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目 次

	· # (1012)
长江经济带 $PM_{2.5}$ 分布格局演变及其影响因素 黄小刚,赵景波,曹军骥,辛未中国东海近海岛屿冬季与夏季气溶胶中水溶性离子化学组分特征及来源解析 方言,曹芳,范美益,章炎阳泉市秋冬季 $PM_{2.5}$ 化学组分及来源分析 王成,闫雨龙,谢凯,李如梅,徐扬,彭泰安市夏季 $PM_{2.5}$ 中正构烷烃和糖类化合物的化学组成及其来源	:令(1013) :龢(1025)
阳泉市秋冬季 PM., 化学组分及来源分析 ····································	林(1025)
泰安市夏季 PM _{2.5} 中正构烷烃和糖类化合物的化学组成及其来源	•••
	璇(1045)
再出秋李天气颗粒初埋化符性 下逸韵,银燕,土红磊,阴 业碎区与滚胺业学属度特征及其与颗粒物浓度的相关性	、魁(1056) - 睉(1067)
淄博市重点工业行业 VOCs 排放特征 ····································	文(1007)
鄂州市大气 VOCs 污染特征及来源解析	1新(1085)
泰安市夏季 PM _{2.5} 中正构烷烃和糖类化合物的化学组成及其来源	字(1093)
基丁工乙以住的金属包衣业 VOCs 行架付值	八氏(1099)
··················· 牛真真,孔少飞,严沁,郑淑睿,郑煌,曾昕,姚立全,吴剑,张颖,吴方琪,程溢,覃思,刘玺,燕莹莹,祁士	:华(1107)
精细化工园区工艺过程 VOCs 产生量核算方法····································	军 (1116)
2017 年春李常州 HONO 观测及对大气氧化能力影响的评估 ····································	年(1123)
中国工住机械使用付征及共耗气排放趋势	: 禺(1132)
在用汽油和柴油车排放颗粒物的粒径分布特征实测	
王瑞宁, 胡磬遥, 任洪娟, 马冬, 徐冲, 赵玺乾, 王孟昊, 徐为标, 安静宇, 黄	成(1151)
参数选取对畜禽养殖业大气氨排放的影响;以长三角地区为例 ************************************	··· ·
乌海市煤矿区及周边春季降尘污染特征及来源分析	·利(1138) ·音(1167)
长江中下游地区丰水期刊、湖水氢氧回位系组成特值 学静,关华武,周水独,赵甲华,土皖龙,祭水久,负斌,陈芠,初	^ (11/6)
伊洛河流域河水来源及水化学组成控制因素 刘松韬 张东 李玉红 杨锦媚 邹霜 干永涛 黄兴字 张忠义 杨伟 曹保	(至 (1184)
城镇化进程中新疆塔城盆地浅层地下水化学演变特征及成因 ····································	釜(1197)
基丁 Sentinei-2 MSI 影像的河湖系统小体态仔初至间开开遥恐监测:以安徽省开壶湖马连妆长江权为例 工作有,工杰, 在玉苗十高原高浑浊水体 CDOM 光学特性及影响因素	.坏(1207) : 涛(1217)
黄土高原高浑浊水体 CDOM 光学特性及影响因素 梁晓文,邵田田,王人工强制混合充氧及诱导自然混合对水源水库水质改善效果分析	
发生。	军 (1227)
滤速与水质对似温含铁锰氨地下水中氨去除的影响····································	·冬(1236) ·举(1246)
汤肝河着牛砫藻群落及其与环境因子的关系	1240)
盐龙湖水源生态净化系统 FG 和 MBFG 演替特征及水质响应性评价 ········ 王莲,李璇,马卫星,邹立航,赵强强,丁成,吴向	月阳(1265)
一、映作区口盆丘小流现刻解和出形态及流大通重	谷(12/6)
化肥减量配施生物灰对紫色土坡耕地解流矢的影响 罗东海,土于芳,龙巢,严冬春,徐国鑫,李娇,局利田供办性方解石作为泛性覆盖材料控制水体内酒碟的移放	i 明(1286) s 龢(1206)
不用的数据为所有作为信任设量材料证明外径内部辨明样及 ************************************	娜(1290)
3种典型多孔高温改性固废材料对磺胺二甲嘧啶的吸附特性 王静,朱晓丽,韩自玉,胡健,秦之瑞,焦文	涛(1319)
新制和老化微塑料对多溴联苯醚的吸附	超(1329)
化肥减量配施生物灰对紧巴土玻耕地瞬流失的影响	. 男(1338)
·····································	飞 (1346)
一个一个一个一个一个一个一个一个一个一个一个一个一个一个一个一个一个一个一个	(锦(1357)
北京某污水处理厂及受纳水体中典型有机磷酸酯的污染特征和风险评估 ············· 张振飞, 吕佳佩, 裴莹莹, 王春英, 郭昌胜, 徐基于短程反硝化厌氧氨氧化的低碳源城市污水深度脱氮特性 ····································	建(1368)
基于短程反硝化戊氧氨氧化的低碳源城市污水深度脱氮特性····································	、臻(1377)
工艺工术 3000 全元 300 17 37 3 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1)晖(1384)
中试 MBBR 反应器启动 CANON 工艺及其短程硝化	航(1393)
抗生素对耐药型反硝化菌反硝化过程及微生物群落结构的影响 ······ 代莎,李彭,彭五庆,刘玉学,王拯,何义亮,沈根祥,胡双	人庆(1401)
多价近小·嗪气的好氧积位1万化余纪头短	、附(1409) 、臻(1418)
加生家对啊约型反硝化菌反硝化过程及微生物群洛结构的影响 "'代沙,辛彭,彭五庆,刘玉字,土拯,何又是,况稂拜,胡欢 多次进水-曝气的好氧颗粒污泥系统实验 "张杰,王玉颖,李冬,刘志诚,曹思 剩余污泥碱性发酵产物对硝化过程及性能的影响 邱圣杰,刘瑾瑾,李夕耀,彭永 硫酸盐对污泥高级厌氧消化过程中甲基汞迁移转化的影响	
一个一个一个一个一个一个一个一个一个一个一个一个一个一个一个一个一个一个一个	送(1425)
上业城市农田土壤重金属时至受异及米源解析····································	· 斌(1432) · 化(1440)
黄河三角洲盐碱土根际微环境的微生物多样性及理化性质分析	達 (1449)
不同土地利用方式对土壤细菌分子生态网络的影响 李冰,李玉双,魏建兵,宋雪英,史荣久,侯永侠,刘厶	瑶(1456)
红壤丘陵区土壤有机碳组分对土地利用方式的响应特征 · · · · · · · · · · 章晓芳,郑生猛,夏银行,胡亚军,苏以荣,陈香	碧 (1466)
有70加一一元70加工的加州的工N2U排放的影响	i新(14/4) i幼(1482)
等镁磷肥对石灰、海泡石组配修复镉污染稻田土壤的影响 ····································	高(1491)
桉树遗态磷灰石材料对铅污染土壤的钝化修复效应 方雅莉,朱宗强,赵宁宁,朱义年,李超,张立	浩(1498)
耕地土壤重金属健康风险空间分布特征	···· · 丰 / 1505 \
	. 肎(1505)
# 1 Not 1 4 # best 1 H 10 / 1 root 1 # 1 root 1 4 # best 1 day 1 (1 root 2 1 day 1	



城镇化进程中新疆塔城盆地浅层地下水化学演变特征 及成因

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摘要:随着经济的发展,新疆塔城盆地地下水开采量持续增大,然而研究区水文地质研究基础薄弱,盆地地下水化学演化趋势及其成因不明,这使得未来的地下水开发利用存在较大风险.本文在对盆地内地下水进行系统采样分析的基础上,基于5种水化学图对地下水化学组分进行异常识别,并对比历史水化学数据,对盆地地下水化学演变进行了深入分析.结果表明:研究区地下水阳离子以 Ca²+和 Na+为主,阴离子以 HCO₃ 和 SO₄ 为主,盆地广泛分布溶解性总固体小于1.0 g·L⁻¹的淡水.从山前淋溶迁移带到地下水径流缓慢的平原区,地下水化学类型由 HCO₃-Ca 和 HCO₃·SO₄-Ca·Mg 型过渡到 SO₄·HCO₃-Na·Ca 型.对比1979年水化学数据,城镇化进程中,由于地下水的过量开采,水位埋深下降,原来的部分排泄区变为径流区,水循环交替加快,致使研究区 HCO₃型和 SO₄·HCO₃型水分布面明显增加,以硫酸根和氯离子为主的高 TDS 水化学类型分布面积明显减少.然而在城镇周边人口密集区地下水中水氯离子和硝酸根离子明显升高,地下水 TDS 和总硬度呈上升趋势,地下水盐化和硬化明显.研究区地下水化学演变主要受潜水流经的含水层介质及地下水流场变化影响,另外排污沟渠污水下渗是影响地下水水质的另一个主要因素.

关键词:塔城盆地;城镇化;地下水化学;离子比;异常识别;控制因素

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Chemical Evolution of Groundwater in the Tacheng Basin of Xinjiang in the Process of Urbanization

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Abstract: With the development of the local economy, the volume of groundwater production has increased continuously in the past decades in the Tacheng Basin of the Xinjiang Uygur Automous Region. Previous studies have not provided a clear pattern of the chemical composition evolution of groundwater and its driving force in this basin, which makes the future development and utilization of groundwater riskier. This study carried out systematic sampling and analysis of groundwater chemistry in this basin, and the chemical evolution of groundwater in the basin was analyzed by comparison with historical hydrochemical data. The results show that Ca^{2+} and Na^{+} are the main cations in the groundwater, HCO_3^{-} , SO_4^{2-} are the main anions in the groundwater, and freshwater is widely distributed. The chemical types of groundwater changed from HCO_3 -Ca and $HCO_3 \cdot SO_4$ -Ca ·Mg in the source zone in front of the mountains to $SO_4 \cdot HCO_3$ -Na ·Ca type in the plain area. In comparison with the hydrochemical data of 1979, HCO_3 and $SO_4 \cdot HCO_3$ type groundwater increased significantly. SO_4 and SO_4^{-1} concentration and total hardness in the groundwater with high total dissolved solids decreased significantly. However, the CI^{-1} and SO_4^{-1} concentration and total hardness in the groundwater around the cities and towns increased. Aquifer material and the change of flowing field are the two controlling factors of groundwater chemical change, but the leakage of waste water from city drainage channels also affects the groundwater chemistry drastically.

Key words: Tacheng Basin; urbanization; hydrogeochemistry; ion ratio; abnormal recognition; controlling factors

西北干旱区内陆盆地,降雨稀少,蒸发强烈,地下水资源对区内生态系统、经济发展发挥着不可替代的作用^[1,2].在"一带一路"背景下,西北地区城镇化进程不断推进^[3],城市化的快速发展和市政建设的严重滞后,给当地地下水环境带来严峻压力,主要表现为地下水位下降、水质恶化^[4,5].水化学特征分析是研究水体化学组成与地下水演化过程的基础^[6~8],水体中主要离子成分常用来分析区域水化学控制因素及物质来源^[9,10].城镇化进程中人类活动对地下水环境的影响在很大程度上可以通过水化

学组分表现出来^[11~14],即出现"水化学异常".对地下水化学组分异常识别是确认人类活动对地下水影响程度的关键. Huang 等^[11]的研究表明,城镇化进程中,化粪池以及生活污水渗漏导致珠江三角洲地区地下水硝酸根离子明显升高,硝酸根在阴离子中

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的毫克当量百分比大于 25%, 硝酸型水在珠江三角 洲地区频繁显现. 李洁等^[12]的研究表明, 城镇化使 贵州岩溶区地下水溶滤作用减弱, 硫酸盐和氯化物 的输入增加. 朱丹尼等^[14]的研究表明, 城镇化进程 中生活污水及垃圾的大量输入导致地下水中三氮质 量浓度明显升高.

目前,国内外学者对城镇化进程中地下水化学 特征及水质演化研究较多,但是上述研究对城镇化 进程中人为输入对地下水影响的识别方法多局限于 数理统计、变异系数离散程度等方法[12,14]. 这些方 法主要根据某几项指标的变异程度评价人类活动对 地下水水质的影响,忽略了反映水化学特征的水化 学指标之间的内在联系,很可能造成评价结果失 准. 水化学方法识别人类活动对地下水水质的影 响程度是根据主要组分在地下水流向分布上出现 的异常,利用 Piper 三线图及各离子比例关系图等 进行异常识别,并最终得出人类活动影响程度[13]. 该方法避免了评价结果受单个指标极值的影响, 且从水化学场进行整体考虑,方法上更为科学合 理. 因此,本文以处于古丝绸之路经济中心带的新 疆塔城盆地为研究区,在对盆地内地下水进行系 统采样分析的基础上,综合多方面水文地质数据, 基于5种水化图对地下水主要组分进行异常识 别[15],探讨城镇化进程中盆地地下水化学演变特 征及控制因素,以期为当地地下水资源的可持续 开发与利用提供参考,并为我国西北干旱脆弱区 地下水化学演化研究提供借鉴.

1 材料与方法

1.1 研究区水文地质概况

塔城盆地位于新疆西北部,三面环山,盆地自东 北向西南倾斜,盆地边缘的山前地貌由互相毗邻的 洪积扇组成(图1和图2),中部为平原区. 塔城盆地 属内陆中温带干旱和半干旱气候区,多年平均气温 6.5℃. 多年平均降水量 300 mm, 集中在 7~9 月, 多 年平均蒸发量为1600 mm. 盆地内有大小河流共58 条,水量充沛,汇向盆地中心的额敏河. 河流的动态 类型为雪水型,春洪径流量约占年总流量的50%. 塔城盆地为新生代断陷盆地,其构造格局由5条区 域大断裂控制,且断裂走向由北东逐渐转为近东西 方向,平面上呈向西撒开、向东收敛的展布特 征[16]. 第四纪以来堆积了厚达几十到 250 m 的卵 石、砾石、砂和土层,组成从四周向盆地中部分布 的含水层(图1和图2).盆地内地下水从山前洪积 平原单一的砂砾石潜水向盆地中心过渡为多层结构 的潜水及承压水;含水层结构由简单到复杂,颗粒 由粗到细,地下水位埋藏深度由深到浅以至溢出地表. 额敏河横贯盆地中部,对盆地地下水起着总排泄作用. 大气降水、山区河床地下潜流的侧向补给、河流及渠道入渗补给是盆地平原区地下水的主要补给来源,人工开采以及额敏河的排泄和蒸发是地下水的主要排泄方式[17].

1.2 研究区城市化进程

塔城盆地农牧业发达,工业基础薄弱、企业布局分散,人类活动形式较为单一^[3].区内人口集中分布在塔城市、额敏县、托里县和裕民县等4个县市及其周边,在城镇周边分布有污水处理厂、垃圾填埋场、排污沟渠等潜在污染源.据塔城市统计年鉴^[19],研究区城镇化率呈逐年递增的趋势,城镇建设用地从2009年的8760.48 hm²发展到2015年的12009.34 hm²,年平均增长率为5.3%.城镇化进程中,城市生活污水和工业废水量逐年增加,但由于地处边疆,市政基础设施严重滞后,生活污水的泄漏和垃圾及人畜粪便的淋滤下渗导致城镇人口密集区地下水水质呈恶化趋势.农药化肥的使用导致区域地下水硝酸根离子呈上升趋势^[20].

1.3 样品采集与分析

2015年7~8月在研究区采集水样品90组(图 1). 其中,地下水样品80组(含泉水12组),地表水 样品 10 组. 地下水样品取自民井、农业灌溉井和 泉,采样井井深3~180 m,地下水位埋深0.5~120 m. 采样设备为离心泵或潜水泵. 采样前,对采样井 进行抽水清洗,等排出水量大于井孔储水量3倍,并 且现场指标稳定后再进行采样,以确保采集的样品 具有代表性. 当采样井中没有水泵时,采用中国地质 科学院水文地质环境地质研究所自主研发的"水斗 定深取样器"[21],通过人工提水方式在微扰动的情 况下,将井下目标深度的地下水直接密封后采出地 面. 所用采样瓶为 2.5 L 的高密度聚乙烯瓶. 取样 后,贴标签并用 Parafilm 封口膜密封避光保存, 7 d 内送达实验室测试. 样品由自然资源部地下水矿泉 水及环境监测中心测定,执行标准参照 GB/T 8538-95^[22]. 利用 DX-120 型离子色谱仪和 ICP-AES 分别 对阴、阳离子进行测定. 现场测试指标酸碱度、溶 解氧和氧化还原电位等采用多功能便携式测试仪 (Multi-340i/SET, 德国 WTW) 进行测试.

1.4 数据分析

运用舒卡列夫分类法分析水化学类型,运用 SPSS 和 AquaChem 等软件分析水化学数据,基于水 化学法和离子比分析法探讨地下水化学组分演变特 征及控制因素.

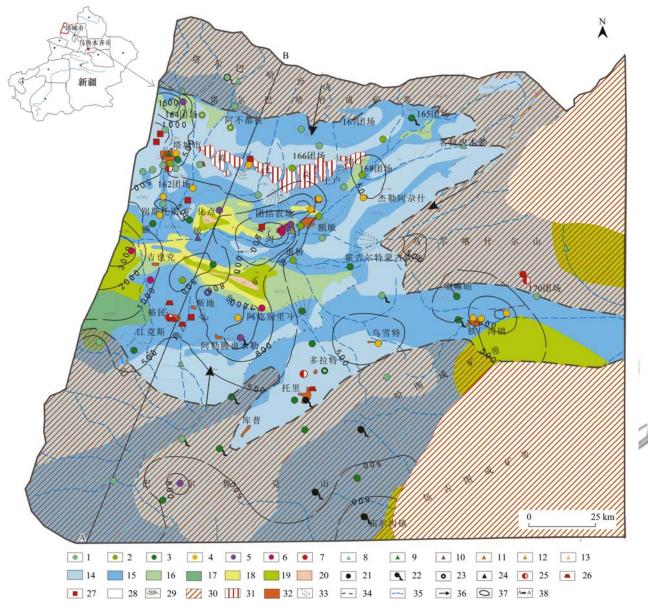


图 1 研究区水化学类型变化对比(1979年与2015年)

Fig. 1 Contrast of groundwater chemical type distribution in 1979 and 2015

2 结果与讨论

2.1 现状条件下地下水水化学组分特征

本次研究在山前洪积扇群采集地下水样品 25 组,在冲洪积平原区采集地下水样品 55 组,在出山口采集地表水 6 组,在排污沟渠下游采集地表污染水样 4 组(图 1). 塔城盆地不同地貌单元地下水总体上属于稳态水,略偏氧化水(图 3). 浅层地下水

pH 介于 7. 1~8. 5 之间, 为中性或弱碱性水, pH 值 从上游到下游呈逐渐降低趋势. 研究区地下水阳离子以 Ca^{2+} 和 Na^{+} 为主, 阴离子以 HCO_3^{-} 和 SO_4^{2-} 为主. 盆地潜水溶解性总固体 (TDS) 质量浓度介于 $161.6~3~881~mg\cdot L^{-1}$, 均值为 $669.0~mg\cdot L^{-1}$, 总体上为 TDS 小于 $1.0~g\cdot L^{-1}$ 的淡水. 从山前到平原, 沿着地下水流向地下水中的各主要离子组分、总硬及 TDS 呈逐渐升高的趋势 (图 4 和图 5) . 从各宏量组

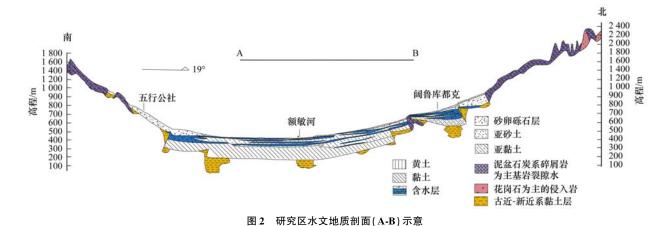


Fig. 2 Hydrogeological profile of the study area(A-B)

表 1 研究区水化学组分统计1)

			Table	1 Chemic	al composition	on of groun	dwater in tl	he study area	ı		
分	<u> </u>	检测项目	K +	Na +	Ca ^{2 +}	Mg^{2+}	Cl -	SO ₄ -	HCO ₃	$\mathrm{NO_3}^-$ -N	H_2SiO_3 TDS
		最小值/mg·L-1	0.35	6. 16	41. 72	6. 16	0. 35	10. 92	172. 8	0. 10	10. 92 198. 0
	山前	最大值/mg·L-1	7. 32	205.60	176. 90	36. 79	114. 30	511.80	471.2	15. 38	21.74 1 116.0
	(n=25)	均值/mg·L-1	2. 07	68. 76	77. 56	19.04	27. 68	156. 95	253. 5	4. 00	2. 74 507. 7
V++ 1.		变异系数/%	89. 5	81. 8	50. 2	49.6	42. 9	51.4	25. 3	115.3	16. 8 58. 0
潜水		最小值/mg·L-1	0.49	6. 14	41. 00	4.40	0. 35	14. 32	78. 5	0. 10	14. 35 161. 6
_	平原区	最大值/mg·L-1	3. 94	770. 80	281. 10	130. 10	404. 40	2 082. 00	519. 5	29. 22	24. 39 3 881. 0
	(n = 46)	均值/mg·L-1	1. 51	109. 41	109. 72	31.01	48. 61	305. 01	275.0	8. 01	19. 52 793. 9
6	9	变异系数/%	43. 2	131.0	54. 6	80.4	131.7	131.6	35. 8	99. 7	13. 3 90. 7
- }	1 6	最小值/mg·L-1	0. 64	9. 38	45. 32	9.06	1. 05	34. 49	158. 3	0.74	14. 20 228. 4
季 □→	平原区	最大值/mg·L-1	1.84	236. 70	137. 90	44. 67	184. 6	533. 3	277. 9	7.75	21. 53 1 289. 0
承压水	(n = 9)	均值/mg·L-1	1. 12	66. 20	68. 11	20.04	40. 9	145. 98	208. 5	3. 91	17. 87 478. 5
7	7	变异系数/%	38. 8	127. 8	42. 6	54. 4	149. 8	129. 8	18. 1	63.7	16. 8 79. 3

1)n表示样品采集数量(组)

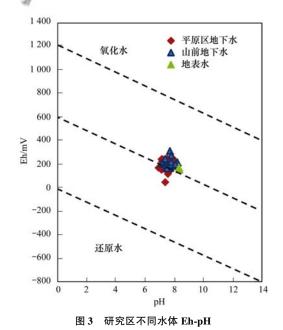


Fig. 3 Eh-pH plots of water samples in the study area

分的变异系数来看,除 pH 值外,其余指标的变异系 数值均大于10%,表现出中等或强变异性,表明各

离子在空间分布上存在较大离散性和波动性. 其 中,平原区地下水大部分离子组分变异系数较山 区大,平原区 Na⁺、NO₃⁻、SO₄²⁻、Cl⁻、Mg²⁺和 TDS 等指标表现出较强的变异性(表1),表明这些组 分在区域上分布不均匀,局部富集程度高.研究区 HCO3 与 Ca2+、Mg2+的变异系数相对较小,指示 其含量主要受地质背景条件控制,受外界人类活 动影响较小.

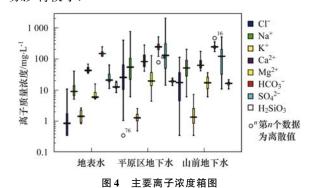


Fig. 4 Box plots showing variations in major ion concentrations in the water samples

2.2 地下水水化学时空演变特征

2.2.1 现状条件下地下水化学空间分布特征

根据舒卡列夫分类法(图 6),研究区地下水化学水平分带明显,潜水在盆地四周的基岩山区主要为 HCO₃·SO₄·Ca·Mg型水;山前丘陵区主要为 HCO₃·SO₄·Ca·Na型,局部为 SO₄·HCO₃·Na·Ca型水,到盆地中部冲积平原地下水化学类型从 HCO₃·SO₄ 型水过渡到 SO₄型水,溶解性总固体也逐渐增高(图 1).研究区承压水各离子组分浓度较潜水低,水化学类型在平原区周边主要为 HCO₃·SO₄·Ca·Mg型水,到平原区中部过渡为 SO₄·Cl·Na·Ca型水,TDS 值介于 228.4~1 289 mg·L⁻¹之间,总体上为 TDS 小于 1.0 g·L⁻¹的淡水.地表水在上游出山口主要为 HCO₃·Ca型水,到盆地中部额敏河一带过

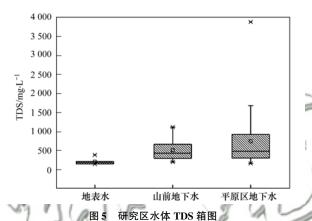


Fig. 5 Box plots showing variations of TDS in the water body of different areas in the study area

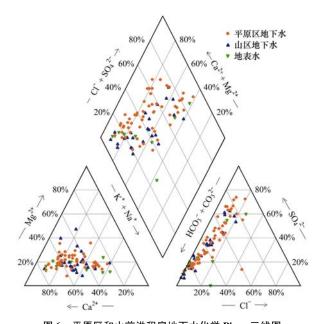


图 6 平原区和山前洪积扇地下水化学 Piper 三线图 Fig. 6 Piper plots of groundwater samples in the plain area and the piedmont area

渡到 HCO₃·SO₄- Ca·Na 型水. 盆地山前地下水与山区地表水具有类似的水化学演化特征(图 6),说明山区地表水体在出山口入渗补给地下水,二者水力联系密切.

塔城盆地南北两侧地下水化学类型具有明显的不对称性(图1),这是由于受下部隐伏构造的控制,盆地北部地下水系统洪积平原潜水由两个循环带组成,且黄土岗地南北含水层由洪积砂砾石层组成.这一独特的水文地质结构导致塔城盆地北部山前洪积平原地下水径流途径较短,水循环较快,以溶滤作用为主,蒸发浓缩作用相对较小,TDS总体较低,地下水以重碳酸、重碳酸硫酸型水为主.而在盆地南部地下水系统由山前到盆地中部地下水溶解性总固体逐渐升高.额敏河近岸,地表水与地下水交替循环,地下水水质较好,以重碳酸硫酸型水为主.然而在额敏河出国境处,基底隆起,隔水湖相黏土层拾升,地下水径流缓慢,河流两侧地下水水质明显变差,以硫酸或硫酸氯型水为主.

2.2.2 地下水化学随时间的演变特征

对比 1979 年水化学数据[18],研究区地下水在 20世纪六七十年代基本保持天然状态,南北部基岩 山区及盆地倾斜洪积平原区地下水大面积分布重碳 酸型、重碳酸硫酸型水,总体上为 TDS 值小于 1.0 g·L-1的淡水. 盆地中部冲积平原区以 Cl·SO4型和 SO₄型水为主的微咸水、咸水和卤水在额敏河两岸 的地势低洼处呈斑块晕状分布,在地下水的最终排 泄区晕中心 TDS 值大于 10 g·L-1. 近 40 年来, 随着 城镇化进程,人口数量逐年增加,城镇化面积逐渐加 大,地下水开采量持续增加,塔城盆地山前倾斜平原 地下水位埋深下降、泉流量减少、溢出带下移. 在 自然条件演化及人类活动输入双重影响下,盆地地 下动力条件发生了较大变化,原来的水岩平衡被打 破,水化学场发生了明显变化. 对比研究区 1979 年 及 2015 年水化学类型图(图1),塔尔巴哈台山前上 覆不同厚度黄土的波状、垅状冰水台地中赋存的潜 水水化学类型由之前的 HCO3·SO4-Ca·Na 型水转变 为 HCO₃-Ca·Mg型水,这是由于随着山前潜水埋深 下降,流经低渗透介质的潜水水量减少,大部分潜水 流经下伏于黄土的粗颗粒冰水堆积物,地下水循环 交替加快,地下水TDS 值下降,水化学类型向HCO。 型水演化. 研究区黄土岗地南侧部分区域浅层地 下水由 HCO3·SO4-Ca·Na 型水转变为 SO4·HCO3-Ca·Na型水,这是由于随着地下水位埋深的下降, 黄土岗以北泉流量减小,地下水以地表潜流的形 式流经黄土岗地,地下水径流速度变慢,在溶滤及 蒸发浓缩作用下,黄土岗地与山前弱倾斜平原交 界处潜水中硫酸根离子升高. 然而在额敏河两岸 地下水排泄区地势低洼处硫酸型及氯化物硫酸型 水分布面积明显减小,TDS 值明显降低,咸水和微 咸水分布面积大范围萎缩,卤水消失.这是由于, 城镇化进程中,额敏河两岸地下水开采量持续增 加,水位埋深下降,额敏河两岸沼泽及沼泽化湿地 面积不断缩小,原有部分排泄区转变为径流区,水 化学交替加强,蒸发浓缩作用较之前明显减弱,原 有的咸水向微咸水演变. 另外,在塔城市及额敏县 周边人类活动密集区,地下水 TDS 和总硬度总体 呈上升趋势,受人类活动影响较大的氯离子和硝 酸根离子质量浓度增加,地下水由 HCO3-Ca·Mg 和 HCO、·SO₄-Ca·Mg 型演化为SO₄·HCO、-Ca·Mg (Ca·Na)和SO₄-Na·Ca型水.南部巴尔鲁克山前人 类活动相对密集区庙尔沟镇地下水由重碳酸硫酸 型演化为 SO4·HCO3-Ca·Na 型水.

总之,从水化学数据的历史变化趋势来看,在过去近40年的城镇化进程中,研究区地下水化学类型发生了明显变化. HCO₃型水和SO₄·HCO₃型水分布面明显增加,以硫酸根和氯离子为主的高溶解性总固体水化学类型分布面积均有不同程度下降. 在城镇周边人类活动密集区,重碳酸型水范围逐渐缩小,而人类活动来源的氯化物硫酸根浓度升高,地下水盐化和硬化明显. 另外,在垃圾填埋场、排污沟渠等大型污染源附近,水体中的 Na⁺、Cl⁻和 NO₃⁻ 明显升高,水化学类型由 HCO₃-Ca 型向 HCO₃·Cl -Na·Ca和 SO₄·Cl -Na·Ca

2.3 地下水化学演变成因分析

2.3.1 基于"水化学法"的异常数据识别

应用"水化学法"识别人为活动输入对地下水 水质的影响程度[23],该方法充分考虑了各水化学指 标之间的内在联系,且能反映溶滤、蒸发、混合、离 子交换等一系列的水化学演化过程[24]. 马氏距离是 1936年由印度统计学家马哈拉诺比斯(P.C. Mahalanobis)提出,表示数据协方差距离,能有效计 算两个未知样本集的相似度[25]. 本文利用反映水化 学演化过程的5种水化学图与马氏距离相结合[15], 识别水化学异常,量化人类活动对地下水水质的影 响程度,预测地下水受污染的潜在风险. 5 种水化 学图主要包含水化学类型 Piper 三线图、离子交换 作用 $(Ca^{2+} + Mg^{2+} - Na^{+} - K^{+}) - (HCO_{3}^{-} - SO_{4}^{2-} -$ Cl⁻)以及3种离子比关系[(Ca²⁺ + Mg²⁺)-HCO₃、 (Na⁺ + K⁺)-Cl⁻、(Ca²⁺ + Mg²⁺)-SO₄²⁻]. 利用水化 学图转换的5幅二维坐标图,计算图中各点所对应 数据集的马氏距离进行异常数据识别. 基于 5 种水 化学图异常数据识别情况见表 2 和图 7.

Piper 三线图二维坐标转换公式:

$$\begin{split} X &= 1.\ 125 - \left(1 - \frac{v m_{\text{Na+K}}}{\sum v m_{\text{e}}}\right) \times 0.\ 5 + \\ &\left(1 - \frac{v m_{\text{HCO}_3 + \text{CO}_3}}{\sum v m_{\text{a}}}\right) \times 0.\ 5 \\ Y &= a + \left(1 - \frac{v m_{\text{Na+K}}}{\sum v m_{\text{e}}}\right) \times b + \left(1 - \frac{v m_{\text{HCO}_3 + \text{CO}_3}}{\sum v m_{\text{e}}}\right) \times b \end{split}$$

式中, m_e 表示各阳离子(Na⁺、K⁺、Ca²⁺、Mg²⁺)的物质的量浓度; m_a 表示各阴离子(HCO₃⁻、CO₃²⁻、SO₄²⁻、Cl⁻)的物质的量浓度; m_{Na+K} 表示钾离子和钠离子物质的量浓度之和; $m_{HCO_3+CO_3}$ 表示重碳酸根和碳酸根物质的量浓度之和;v 表示对应离子的电荷数; $a=0.25\times(3^{0.5})/2$; $b=(3^{0.5})/2$.

马氏距离公式:
$$\left(X - \overline{X}\right)'S^{-1}\left(X - \overline{X}\right) = Da^{2}$$

式中,X 表示数组毫克当量浓度;X表示数组毫克当量浓度平均值; S^{-1} 表示协方差逆矩阵; Da 表示马氏距离.

得到 5 种二维坐标水化学图对应的马氏距离后,分别利用马氏距离的均值加减 3 倍标准差作为临界距离(Di²),比较 Di²与 Da²的大小,将 Da²值超出临界距离(Di²)的样品全部剔除,即所有剩余样品对应的马氏距离都处于均值加减 3 倍标准差内.然后重新计算所剩样品的均值、协方差矩阵、马氏距离,再次比较 Di²与 Da²的大小,反复循环,直到剩余样品的马氏距离均小于临界距离 Di²,即无离群样品.

表 2 各水化学图异常数据识别1)

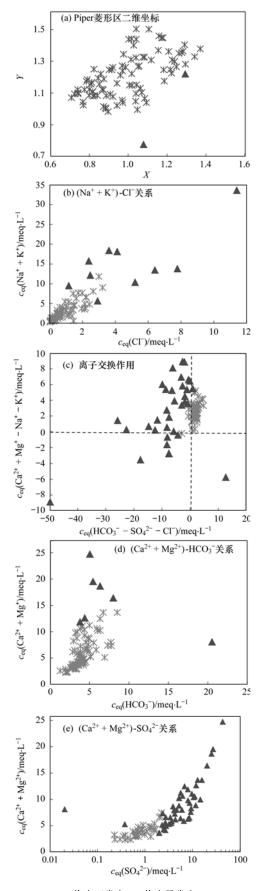
Table 2 Identification of abnormal water chemical data

from the nythoc	memicai maps
水化学图	异常数据识别(组)
Piper 三线图	2
离子交换图	33
(Na $^+$ + K $^+$) - Cl $^-$	10
$(Ca^{2+} + Mg^{2+}) - HCO_3^{-}$	7
$(Ca^{2+} + Mg^{2+}) - SO_4^{2-}$	50
合计	50

1)合计为各方法识别的总数含有重复项

2.3.2 异常数据主要影响指标及其演变成因分析

为探明地下水水质与人类活动影响的关系,对比 2017 年发布的《地下水质量标准》(GB/T 14848-2017) [26] 中各指标的限值及推荐的地下水水质单因子评价法,选取 Na⁺、Ca²⁺、Mg²⁺、Cl⁻、SO₄²⁻、HCO₃⁻、NO₃⁻、NO₂⁻、NH₄⁺、Fe、Mn、Zn、Cu、COD、TDS、总硬度、F⁻、I⁻、Cd、Hg、As、Se 和 Pb 等 23



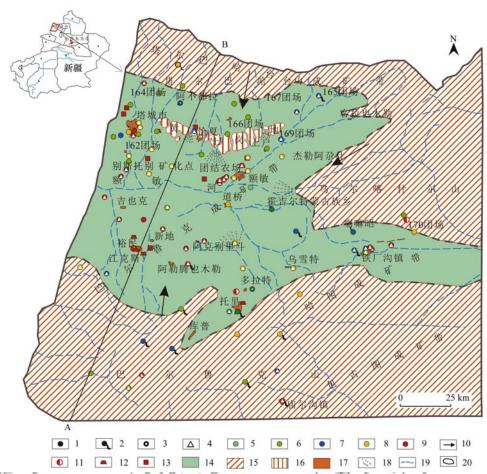
※代表正常点,▲代表异常点图 7 5 种水化学图法异常数据识别

Fig. 7 Abnormal data identification figures from five water chemical diagram methods

项无机指标对研究区地下水进行质量评价. 由图 8 可见,地下水水质超 II 类水与基于"水化学法"识别出的受人类活动影响的异常点分布几乎一致. 主要分布在盆地中部冲洪积平原人类活动密集区及其大型污染源周边. 然而在南北两侧山体受原生沉积环境影响,也分布有地下水水质超 II 类水,其超标组分主要为 Fe、Mn 和 F⁻. 位于盆地北部及南部山前倾斜洪积平原区部分样品点为异常点,但地下水水质未超出 II 类水标准. 由此可见,山前地带部分区域地下水已经受人类活动影响,虽然其地下水水质相对较好,但是可以推测该区域地下水化学正在发生变化,其受污染的潜在风险较大.

研究区识别的异常点地下水水质超Ⅲ类水主要 影响指标为 SO₄²⁻、总硬度、TDS、Fe、F⁻、Na⁺、 Mg²⁺和 NO₃ 等 15 项(图 9). 其中,SO₄ 、总硬度和 TDS的贡献率较大,贡献率均超过了50%.其次是 Fe、F⁻、Na⁺、Mg²⁺和NO₃ 等,贡献率介于27%~ 42%. 另外, Mn、NH₄ 、Cl 、COD、I 、NO₂ 和 As 等指标贡献率3%~7%.由此可见,影响塔城盆地地 下水水质的指标呈多样性,除人类活动输入造成的 "三氮"、CI⁻和 COD 以外,受原生沉积环境所形成 的高硫酸盐、高硬度、高TDS、高铁锰、高氟等天 然劣质水是研究区地下水水质较差的主要原因. 研 究区南北两侧山体高铁劣质水主要受盆地地层中铁 锰矿物风化溶解所致,盆地中部人类活动强烈区额 敏河两侧地下水中铁锰超标点主要为天然劣质水受 人类活动的影响其铁锰离子浓度值升高而发生进一 步劣变恶化[27].

城镇化进程中,水资源需求量增大,长期以来地 下水的过量开采及人类活动产生的污染物直接作用 于地下水,对研究区地下水水质造成威胁.研究区一 方面由于过量开采地下水,改变了地下水循环途径, 天然水盐平衡遭到破坏,致使研究区水化学特征发 生变化. 如前文所述,流经塔尔巴哈台山前地下水, 由于潜水位埋深下降,流经的含水层介质颗粒变粗, 水循环加快,SO²⁻ 离子浓度降低,HCO; 离子浓度 升高,地下水溶解性总固体降低,呈淡化趋势.而盆 地北部黄土岗地以南弱倾斜平原区,由于水位下降, 流经的含水层介质为含石膏的二叠系、石炭系砂岩 组成的冲洪积物[28],地下水在径流过程中可能发生 了石膏和泻盐的淋溶作用,地下水中 SO₄ 高子浓 度升高. 而在地下水的总排泄区额敏河两侧,由于过 量开采,地下水位持续下降,原来的部分沼泽湿地逐 渐消失,蒸发浓缩作用明显减弱,地下水溶解性总固 体降低明显,原有的咸水向微咸水演变.另一方面, 塔城盆地城镇人口密集区污染物的直接输入导致地



取样点类型:1. 潜水; 2. 泉; 3. 承压水; 4. 异常点; 水质级别: 5. I 类水; 6. II 类水; 7. II 类水; 8. IV 类水; 9. V 类水; 边界及其他:10. 地下水流向; 11. 污水处理厂; 12. 垃圾填埋场; 13. 农业污染源; 14. 松散岩类孔隙水 15. 山区基岩裂隙水; 16. 黄土岗地; 17. 城镇区; 18. 洪 积扇; 19. 地表水系; 20. 研究区边界

图 8 地下水质量评价及取样点(含异常点)分布示意

Fig. 8 Groundwater quality assessment and sampling points (including abnormal points) distribution map

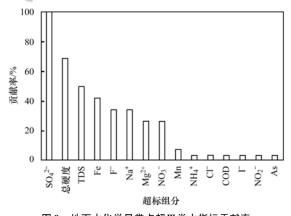


图 9 地下水化学异常点超Ⅲ类水指标贡献率

Fig. 9 Contribution rate of over-standard components of groundwater chemical abnormal data

下水水质恶化,主要表现在 Cl⁻、NO₃⁻、NO₂⁻、NH₄⁺和 COD 等指标发生变化,出现超Ⅲ类地下水.受城市生活污水以及垃圾渗滤液等点状污染物的输入影响,城镇周边人口密集区地下水 NO₃⁻ 离子和耗氧量升高明显. 塔城盆地工矿业薄弱,这可能是地下水不存在重金属超标的主要原因. 总之,从超标组分的贡

献程度来看,硫酸根、溶解性总固体、铁和氟离子等指标为塔城盆地地下水水质的主要影响指标,而硝酸根、化学需氧量等为研究区地下水水质的次要影响指标. 塔城盆地劣质地下水的劣质组分主要来源于天然,多为常规指标,大部区域地下水水质受原生沉积环境所控,人类活动诱导而形成的劣变水,也有部分区域叠加人类活动污染输入进一步劣变恶化.

3 结论

(1)塔城盆地地下水化学类型以 HCO₃-Ca(Ca·Mg)、HCO₃·SO₄-Ca·Mg、SO₄·HCO₃-Na·Ca 型为主,沿着地下水流向,地下水由山前的 HCO₃-Ca 和HCO₃·SO₄-Ca·Mg 型水过渡到平原区的 SO₄·HCO₃-Na·Ca 型水,地下水 TDS 也逐渐升高. 对比历史水化学数据,城镇化进程中,研究区地下水化学类型呈复杂的演化特征. 区域上 HCO₃ 型水和 SO₄·HCO₃型水分布面明显增加,盆地中部额敏河两侧以硫酸根和氯离子为主的高 TDS 水化学类型分布面积明显

减少. 而在城镇周边人口密集区,地下水由 HCO₃-Ca·Mg 和 HCO₃·SO₄-Ca·Mg 型演化为 SO₄·HCO₃-Ca·Mg(Ca·Na)和 SO₄-Na·Ca 型水,地下水呈盐化、硬化趋势.

(2)研究区地下水化学演化受天然和人为因素双重作用影响,其中地下水超采、岩石风化溶解和蒸发浓缩作用是地下水演化的主导因素. 塔城盆地超Ⅲ类地下水劣质组分多为常规指标,主要来源于天然,大部分区域地下水水质受原生沉积环境所控,人类活动诱导而形成劣变水. 研究区地下水化学演变主要受潜水流经的含水层介质及地下水流场变化影响,另外城镇化进程中,排污沟渠污水以及垃圾渗滤液下渗是影响地下水水质的另一重要因素.

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HUANJING KEXUE

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CONTENTS

Evolution of the Distribution of PM _{2, 5} Concentration in the Yangtze River Economic Belt and Its Influencing Factors	
Chemical Characteristics and Source Apportionment of Water-Soluble Ions in Atmosphere Aerosols over the East China Sea Island D	Turing Winter and Summer
	FANG Yan, CAO Fang, FAN Mei-yi, et al. (1025)
Analysis of Chemical Components and Sources of PM _{2.5} During Autumn and Winter in Yangquan City ·····	WANG Cheng, YAN Yu-long, XIE Kai, et al. (1036)
Chemical Compositions and Sources of n-Alkanes and Saccharides in PM _{2.5} from Taian City During the Summer	······ YI Ya-nan, HOU Zhan-fang, YANG Qian-cai, et al. (1045)
Physical and Chemical Characteristics of Atmospheric Particles in Autumn in Mt. Huangshan	
Characteristics of Aerosol Optical Depth in the Urban Area of Beibei and Its Correlation with Particle Concentration	
Emission Characteristics of Volatile Organic Compounds from Typical Industries in Zibo	
Analysis of Pollution Characteristics and Sources of Atmospheric VOCs in Ezhou City	
Producing Coefficients and Emission Coefficients of Volatile Organic Compounds from the Automobile Manufacturing Industry in Zhe	giang Province
$Pollution\ Characteristics\ of\ Volatile\ Organic\ Compounds\ Emission\ from\ the\ Metal\ Packaging\ Industry\ Based\ on\ Analysis\ of\ Process$	WANG Di, NIE Lei, ZHAO Wen-juan, et al. (1099)
Profile Characteristics of VOCs from Wood and Economic Crop Burning	
Accounting Methods of VOCs Emission Associated with Production Processes in a Fine Chemical Industrial Park	
HONO Observation and Assessment of the Effects of Atmospheric Oxidation Capacity in Changzhou During the Springtime of 2017	
Analysis of Activity and Its Emissions Trend for Construction Equipment in China	
Air Pollutant Emission Inventory from LTO Cycles of Aircraft in the Beijing-Tianjin-Hebei Airport Group, China	
Particle Size Distribution of PM Emission from In-use Gasoline and Diesel Vehicles	
Impact of Parameterization on the Estimation of Ammonia Emissions; A Case Study over the Yangtze River Delta	
Characteristics and Source Apportionment of Dustfall Pollution in the Coal Mine Area and Surrounding Areas of Wuhai City in Spring	g 5
	···· WU Hong-xuan, SHI Chang-qing, ZHANG Yan, et al. (1167)
Variations of Stable Oxygen and Deuterium Isotopes in River and Lake Waters During Flooding Season Along the Middle and Lower	Reaches of the Yangtze River Regions ·····
	LI Jing, WU Hua-wu, ZHOU Yong-qiang, et al. (1176)
Water Sources and Factors Controlling Hydro-chemical Compositions in the Yiluo River Basin	
Chemical Evolution of Groundwater in the Tacheng Basin of Xinjiang in the Process of Urbanization	
Remote Sensing Monitoring on Spatial Differentiation of Suspended Sediment Concentration in a River-Lake System Based on Sentin	
Yangtze River Section in Anhui Province	······ WANG Hang-hang, WANG Jie, CUI Yu-huan (1207)
CDOM Optical Characteristics and Related Environmental Factors of High-turbidity Waters on the Loess Plateau	LIANG Xiao-wen, SHAO Tian-tian, WANG Tao (1217)
Effects of Artificial Destratification and Induced-natural Mixing on Water Quality Improvement in a Drinking Water Reservoir	
Effect of Filter Speed and Water Quality on Ammonia Removal in Groundwater Containing Iron, Manganese, and Ammonia at Low	Temperature
Long-term Variation Characteristics of Zooplankton Community Structure in Meiliang Bay, Lake Taihu	······ YANG Jia, ZHOU Jian, QIN Bo-qiang, et al. (1246)
Community of Benthic Diatoms and Their Relationship with Aquatic Environmental Factors in the Tangwang River, China	······ XUE Hao, WANG Ye-yao, MENG Fan-sheng, et al. (1256)
Succession Characteristics and Water Quality Responsiveness Evaluation of FG and MBFG in Yanlong Lake Water Source Ecologica	l Purification System ·····
Succession Characteristics and Water Quality Responsiveness Evaluation of FG and MBFG in Yanlong Lake Water Source Ecologica	l Purification System
Characteristics of Nitrogen and Phosphorus Output and Loss Flux in the Shipanqiu Watershed, Three Gorges Reservoir Area	
Characteristics of Nitrogen and Phosphorus Output and Loss Flux in the Shipanqiu Watershed, Three Gorges Reservoir Area Effect of Optimized Fertilization and Biochar Application on Phosphorus Loss in Purple Soil Sloping Farmland Use of Iron-modified Calcite as an Active Capping Material to Control Phosphorus Release from Sediments in Surface Water Bodies	
Characteristics of Nitrogen and Phosphorus Output and Loss Flux in the Shipanqiu Watershed, Three Gorges Reservoir Area Effect of Optimized Fertilization and Biochar Application on Phosphorus Loss in Purple Soil Sloping Farmland Use of Iron-modified Calcite as an Active Capping Material to Control Phosphorus Release from Sediments in Surface Water Bodies Preparation of Tea Waste Biochar and Its Application in Tetracycline Removal from Aqueous Solution	
Characteristics of Nitrogen and Phosphorus Output and Loss Flux in the Shipanqiu Watershed, Three Gorges Reservoir Area Effect of Optimized Fertilization and Biochar Application on Phosphorus Loss in Purple Soil Sloping Farmland Use of Iron-modified Calcite as an Active Capping Material to Control Phosphorus Release from Sediments in Surface Water Bodies Preparation of Tea Waste Biochar and Its Application in Tetracycline Removal from Aqueous Solution	
Characteristics of Nitrogen and Phosphorus Output and Loss Flux in the Shipanqiu Watershed, Three Gorges Reservoir Area Effect of Optimized Fertilization and Biochar Application on Phosphorus Loss in Purple Soil Sloping Farmland Use of Iron-modified Calcite as an Active Capping Material to Control Phosphorus Release from Sediments in Surface Water Bodies Preparation of Tea Waste Biochar and Its Application in Tetracycline Removal from Aqueous Solution	
Characteristics of Nitrogen and Phosphorus Output and Loss Flux in the Shipanqiu Watershed, Three Gorges Reservoir Area Effect of Optimized Fertilization and Biochar Application on Phosphorus Loss in Purple Soil Sloping Farmland Use of Iron-modified Calcite as an Active Capping Material to Control Phosphorus Release from Sediments in Surface Water Bodies Preparation of Tea Waste Biochar and Its Application in Tetracycline Removal from Aqueous Solution	
Characteristics of Nitrogen and Phosphorus Output and Loss Flux in the Shipanqiu Watershed, Three Gorges Reservoir Area Effect of Optimized Fertilization and Biochar Application on Phosphorus Loss in Purple Soil Sloping Farmland Use of Iron-modified Calcite as an Active Capping Material to Control Phosphorus Release from Sediments in Surface Water Bodies Preparation of Tea Waste Biochar and Its Application in Tetracycline Removal from Aqueous Solution Adsorption Characteristics of Sulfamethazine on Three Typical Porous High-temperature Modified Solid Waste Materials Sorption of Polybrominated Diphenyl Ethers by Virgin and Aged Microplastics Spatial Distribution and Risk Assessment of Heavy Metals in Sediments of the Ruxi Tributary of the Three Gorges Reservoir Distribution and Risk Assessment of OCPs in Surface Water, Sediments, and Fish from Lake Gucheng and Inflow and Outflow Rive	
Characteristics of Nitrogen and Phosphorus Output and Loss Flux in the Shipanqiu Watershed, Three Gorges Reservoir Area Effect of Optimized Fertilization and Biochar Application on Phosphorus Loss in Purple Soil Sloping Farmland Use of Iron-modified Calcite as an Active Capping Material to Control Phosphorus Release from Sediments in Surface Water Bodies Preparation of Tea Waste Biochar and Its Application in Tetracycline Removal from Aqueous Solution	
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Characteristics of Nitrogen and Phosphorus Output and Loss Flux in the Shipanqiu Watershed, Three Gorges Reservoir Area Effect of Optimized Fertilization and Biochar Application on Phosphorus Loss in Purple Soil Sloping Farmland Use of Iron-modified Calcite as an Active Capping Material to Control Phosphorus Release from Sediments in Surface Water Bodies Preparation of Tea Waste Biochar and Its Application in Tetracycline Removal from Aqueous Solution Adsorption Characteristics of Sulfamethazine on Three Typical Porous High-temperature Modified Solid Waste Materials Sorption of Polybrominated Diphenyl Ethers by Virgin and Aged Microplastics Spatial Distribution and Risk Assessment of Heavy Metals in Sediments of the Ruxi Tributary of the Three Gorges Reservoir Distribution and Risk Assessment of OCPs in Surface Water, Sediments, and Fish from Lake Gucheng and Inflow and Outflow Rive Occurrence and Ecological Risk Assessment of Typical Persistent Organic Pollutants in Hengshui Lake Pollution Characteristics and Risk Assessment of Typical Organophosphate Esters in Beijing Municipal Wastewater Treatment Plant and the surface of the properties of the properties of the properties of the properties of the Pollution Characteristics and Risk Assessment of Typical Organophosphate Esters in Beijing Municipal Wastewater Treatment Plant and Pollution Characteristics and Properties of the Pollution Characteristics and Properties of Typical Organophosphate Esters in Beijing Municipal Wastewater Treatment Plant and Properties of Pollution Characteristics and Properties of Pollution Charac	WANG Lian, LI Xuan, MA Wei-xing, et al. (1265) CHEN Shi-qi, LONG Yi, YAN Dong-chun, et al. (1276) LUO Dong-hai, WANG Zi-fang, LONG Yi, et al. (1286) BAI Xiao-yun, LIN Jian-wei, ZHAN Yan-hui, et al. (1296) FAN Shi-suo, LIU Wen-pu, WANG Jing-tao, et al. (1308) WANG Jing, ZHU Xiao-li, HAN Zi-yu, et al. (1319) XU Peng-cheng, GUO Jian, MA Dong, et al. (1329) FANG Zhi-qing, WANG Yong-min, WANG Xun, et al. (1338) KAN Ke-cong, GU Xiao-hong, LI Hong-min, et al. (1346) ZHANG Jia-wen, WEI Jian, LÜ Yi-fan, et al. (1357) and the Receiving Water ZHANG Zhen-fei, LÜ Jia-pei, PEI Ying-ying, et al. (1368)
Characteristics of Nitrogen and Phosphorus Output and Loss Flux in the Shipanqiu Watershed, Three Gorges Reservoir Area Effect of Optimized Fertilization and Biochar Application on Phosphorus Loss in Purple Soil Sloping Farmland Use of Iron-modified Calcite as an Active Capping Material to Control Phosphorus Release from Sediments in Surface Water Bodies Preparation of Tea Waste Biochar and Its Application in Tetracycline Removal from Aqueous Solution Adsorption Characteristics of Sulfamethazine on Three Typical Porous High-temperature Modified Solid Waste Materials Sorption of Polybrominated Diphenyl Ethers by Virgin and Aged Microplastics Spatial Distribution and Risk Assessment of Heavy Metals in Sediments of the Ruxi Tributary of the Three Gorges Reservoir Distribution and Risk Assessment of OCPs in Surface Water, Sediments, and Fish from Lake Gucheng and Inflow and Outflow Rive Occurrence and Ecological Risk Assessment of Typical Persistent Organic Pollutants in Hengshui Lake Pollution Characteristics and Risk Assessment of Typical Organophosphate Esters in Beijing Municipal Wastewater Treatment Plant Advanced Nitrogen Removal Characteristics of Low Carbon Source Municipal Wastewater Treatment via Partial-denitrification Couple Stable Nitrite Accumulation and Phosphorus Removal from High-nitrate and Municipal Wastewaters in a Combined Process of Partia	WANG Lian, LI Xuan, MA Wei-xing, et al. (1265) CHEN Shi-qi, LONG Yi, YAN Dong-chun, et al. (1276) LUO Dong-hai, WANG Zi-fang, LONG Yi, et al. (1286) FAN Shi-suo, LIU Wen-pu, WANG Jing-tao, et al. (1308) WANG Jing, ZHU Xiao-li, HAN Zi-yu, et al. (1319) YU Peng-cheng, GUO Jian, MA Dong, et al. (1329) FANG Zhi-qing, WANG Yong-min, WANG Xun, et al. (1338) KAN Ke-cong, GU Xiao-hong, LI Hong-min, et al. (1346) ZHANG Jia-wen, WEI Jian, LÜ Yi-fan, et al. (1357) and the Receiving Water ZHANG Zhen-fei, LÜ Jia-pei, PEI Ying-ying, et al. (1368) ad with ANAMMOX MA Bin, XU Xin-xin, GAO Mao-hong, et al. (1377) I Denitrification and Denitrifying Phosphorus Removal
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Characteristics of Nitrogen and Phosphorus Output and Loss Flux in the Shipanqiu Watershed, Three Gorges Reservoir Area Effect of Optimized Fertilization and Biochar Application on Phosphorus Loss in Purple Soil Sloping Farmland Use of Iron-modified Calcite as an Active Capping Material to Control Phosphorus Release from Sediments in Surface Water Bodies Preparation of Tea Waste Biochar and Its Application in Tetracycline Removal from Aqueous Solution Adsorption Characteristics of Sulfamethazine on Three Typical Porous High-temperature Modified Solid Waste Materials Sorption of Polybrominated Diphenyl Ethers by Virgin and Aged Microplastics Spatial Distribution and Risk Assessment of Heavy Metals in Sediments of the Ruxi Tributary of the Three Gorges Reservoir Distribution and Risk Assessment of OCPs in Surface Water, Sediments, and Fish from Lake Gucheng and Inflow and Outflow Rive Occurrence and Ecological Risk Assessment of Typical Persistent Organic Pollutants in Hengshui Lake Pollution Characteristics and Risk Assessment of Typical Organophosphate Esters in Beijing Municipal Wastewater Treatment Plant Advanced Nitrogen Removal Characteristics of Low Carbon Source Municipal Wastewater Treatment via Partial-denitrification Couple Start-up of CANON Process and Short-cut Nitrification in a Pilot-scale MBBR Reactor	WANG Lian, LI Xuan, MA Wei-xing, et al. (1265) CHEN Shi-qi, LONG Yi, YAN Dong-chun, et al. (1276) LUO Dong-hai, WANG Zi-fang, LONG Yi, et al. (1286) FAN Shi-suo, LIU Wen-pu, WANG Jing-tao, et al. (1308) WANG Jing, ZHU Xiao-li, HAN Zi-yu, et al. (1319) XU Peng-cheng, GUO Jian, MA Dong, et al. (1329) FANG Zhi-qing, WANG Yong-min, WANG Xun, et al. (1338) KAN Ke-cong, GU Xiao-hong, LI Hong-min, et al. (1346) ZHANG Jia-wen, WEI Jian, LÜ Yi-fan, et al. (1357) and the Receiving Water ZHANG Zhen-fei, LÜ Jia-pei, PEI Ying-ying, et al. (1377) I Denitrification and Denitrifying Phosphorus Removal WANG Qiu-ying, YU De-shuang, ZHAO Ji, et al. (1384) FU Kun-ming, YANG Zong-yue, LIAO Min-hui, et al. (1393) mmunity Structure DAI Sha, LI Peng, PENG Wu-qing, et al. (1401) ZHANG Jie, WANG Yu-ying, LI Dong, et al. (1418)
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