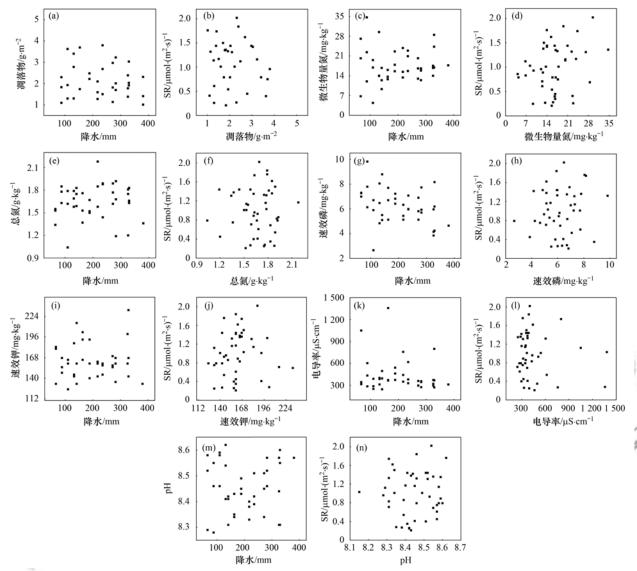


图 6 环境因素随着降水变化对土壤呼吸无显著影响

Fig. 6 Environmental factors that had no significant effect on soil respiration with changes in precipitation

生物的呼吸速率;另外,增温可能通过降低土壤含水量与充足的降雨产生拮抗的交互作用影响微生物群落结构直接影响微生物的呼吸^[32].因此,先前降雨事件和测量当日气温变化是影响 SR 日变化的主要因素.在 2015 年 8 月, SR 日变化随着时间的变化呈下降的趋势,且峰值出现在 08:00 左右,而气温随着时间的变化呈上升的趋势,且峰值出现在 11:00~15:00.因为在08:00温度较低蒸发很弱且露水充足,又因为呼吸测量前一周内又出现了较大的降水事件(图 8),所以这一时间段呼吸值较高,但随着温度的增加,土壤呼吸逐渐降低.因此,在该月降水和气温对 SR 日变化影响的具体过程和 2015 年 7 月类似.

一般地,土壤温度和湿度通常是控制 SR 最重要的环境因子^[5]. SR 在湿润的环境下对土壤温度的

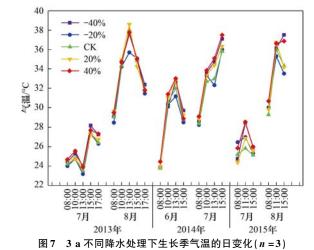


Fig. 7 Diurnal change of air temperature during the growing season of three years, under different precipitation treatments (n = 3)

变化更敏感[33],而在干燥的环境下对土壤含水量的

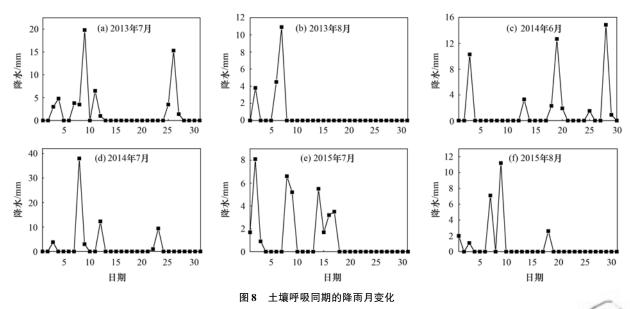


Fig. 8 Monthly variation in soil respiration during the same period of rainfall

变化更敏感^[13,14].内蒙古干旱半干旱典型草原上的研究也指出,SR与土壤含水量、土壤温度均具有显著的正相关关系,但主要由土壤含水量决定^[30],这与本研究中的结果一致,即 SR与土壤含水量之间存在着显著的正相关关系(图 4),而与土壤温度之间的关系并不明显,这反映出在干旱条件下,土壤含水量是控制 SR 的一个重要因素.在本文中温度对呼吸没有产生明显作用的原因可能主要是在干旱的环境中,土壤水分是限制温带草原生态系统最主要的因子^[34],因此增温将导致土壤进一步干旱化,然后对植物产生水分胁迫,从而抵消了增温对 SR 带来的促进作用,甚至对 SR 产生抑制作用^[29].

3.2 降水变化对 SR 年际变化的影响及其控制 因素

在降水处理后的第一年,增减水处理下的 SR 与对照并无显著差异(P>0.05),而在第二年和第 三年相比对照表现出显著差异(P<0.05),其中,降 水处理第二年(偏湿年份), 40% 处理下 SR 显著增 加 33.0%, -40% 处理显著降低 SR 38.9%; 第三 年(偏干年份), 40% 处理显著增加 SR 达 86.3%,而 减水处理并未显著降低 SR(图3). 这表明,降水变 化对 SR 的影响具有遗留效应,也反映出 SR 对降水 变化的响应与当年降水背景值有关,即偏湿年 SR 对减水的响应更明显,而偏干年对增水的响应更明 显. SR 主要由自养呼吸和异养呼吸所组成,其中,自 养呼吸主要为根系呼吸和根际微生物呼吸,而异养 呼吸包含微生物呼吸.有研究表明,降水变化对半干 早草原地上生物量的影响具有滞后性[35],因此,降 水变化对植物生长状况的滞后性可能导致根系呼吸 及根际微生物呼吸也表现出滞后性. 另外,也有研究 发现,土壤微生物群落结构及多样性对气候变化的响应存在滞后性^[36,37],因此,土壤微生物群落对气候变化的滞后响应可能会导致异养呼吸也表现出滞后响应.由此可知,植物群落和土壤微生物群落对降水变化的滞后响应可能是形成 SR 表现出遗留效应的主要原因.

SR 在不同干湿年对降水变化的响应规律与多 数的研究结果一致[13,25,38],即偏湿条件下 SR 对降 水减少更敏感,而偏干条件下对降水增加更敏感.偏 湿年份,降水减少对呼吸的作用比降水增加更明显, 可能与偏湿年份自然降水量较多,植物生长以及微 生物活性并未受到水分限制,因此增加降水对植物 和微生物活性的促进作用有限,从而使 SR 也并未 表现出更大幅度的增加[25,39]. 但减少降水,对植物 生长和微生物活性的抑制作用更明显,本研究结果 也反映出,植物地上生物量在-40%条件下显著最 低. 有研究发现植物 AGB 与 SR 存在显著正相关关 系,因此,在评价生态系统碳循环时还应考虑植物生 长对 SR 的影响[40], 植物生长的降低可能直接导致 根系自养呼吸的减弱,并间接通过碳输入的减少抑 制微生物活性,进而也降低异养呼吸[41,42],本研究 结果也反映出 SOC、C/N 与 SR 存在显著的正相关 关系. 这与在内蒙古 3 种典型草原的研究结果一致, 降水后,微生物可用底物的增加,加快了其土壤呼吸 速率[43],有研究也发现在草地和森林土壤中,微生 物呼吸速率与底物呈正相关关系,此外,底物的化学 计量比(可溶性有机碳氮比)也可能会影响微生物 对底物分解的有效性,从而影响呼吸中的异养呼 吸[44]. 然而,另一项位于内蒙古半干旱草原的研究 发现随着降水增加基质有效性降低会进一步增强土

壤呼吸,这可能是由于可溶解碳在干燥的土壤中扩 散和周转率较低,而在湿土中的分解率较高,所以在 干燥的土中,可溶性有机碳较多从而有助于增强土 壤呼吸[45]. 偏干年份, SR 对增水的响应更明显. 偏 干年份的自然降水量本身偏少,甚至接近极端干旱, 所以自然降水条件本身对植物生长和微生物活性产 生抑制作用[41,42]. 因此,降水的进一步减少对 SR 的 抑制作用有限.有研究表明, SR 在湿润的环境下对 土壤温度的变化更敏感[33],而在干燥的环境下对土 壤含水量的变化更敏感[13,14]. 本研究结果也表明, 土壤含水量与 SR 存在显著正相关关系,且解释率 达67%.这可能是由于在干土再润湿后,土壤会通 过裂解活微生物细胞释放更多的碳,细胞溶质的释 放和受保护的有机物暴露在微生物中及碳的可用性 都可能是增加碳释放的原因之一[46].因此,干旱条 件下土壤水分增加对植物和微生物活性的刺激作用 尤为明显,进而促进根系呼吸和微生物呼吸,最终导 致 SR 对增水的响应更为明显[10,13,14]. 由于每年仅 通过两次的观测数据来分析不同处理间 SR 的显著 性以及年际差异确实可能会因观测次数较少而影响 结果的比较,在今后的研究中会完善相应的研究.

3.3 降水变化对土壤碳库的可能影响

土壤碳库作为固定大气 CO, 的主要形式,其含 量的变化调控着大气 CO, 的浓度, 进而影响全球气 候变化[4]. 碳库主要包括植物生物量碳库和土壤碳 库(土壤有机碳库、微生物生物量碳库)[24,25]. 本研 究结果表明,在正常年,增水20%可能有利于土壤 碳固存,由于植物生物量输入幅度大于 SR 输出幅 度,而其他降水处理均不利于碳固碳,由于植物生物 量输入的幅度小于 SR 输出的幅度; 无论在偏湿还 是偏干年, SR 输出幅度均大于植物生物量输入幅 度,表明降水变化并不利于土壤碳固存;另外,偏干 年 SR 输出幅度比偏湿和正常年份均要明显,表明 在降水偏少的背景下降水变化对土壤碳库的扰动作 用更强. 综上所述, AGB 和 SR 对降水变化的不同 响应可能会影响荒漠草原的碳输入和输出,从而对 土壤碳平衡产生显著的影响.未来,极端湿润或干旱 可能会使荒漠草原生态系统从碳平衡变成碳负债. 本研究将有助于对干旱半干旱生态系统进行碳预算 和碳循环的评估,长时间序列控制降水实验对于荒 漠草原生态系统碳平衡可能提供更好地支撑.

4 结论

通过在荒漠草原上为期 3 a 的控制降水实验, 本研究发现不同降水处理下 SR 的日变化随时间呈 单峰和双峰模式,这主要与土壤水分状况和气温有 关. SR 的年际变化整体上均随降水的增加呈增加趋势,偏湿年减少降水对 SR 的抑制作用更为明显,而偏干年增加降水对 SR 的促进作用尤为明显. SR 主要受土壤含水量、土壤有机碳、地上生物量、C/N、微生物生物量碳的影响. 降水变化下,碳库输出高于碳库输入,尤其在偏干年份更为明显,这可能会使荒漠草原土壤成为碳源.

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《环境科学》多项引证指标名列前茅

2020年12月29日,中国科学技术信息研究所在中国科技论文统计结果发布会上公布了2019年度中国科技论文统计结果. 统计结果显示《环境科学》2019年度总被引频次12057,影响因子2.256,多项引证指标位居环境科学技术及资源科学技术类科技期刊前列.

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