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### 千岛湖湖泊区水体季节性分层特征研究

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摘要: 千岛湖是典型的亚热带人工深水水库. 为了解水库的物理特征,于 2011 年对水库温度、溶解氧(DO)、pH 值、浊度、有效光合作用辐照强度(PAR)和叶绿素 a 的季节动态和垂直分层结构进行分析,探讨水库水体季节性分层特征. 结果表明,千岛湖水库表层水温在 10.4~32.7℃,在 4~12 月间出现水体分层现象,在夏季形成明显分层,是典型的亚热带单混合水体.水温分层结构对 DO、pH 值、浊度、PAR 和叶绿素 a 的垂直剖面变化有一定的影响作用. 水体溶解氧出现显著的分层现象,且温跃层内 DO 浓度较低,冬季底层溶解氧出现暂时周期性缺氧;浊度在温跃层内达到最大值,其表层大小主要受降水的影响;叶绿素 a 含量表层含量较高,且温跃层以下水深维持较低水平,这说明水体分层对水库水质的变化起重要作用. 此外,水体光合作用最强处在 5~10 m 左右,此深度适合藻类生长,这为分析千岛湖水体蓝藻暴发取样和分析提供了依据.

关键词:千岛湖水库;水体分层;亚热带;季节性;物理特征

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# Study on Seasonal Characteristics of Thermal Stratification in Lacustrine Zone of Lake Qiandao

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Abstract: Lake Qiandao is a typical subtropical man-made reservoir in China. The investigation on the seasonal and vertical dynamics of water temperature, dissolved oxygen (DO), pH value, turbidity, photosynthetic available radiation (PAR) and chlorophyll a was conducted in 2011 in order to find out the physical characteristics of Lake Qiandao. The average surface water temperature ranged from 10.4 to 32.7°C. A monomictic thermal stratification was observed in Lake Qiandao, initiating in April and lasting until December. The results showed that thermal stratification had influences on vertical distribution of DO, pH value, turbidity, PAR and chlorophyll a. Very strong stratification of DO was found, inducing lower oxygen concentration in the thermocline layer and temporal hypoxia in the bottom water. The maximum turbidity was found in the thermocline layer and the precipitation affected the surface turbidity value. Moreover, the chlorophyll a concentration was higher in the surface water and lower in the bottom water as found in this study, implying that water quality was affected by stratification. Besides, the maximum photosynthesis rate and algal growth rate were found at the depth 5-10 m below the water surface. Therefore, the results can provide theoretical support for the sampling and analysis of algal blooms in Lake Qiandao.

Key words: Lake Qiandao; water stratification; subtropics; seasonality; physical characteristics

水库(特别是深水水库)水温由于受季节变化而引起的分层和对流现象是水库的一个重要特征,在深水型湖库中水温分层的温差大、持续时间长,呈现季节性分层的现象,影响着水体的物理特性、化学反应过程和生物活动<sup>[1]</sup>.一般来说,深度超过7 m的水库就可能形成温跃层<sup>[2]</sup>.在温跃层内水深每下降1 m,水温下降1℃或1℃以上<sup>[3]</sup>.分层湖库中水温分布、温跃层的形成和湖泊或水库内水体的上下交换,对水体的溶解氧垂向分布、营养物质分布、水生生物,特别是浮游植物的垂向分布产生重大影响<sup>[4-9]</sup>.稳定的热分层可阻碍上下水层水体的混合和物质交换,导致了水化学的分层,形成化跃

层<sup>[10]</sup>. 突发性水质恶化事件和温跃层及化跃层的改变有关<sup>[11]</sup>. 因为气温突降使水体分层失稳,上下层水体发生等温对流,下层富集的还原性物质及营养盐向上迁移,从而发生缺氧等水质恶化事件,并对水体水质和水环境变化起着重要作用<sup>[12~14]</sup>. 王雨春等<sup>[4]</sup>通过研究也发现水体分层现象的发生和变化直接控制水库的水环境性质,百花湖水库近年来

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经常发生的季节性水质恶化事件,是由于秋季突然降温导致的水体分层结构失稳,上下水团在垂直交替过程中耗氧和复氧机制不平衡引起的. 因此,对于有一定水深的水库,水体的季节性分层对深水水库理化指标的垂向分布起着控制或者影响作用,直接影响着水环境质量的变化. 气温突降是水质恶化事件发生的诱因,而水体的分层是发生的条件[12]. 因此,对深水水库的水环境的研究必须首先明确温度分层的形成和发展过程,因为温度分层的发生和破坏会引起水化学参数的变化,进而为水华的发生或水质的污染创造条件.

千岛湖是典型的亚热带人工深水水库,是钱塘 江的"生态源",是杭州市乃至浙江省的重要饮用水 源地(其淡水蓄水量约是太湖的4倍). 然而,随着 库区及上游流域社会经济的不断发展和资源的不合 理利用,水库的污染负荷的增加,水体的营养状况发 生了显著变化,局部水域出现了明显的富营养化迹 象,营养水平从建湖时期的贫营养逐步提高. 近几 年千岛湖库区水环境质量逐渐下降,据《中国环境 状况公报》数据表明,千岛湖局部水域水体从2000 年的Ⅰ类水质变为2009年的Ⅲ类水质,富营养化进 程加剧,总磷由 2000 年的 0.018 mg·L-1上升到目 前 0.037 mg·L<sup>-1</sup>, 总氮由 0.8 mg·L<sup>-1</sup>上升到 1.5 mg·L<sup>-1</sup>,叶绿素 a 由 2000 年的 2 μg·L<sup>-1</sup>上升到目 前 6 μg·L<sup>-1</sup>; 局部水域蓝藻水华频繁暴发. 但以往 对千岛湖的研究主要着重水质的变化或者仅关注浮 游植物的变化,遗憾的是,有可能引起水质变化的水 体分层现象的研究鲜有报道.

千岛湖水库的季节性分层及相应的水化学响应研究仍是空白.本研究以千岛湖水库湖泊区大坝采样点为对象,以 2011 年度水体温度、溶解氧(DO)、pH 值等水化学指标垂向分布特征为内容,阐明了千岛湖水库湖泊区水体分层结构发生、发展和消亡的年际演化过程,探讨水温分层对水库物化特征的影响作用,以期为亚热带特大型人工深水水库水环境特征的研究和水体富营养化的控制提供科学的参考.

#### 1 材料与方法

#### 1.1 水库概况

千岛湖原名新安江水库,位于浙江省淳安县境内(29°22′~29°50′N,118°34′~119°15′E),为1959年新安江水电站大坝建成后所形成的特大山谷型深水水库.新安江流域位于亚热带北缘、东南沿海季

风区,气候温暖,雨量充沛,年平均气温 16.90℃. 千岛湖正常高水位 108 m,库容 178. 4×10<sup>8</sup> m³,库水面 580 km²,平均水深 30.44 m,多年平均入库水量 94.5×10<sup>8</sup> m³,出库水量为 91.07×10<sup>8</sup> m³,水体交换周期长达 2 a. 千岛湖具有多年调节性能,正常情况下,出流量很小. 千岛湖泄水口高度在 20~30 m 水深处,但下泄水量发电属于调峰型电站,只在用电高峰期才运行. 千岛湖水量主要来自地表径流,共有大小入库支流 30 余条,其中新安江是最主要的入库地表径流,约占入库总径流量的 60%.

#### 1.2 研究方法

#### 1.2.1 数据的采集

2011年1~12月(采样时间为每月1号±2d 下午 12:00~13:00) 在千岛湖湖泊区大坝前 (29°30′36″N,119°12′37″E) 进行连续的监测(见图 1). 采样区域大坝前离泄水口至少5 km 以上,且在 近年的观测中,采样点处区域宽广,水体交换时间 长,出流量较小,理化性质和生物特征均较稳定. 因 此,采用湖泊区坝前作为研究样点,具有代表性和可 靠性. 本研究采用加拿大 RBR 公司的 XRX-620 快 速多参数水质剖面仪对垂直水体的温度、有效光合 作用辐照强度(PAR)、浊度、pH值、DO和叶绿素 a 进行数据的连续采集. XRX-620 是加拿大 RBR 公 司生产的可多参数集成的高精度仪器,采样速度快; 测量精度高;数据储存量大;仪器自动感应,下水 自动开始工作、出水自动停止工作; 具备水下电 缆,实时传输数据并供电,被广泛用于近海水质快速 调查、水库水质监测和湖泊水质监测等. 因此, XRX-620 仪器保证了数据的稳定性和可靠性.

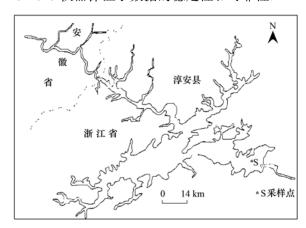


图 1 千岛湖地形图及采样点示意

Fig. 1 Topographical map and sampling site in Lake Qiandao

#### 1.2.2 数据处理

用 Microsoft Excel 2007 软件导出和处理数据,

用 ORIGIN 8.0 软件制图. 由于水质剖面仪测得的位置深度可精确到 ±0.01 m,为了方便比较,对导出的数据每米取平均值.

#### 2 结果与分析

#### 2.1 水温垂直分层与季节动态

水库水温在垂线方向上的变化,有在上层和下层缓慢下降,而在中层发生急剧变化的现象,此突变层称为温跃层.由于风的作用,上层水发生混合和温度变化,下层水始终保持稳定的状态,上下层之间便会产生温度急变的跃层.

一般来说,表层水体的水温受外界气候影响较大,且水库表层温度的变化会通过水体运输往下传递,但是底层水温变化主要受水库水体混合特性的影响.由图2可见,千岛湖水温季节分层特征明显,库区垂直水体温度属于暖单循环湖,即一年"翻库"一次.

冬季随着气温的快速下降,上层水温随之降低, 上下水层水体混合,分层现象消失,表层水和底层水 温差很小,形成较均匀的水体温度,水库在2月初进 人冬季混合期,整个库区水温为  $12.0\% \pm 0.97\%$ , 温度梯度很小,水深 5 m 以下温度变化小于 0.01 ℃·m<sup>-1</sup>,这表明千岛湖库区水体完全混合. 水库在 4月后随着气温的升高和太阳辐射的迅速增强,表 层温度开始逐渐升高,表层与底层的水温差距开始 增大,下层水温度变化很小,水体中逐渐出现水温分 层现象. 4 月开始温差接近 4℃,最大温度梯度 0.34℃·m<sup>-1</sup>出现在 5 m 处,水体热分层开始形成. 随着气温的进一步升高,千岛湖水库夏季的正向分 层则更加显著和稳定. 从图 2 中可以看出,5~8 月 千岛湖已形成较明显的温度梯度,由上至下的温度 梯度在5~15 m 水深范围内均达到最大值,最大温 度梯度 2.11 ℃·m<sup>-1</sup> 出现在 8 月水深 8~9 m 处.5 月和6月在5~20 m间形成了一个较为明显的温度 跃层; 7~8月在5~35 m处出现了明显的温跃层, 表层温度分别为 27.2℃和 32.7℃,且和底层水体的 温差均大于 16℃,8 月表底层温差达到 23℃,形成 相对稳定的"上高下低"水温分层结构,并在一定程 度上限制了上下层水团的交换. 9 月后表层温差逐 渐变小,变温层变厚,这是由于秋季随着气温的降 低,水库表层温度逐渐降低,较重冷水向下运输,温 度较高的底层水接受了上层水的冷却,温度分层逐 渐弱化. 9~10 月在 10~40 m 处依然存在较明显的 温跃层,且最大温度梯度发生在 10~15 m 水深处.

随着气温的进一步降低,11~12月千岛湖上混合层温差已基本消失,但中下层的温差和跃层现象依然存在,温跃层厚度开始变薄.

#### 2.2 水质剖面的垂向变化特征和季节动态

本研究水质测量都是从表层到最深处,图 3 和图 4 是 2011 年度水质参数(DO、pH值、浊度、PAR和叶绿素 a)的垂直剖面变化特征.可以看出,各个水质参数的垂向分布均存在较明显的季节性特征.

#### **2. 2. 1** PAR

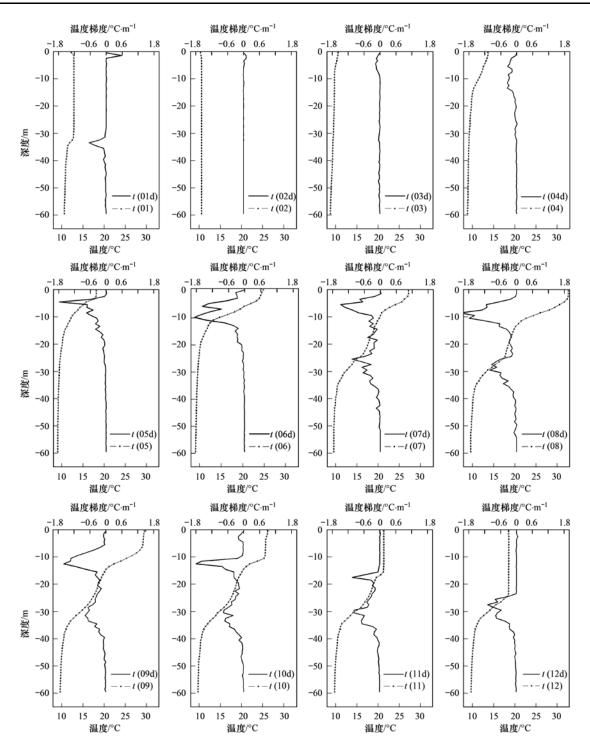
千岛湖水体纵向 PAR 变化特征见图 3. PAR 由于受水体中无机悬浮物、浮游植物的吸收和散射作用而发生衰减,其强度随着深度的增加而逐渐减低. 从图 3 中可以看出,PAR 在表层和中层水体对数递减,30 m 水层后基本接近 0. 以 8 月为例,PAR 在表层为 789. 41  $\mu$ mol·(s·m²)  $^{-1}$ , 10 m 水层达到 26. 51  $\mu$ mol·(s·m²)  $^{-1}$ , 到 20 m 水层降为 1. 50  $\mu$ mol·(s·m²)  $^{-1}$ ,在 30 m 水层以下基本接近 0. 从时间上来看,PAR 强度在 2 月最弱,最大值仅为 8. 85  $\mu$ mol·(s·m²)  $^{-1}$ , 10 月份表层 PAR 达到最大值,为1 413. 19  $\mu$ mol·(s·m²)  $^{-1}$ .

#### 2.2.2 浊度

千岛湖水体纵向浊度变化特征见图 3. 从中可以看出,浊度在垂直方向上是个比较稳定的参数,1~5月和11月、12月垂向变化非常小. 从中也可发现,7~8月温跃层内浊度达到全年的最大值,并且比其它月份高几倍,这是由于2011年6月月平均雨量达到137 mm,雨量充沛,较重的浊水流入水体而引起7月水库浊度的增加,并达到全年最大值,但7月30 m以下水体浊度仍保持较低水平,与其它月份相当,这也说明浊度的纵向变化受到水体分层的影响.

#### 2.2.3 pH 值

千岛湖水体纵向 pH 值的变化见图 3. 千岛湖水体 pH 值的垂向变化不是很明显,表层水呈微碱性,底层水偏微酸性. 这是因为水体中 pH 值的变化主要与藻类生长过程的光合作用有关. 表层水藻类通过光合作用,消耗水中的 CO<sub>2</sub>,破坏了 CO<sub>2</sub>/CO<sub>3</sub><sup>2-/</sup>HCO<sub>3</sub><sup>-</sup> 的平衡,使水体 pH 值呈微碱性<sup>[15]</sup>. 在分层水体中温跃层下 CO<sub>2</sub> 的累积使水体 pH 值略微降低,这是由于光不能穿透温跃层,下层水体 CO<sub>2</sub> 未被光合利用消耗<sup>[16]</sup>. 从中可以看出,pH 值在 0~5 m 出现递增趋势,并在 5 m 处达到一个极大值. 这是因为水体表面一般会出现光抑制现象,光合速率降低,而初级生产力最大值一般出现在次表层<sup>[17]</sup>. 千岛湖水库 pH 值在 5 m 处达到一个极大值,这是



t(数字)表示某数字月的水体温度,t(数字 d)表示某数字月的水体温度梯度

#### 图 2 千岛湖水温及温度梯度垂直剖面变化

Fig. 2 Vertical profile of water temperature and temperature gradient in Lake Qiandao

由于此深度范围内藻类生长最为旺盛,光合作用较强,消耗了水中的  $CO_2$ ,使水体 pH 值呈碱性,并达到一个极大值. 此外,  $1\sim3$  月在水库冬季混合期 5 m 深度以下 pH 值变化不大. 在  $4\sim8$  月水库水温分层形成和稳定期, pH 值的变化和水温的变化趋势基本一致. 在  $4\sim8$  月,  $0\sim5$  m 出现递增趋势,不同月

份表层水 pH 值逐渐增加,分别为 8.05、8.57、9.11、9.14 和9.73. 水深 5 m(温跃层上界)到 15 m 或 25 m 随着水深的增加而减小,这是由于此层内光合作用减弱,并且 CO, 的累积.

#### 2.2.4 溶解氧

千岛湖水体垂向 DO 分层存在明显的季节性分

布特征(见图 3). 1月、11月和12月表现出上层DO高、下层DO低即"上高下低"的现象,且分别在

30~38 m、15~20 m 和 25~30 m 水深处存在跃层, 其最大溶解氧梯度分别为 0.94、1.40 和 1.33

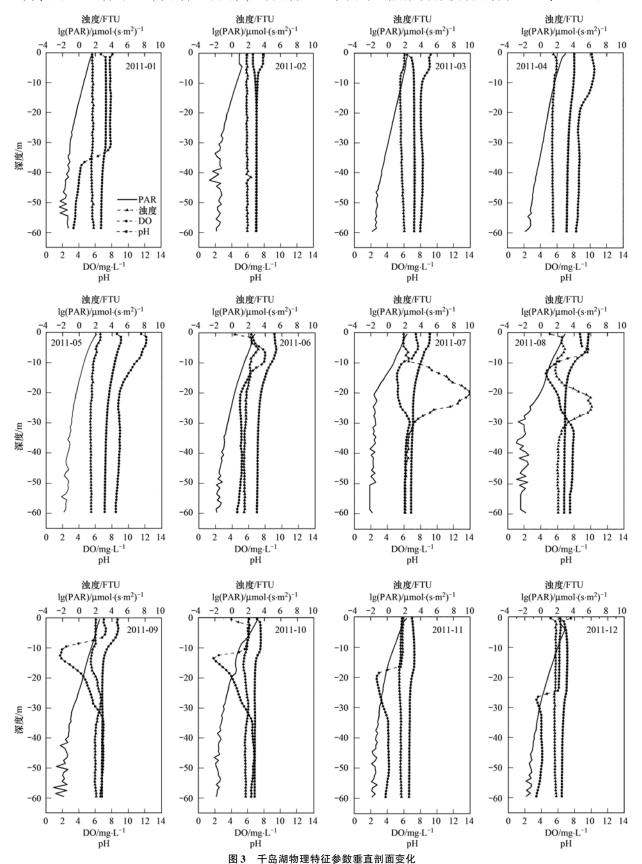
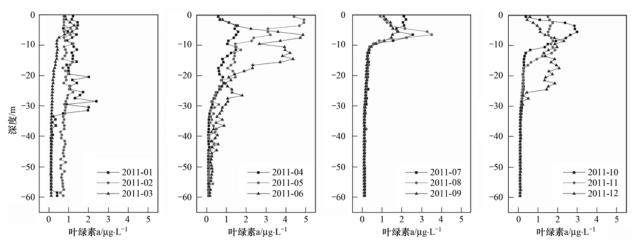


Fig. 3 Vertical profile of physical characteristics in Lake Qiandao

mg·(L·m) <sup>-1</sup>. 此外,底层 DO 浓度小于 4 mg·L<sup>-1</sup>, 进入缺氧状态. 2 月和 3 月上下水层 DO 浓度相差不大,其均值分别为 7. 13 mg·L<sup>-1</sup> ± 0. 22 mg·L<sup>-1</sup>和 8. 23 mg·L<sup>-1</sup> ± 0. 31 mg·L<sup>-1</sup>. 4~10 月水体存在了显著的 DO 纵向梯度:表层水体(10 m 以内)都接近饱和浓度,并且在此水层内 DO 逐渐递增;随着水深的增加,DO 浓度降低,并且在温跃层水深范围内急剧减小;温跃层以下深度 DO 基本恒定. 总体来说,DO 在 5~15 m 处骤降,而在 15~30 m DO 递增,30 m 下保持恒定,DO 随着深度的变化表现出先递减,再增加,而后保持恒定的趋势,底层 DO 浓度与表层相当.

#### 2.2.5 叶绿素 a

叶绿素 a 是表征浮游植物生物量的重要指标.



温较低的影响[12].

图 4 千岛湖水体叶绿素 a 垂直剖面变化

Fig. 4 Vertical profile of chlorophyll a in Lake Qiandao

#### 3 讨论

对于有一定水深的水库,水体的季节性分层可能对水库内几乎所有水环境演化过程起着控制或者影响作用.因此,对千岛湖水体水温分层的研究至关重要.从季节上看,千岛湖1~3月属于水库冬季混合期,水体分层完全消失,水库接近完全混合;4月属于千岛湖水库水体温度分层形成期,表层与底层的温度差逐渐增大;5~10月属于水库水体稳定分层期,水温在纵向上分为明显的3层:变温层、温跃层和恒温层;11~12月属于水库水体温度分层减弱期,温跃层的厚度逐渐变薄.一般判断水库水温分层类型采用指标法:

#### α = 年入库流量 / 总库容

当 $\alpha$ <10 时,为稳定分层型;当 $\alpha$ >20 时,为完全混合型<sup>[18]</sup>. 千岛湖库容 178.4×10<sup>8</sup> m<sup>3</sup>,多年平均入库

水量 94.  $5 \times 10^8$  m³,可以得到  $\alpha = 0.53$ ,远远小于 10,因此,千岛湖垂向水温应该存在稳定的分层. 这 与图 3 中所得的结果一致. 因此,千岛湖水体全年 水温分层时间长(4~12 月),混合期短(1~3 月), 是典型的亚热带单混合水体.

图 4 是千岛湖水体叶绿素 a 垂直剖面变化特征. 从

中可以看出,千岛湖叶绿素 a 分布出现明显且稳定

的分层现象,且和温跃层的变化特征较一致. 除了2

月外,30 m 水深以下的水体叶绿素 a 含量基本恒

定,且随季节变化幅度不大,最大值出现在冬季混合

期的 2 月(0.728  $\mu g \cdot L^{-1} \pm 0.089 \ \mu g \cdot L^{-1}$ ),紧随其后的是 6 月(0.455  $\mu g \cdot L^{-1} \pm 0.221 \ \mu g \cdot L^{-1}$ )和 5 月

(0.235 μg·L<sup>-1</sup> ± 0.116 μg·L<sup>-1</sup>),其余月份均维持

在 0.1 μg·L<sup>-1</sup>附近. 但表层水体的叶绿素 a 含量每

年会形成2个高峰,即春季末的高峰和夏末秋初的

高峰. 5 月和 6 月表层叶绿素 a 含量较高,最大值分别达到 4.932  $\mu g \cdot L^{-1}$  和 4.873  $\mu g \cdot L^{-1}$ . 冬季混合期

间没有引起浮游植物生物量的升高,这可能是受水

千岛湖水温分层结构,对水化学参数的垂直剖面变化有一定的影响. 浊度在温跃层内达到最大值也可以说明温跃层限制了物质向下传递,但是在此层内却得到了累积,这也说明浊度的纵向变化受到水体分层的影响[16]. 但从时间性来说,浊度大小主要受降雨所带来的外界输入的影响.

张运林等<sup>[7]</sup>对天目湖的研究中发现湖水温度 的变化对溶解氧的影响十分明显,温跃层存在时,下 层水的溶解氧明显低于上层水,当温跃层消失时,上 下层水的溶解氧差异不大.对千岛湖的研究也发 现,温度分层对溶解氧的影响作用显著.千岛湖水

体冬季上下水层溶解氧浓度相差不大,这主要是由 于冬季是千岛湖水体混合期,上下水体的混合复氧, 溶解氧浓度较高,2月和3月均值在7.13 mg·L<sup>-1</sup> ± 0. 22 mg·L<sup>-1</sup>和 8. 24 mg·L<sup>-1</sup> ± 0. 31 mg·L<sup>-1</sup>. 但随 着温度分层的形成,水体溶解氧也出现较明显的分 层现象. 一般来说,表层水体(10m 以内)都接近饱 和浓度,并且在此水层内溶解氧逐渐递增;随着水 深的进一步增加,溶解氧浓度逐步降低,并且在温跃 层范围内急剧减小;温跃层以下水体溶解氧变化不 大. 水体分层对溶解氧有一定的影响作用[12,13],特 别是水体存在稳定的温跃层时,上下水体不能进行 交换,阻止了上层水体的溶解氧向下层水体传递,上 层水体的气体交换和中上层浮游植物的光合作用为 上层水体提供了溶解氧,而深层的湖水则由于沉积 物有机质的矿化降解和大量有机体死亡下沉分解消 耗大量水体中的溶解氧,因此形成了显著的溶解氧 纵向梯度,出现上层、中层和下层水体不同的变化 特征,也会引起下层水体暂时性缺氧问题. Foley 等[19] 对英国某湖泊 41 a 数据进行分析,明确了温 度分层是底层缺氧的主导因素. Obenour 等[20] 和 Lee 等[21] 均指出温度分层是底层水体缺氧的一个主 要因素. 千岛湖水体存在稳定的温度分层现象,引 起溶解氧的分层,特别在冬季,下层 DO 浓度出现小 于4 mg·L-1,带来暂时周期性缺氧,这不利于水质 量的保障和水生生态系统的稳定. 此外,从本研究 中也可以发现(图 2 和图 3),温跃层内 DO 浓度较 低,且垂向 DO 最小值均出现在温跃层内,除了水温 分层结构限制了上下水体的交换,从而阻断了温跃 层水体 DO 的输入外,变温层有机物质沉降进入到 温跃层,消耗了温跃层的 DO,使得温跃层 DO 出现 一个最小值,这说明温度分层对水体溶解氧的有一 定的影响作用,分析千岛湖水环境状况需明确温跃 层内 DO 浓度.

另外,在适合藻类生长的春夏季节,从研究中发现水深 5 m 范围,溶解氧浓度出现了一个极大值,这是由于此深度范围藻类生长最为旺盛,光合作用增加了水体溶解氧浓度,pH 值在 5 m 水深处也达到了一个极大值,说明千岛湖水体 5 m 水深处适合藻类生长.此外,光抑制现象也可解释此结果[17].光合作用最强处一般不是出现在表层,或者说初级生产力或者浮游植物浓度最大值一般出现在次表层,这是因为在水体表面一般会出现光抑制现象,光合速率降低.从图 3 中可知,PAR 在表层出现最大值,且在水体中显示了固有的衰减规律,其衰减情况直接

影响到浮游植物光合作用过程和水体初级生产力的分布. 但水体中 PAR 最大点并不等同光合作用最强处. 吴志旭等<sup>[22]</sup>对千岛湖的研究中发现,浮游植物高密度区最大值出现在水深 5~10 m 处. 从 DO浓度和 pH 值的垂向分析也可知(图 3),光合作用最强处不是在 PAR 最大值处,这也间接说明了光抑制现象. 邱二生<sup>[23]</sup>对黑河和苏玉萍等<sup>[15]</sup>对山仔水库的研究也得到类似的结果. 因此,从以上分析中可以推断,由于光抑制存在,千岛湖水体光合作用最强处在 5~10 m 左右,此深度适合藻类生长,这为分析千岛湖水体蓝藻发生时取样和分析提供了依据.

从图 4 中可以看出,至春季温度一升高,表层叶 绿素 a 含量就直线上升,这可能是由于表层水温升 高和光照导致表层藻类大量增殖的缘故[12]. 随着 温度分层的形成和稳定,温跃层以上水深叶绿素 a 浓度平均值是温跃层以下水深的20倍,差值达2 μg·L-1,温跃层以下水深叶绿素 a 浓度较低且稳定 在 0.1 μg·L<sup>-1</sup>左右. 温跃层下 PAR 很小,浮游植物 光合作用较弱,温跃层限制了上下水体的交换,即限 制了浮游植物的垂向分布,使得温跃层下水深处叶 绿素 a 浓度得不到补充,并稳定在较低水平,这说明 深水湖库水体温跃层出现后,物质的垂直迁移受到 温跃层的阻碍作用,导致了叶绿素 a 浓度的分层. 此外,水体中的藻类组成和数量,受营养盐的上行作 用影响,分层水体中的营养盐状况是藻类在春夏暴 发的一个主要因素[24]. 千岛湖水库水温混合出现 在温度较低的冬季,混合期引起水中的营养物质沿 垂向均匀分布,使底层较高的营养物质上升到表层, 又因为温度较低,水中生化反应缓慢,消耗较少的营 养物质,至春季时沉积了大量的营养物质,为藻类快 速生长提供了物质基础,5~6月表层叶绿素 a浓度 达到全年最大值. 刘流等<sup>[25]</sup>和 Liu 等<sup>[26]</sup>对三峡水 库库首最大的支流香溪河库湾进行原位定点连续监 测,发现春季水温分层发育是春季水华暴发的直接 诱因. 因此,对千岛湖藻类增长的分析和探讨需建 立在水体分层研究的基础上.

#### 4 结论

- (1)千岛湖湖泊区水体全年水温分层时间长(4~12月),混合期短(1~3月),是典型的亚热带单混合水体.
- (2)水温分层结构对 DO、pH 值、浊度、PAR 和叶绿素 a 的垂直剖面变化有一定的影响作用,特别是夏季出现的水温显著分层现象,有效地限制了

上下水体的垂直交换,引起了水化学的分层,对水库水环境演化过程起着重要影响,尤其是 DO 和叶绿素 a.

- (3)湖泊区温跃层内 DO 浓度较低,冬季下层水体 DO 浓度小于 4 mg·L<sup>-1</sup>,出现暂时周期性缺氧,不利于水质量的保障和水生生态系统的稳定.
- (4)千岛湖湖泊区水体光合作用最强处在5~10 m 左右,此深度适合藻类生长,叶绿素 a 浓度在此深度范围内达到最大值,温跃层以下叶绿素 a 含量几乎一致且维持较低水平,这为分析千岛湖水体蓝藻暴发取样和分析提供了依据.

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