

湖泊热结构和蒸发的模拟计算

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摘要 应用一维涡扩散模型计算一年内不同时期的湖泊垂直温度分布和湖水蒸发率。主管方程是同一水平面内均温的一维非定常热传导方程。模型不要求特定的湖泊拟合参数。模型中涡扩散系数通过 Richardson 数计算, 水面热交换用能量平衡法计算, 湖底假定为绝热, 主管方程用有限差分法求解。对 Colorado City 湖和 Calhoun 湖的模拟计算表明, 水面温度和温度剖面的计算值与实测值吻合很好, 最大差值均在 2℃ 范围内。湖泊年蒸发量的计算值与实测值亦很接近, 差值仅为实测值的 5%。

关键词 湖泊, 热结构, 蒸发。

温度是重要的水质参数。水温的变化会影响水环境中各种物理、化学和生物过程, 影响溶解氧的含量。因此在湖泊的开发、管理和污染控制中都需要确定水温。

湖泊水温随季节变化; 对于中纬度地区较深的湖泊, 夏季前后一段较长的时间里水体还表现出温度分层的特点: 上层为变温层(Epilimnion), 下层为均温层(Hypolimnion), 中间是斜温层。秋季上层水体冷却下沉, 产生对流混合, 在冬季某一时候全湖才达到均温状态。

影响湖泊水温的因素很多, 一方面上层水体接收太阳和大气的辐射获得热量, 并通过扩散、流动和太阳辐射的水下吸收将热量向下层输运; 另一方面传导、长波辐射和蒸发又使湖水失去热量。

本文的目的在于考虑各种因素对湖泊水温的影响、应用一维涡扩散模型来计算湖泊热结构和蒸发。模型中涡扩散系数通过 Richardson 数计算, 水面热交换用能量平衡法计算, 蒸发用传质理论计算。方程用有限差分法求解。计算求得的温度值和蒸发量分别与实测值进行了比较。

1 数学模型

一维涡扩散模型的主管方程是同一水平面内均温的非定常热传导方程^[1]:

$$\frac{\partial T}{\partial t} = \frac{1}{A} \frac{\partial}{\partial z} [A(\alpha + K) \frac{\partial T}{\partial z}] - \frac{1}{A\rho c} \frac{\partial A\Phi}{\partial z} \quad (1)$$

式中, T 为水温(℃), t 是时间(d), z 是垂直坐标(m), 从水面向下为正, $A=A(z)$ 是深度 z 处湖泊横切面积(m²), α 是水的分子扩散系数(m²/d), K 是涡扩散系数(m²/d), ρ 是水的密度(kg/m³), c 为水的比热(J/(kg·℃)), Φ 是热源项(J/(m²·d)), 表示深度 z 处水吸收的太阳辐射热, 由下式确定^[2]:

$$\Phi = (1 - \beta)\theta_0 e^{-\eta z} \quad (2)$$

式中, β 是被表层吸收的太阳辐射部分, 其值随湖泊而变。 θ_0 是水面接收的太阳辐射热(J/(m²·d)), η 是光在水中的衰减系数(1/m), 其值取决于水的浊度。

方程(1)的初始条件可取为:

$$T(t, 0) = T_0 \quad (3)$$

这种条件可在春季找到, 这时全湖仍是均温的。边界条件, 在湖底取绝热条件:

$$\left. \frac{\partial T}{\partial z} \right|_{z=b} = 0 \quad (4)$$

在湖面边界条件是: 从水面向下扩散出去的热量等于表面水体吸收的太阳热 $\beta\theta_0$ 加上大气辐射热减去水面由于辐射、蒸发和传导引起的热损失:

$$-\rho c(\alpha + K) \frac{\partial T}{\partial z} \Big|_{z=0} = \beta Q_0 + (Q_a - Q_{ar}) - (Q_{br} + Q_e + Q_c) \quad (5)$$

式中, $Q_0 = Q_s - Q_{ar}$, Q_s 是到达水面的太阳辐射值, Q_{ar} 是被反射出去的阳光辐射值, Q_a 是大气长波辐射值, Q_{ar} 是被反射出去的大气长波辐射值, Q_{br} 是水面发出的长波辐射值, Q_e 和 Q_c 分别是蒸发和传导引起的热损失。各辐射值和热损失的单位均为 $(J/(m^2 \cdot d))$ 。

2 参数计算

2.1 涡扩散系数

湖泊中热的涡扩散可用涡扩散系数描述, 涡扩散系数由下式计算^[3]:

$$K = \frac{kw^*z}{P_0} \exp(-k^*z)(1 + 37R_i^2)^{-1} \quad (6)$$

式中, k 为 Von karman 常数, $k=0.4$, P_0 是 Prandtl 数, $P_0=1$, w^* 是水面摩擦速度 (m/s) , k^* 是 Ekman 剖面常数 $(1/m)$, R_i 是 Richardson 数。各参数计算如下:

$$w_* = 1.2 \times 10^{-3} u_2 \quad (7)$$

u_2 为水面上方 2m 处的风速 (m/s) ,

$$k^* = 6.6 \sqrt{\sin \Phi} u_2^{-1.84} \quad (8)$$

Φ 是湖泊所在地的纬度 (rad) ,

$$R_i = \frac{[1 + \frac{40N^2k^2z^2}{w^* \exp(-2k^*z)}]^{1/2} - 1}{20} \quad (9)$$

式中, N 是 Brunt-Vaisala 频率, 可表示为:

$$N^2 = \frac{g}{\rho} \frac{\partial \rho}{\partial z} \quad (10)$$

密度 ρ 是水温 (K) 的函数:

$$\rho = (1 - 1.9549 \times 10^{-5} |T - 227|^{1.68}) \times 10^3 \quad (11)$$

g 是重力加速度, $g=9.8m/s^2$ 。

2.2 表面热流

2.2.1 净太阳短波辐射 Q_0

Q_s 可用下式计算^[4]

$$Q_s = (0.25 + 0.50 \frac{n}{N_m}) R_\infty \quad (12)$$

式中, $\frac{n}{N_m}$ 是实际日照数 n 与最大可能日照数 N_m 之比, R_∞ 是大气顶层处的太阳辐射值 (W/m^2) :

$$R_\infty = 433.7477 \bar{d}^2 (H \sin \Phi \sin \delta + \cos \Phi \cos \delta \cos H) \quad (13)$$

此值是一天 24h 内的平均值, H 是半天长 (rad) , $\bar{d}$ 是日地间平均距离与瞬时距离之比, δ 为太阳倾角。 $\bar{d}$ 和 δ 可计算如下:

$$\bar{d} = [1 - 0.01673 \cos \frac{2\pi I}{365}]^{-1} \quad (14)$$

$$\delta = 0.4093 \sin \frac{2\pi(I - 79.75)}{365} \quad (15)$$

I 是 Julian 日, 半天长 H 用下式确定:

$$\cos H = -\tan \Phi \tan \delta \quad (16)$$

太阳短波辐射在水面的反射值 Q_{ar} 为:

$$Q_{ar} = aQ_s \quad (17)$$

a 为水的反射率, 对淡水本文取 $a=0.07$, 于是湖面上净的太阳短波辐射值 Q_0 为:

$$Q_0 = Q_s - Q_{ar} \quad (18)$$

2.2.2 净大气长波辐射

大气长波辐射 Q_a 可用 Stefan-Boltzman 定律来计算:

$$Q_a = \varepsilon_a \sigma (T_a + 273)^4 \quad (19)$$

式中, T_a 为空气温度 $(^\circ C)$, σ 是 Stefan-Boltzman 常数, $\sigma=4.916 \times 10^{-3} (J/(m^2 \cdot d \cdot K^4))$, ε_a 是大气辐射系数:

$$\varepsilon_a = 0.84 - (0.1 - 9.973 \times 10^{-6} e_a) \frac{n}{N_m} + 3.491 \times 10^{-5} e_a \quad \frac{n}{N_m} = 0.4 \quad (20)$$

$$e_a = 0.87 - (0.175 - 29.92 \times 10^{-6} e_a) \frac{n}{N_m} + 2.693 \times 10^{-5} e_a \quad \frac{n}{N_m} < 0.4 \quad (21)$$

式中, e_a 是空气蒸气压 (Pa)

$$e_a = \psi e_{as} \quad (22)$$

ψ 是相对湿度, e_{as} 是饱和蒸气压 (Pa) 。水面反射的长波辐射 Q_{ar} 为:

$$Q_{ar} = A_r Q_a \quad (23)$$

A_r 是长波反射率, 在水面 $A_r=0.03$, 于是净长波辐射为 $Q_a - Q_{ar}$ 。

2.2.3 水面长波辐射 Q_{br}

$$Q_{br} = \varepsilon \sigma (T_s + 273)^4 \quad (24)$$

ε 是水的辐射系数, 取 $\varepsilon = 0.97$, T_s 是水面温度 ($^{\circ}\text{C}$)。

2.2.4 蒸发热流

蒸发从水面失去的热流 Q_e 为:

$$Q_e = \rho L_e E \quad (25)$$

式中, L_e 是蒸发潜热 (J/kg), 可用下式计算^[4]:

$$L_e = 1.91846 \times 10^6 \left[\frac{T_s + 273}{T_s + 239.09} \right]^2 \quad (26)$$

E 是蒸发率 (m/d), 由传质理论计算:

$$E = NU_2(e_0 - e_a) \quad (27)$$

式中, N 是经验确定的传质系数 ($\text{s}/(\text{Pa} \cdot \text{d})$):

$$N = 1.4581A_0^{-0.05} \times 10^{-6} \quad (28)$$

A_0 是湖泊水面面积 (km^2), e_0 是水面饱和蒸气压 (Pa), e_a 为空气蒸气压 (Pa)

$$e_0 = 101325 \exp(13.3185t_k - 1.976t_k^2 - 0.6445t_k^3 - 0.1299t_k^4) \quad (29)$$

式中, $t_k = 1 - \frac{373.15}{T_s + 273}$, e_a 用下式计算:

$$e_a = e_a^* - \gamma(T_s - T_{wb}) \quad (30)$$

e_a^* 是空气湿泡温度时的饱和蒸气压 (Pa), 用 (29) 式计算, γ 是湿度计常数, 本文取 $\gamma = 67.48245 (\text{Pa}/^{\circ}\text{C})$, T_s 是空气干泡温度 ($^{\circ}\text{C}$), T_{wb} 为湿泡温度 ($^{\circ}\text{C}$)。

3.2.5 对流热流

水面对流损失的热流 Q_c 为:

$$Q_c = RQ_s \quad (31)$$

R 为 Brown 比:

$$R = \gamma(T_s - T_a)(e_0 - e_a) \quad (32)$$

3 数值方法

热传导方程 (1) 可用显式有限差分法求解, 和隐式比较, 显式更方便, 计算中产生的数值误差亦可控制在允许的范围内。本文采用显式, 计算中取距离步长 1m , 最大时间步长为 1d , 为了满足数值稳定性条件, 本文还使用了变时间步长的方法。稳定性条件为:

$$\Delta t \leq \frac{\Delta z^2}{2(\alpha + K)} \quad (33)$$

式中, $\Delta t(\text{d})$ 、 $\Delta z(\text{m})$ 分别是时间和深度的计算步长。

在对方程 (1) 的离散化过程中, 时间的偏微分用前差近似, 空间的偏微分用中心差分近似:

$$\frac{\partial T}{\partial t} = \frac{T_i^{j+1} - T_i^j}{\Delta t} \quad (34)$$

上标 j 和下标 i 分别表示时间层和空间层。令

$$K_e = \alpha + K \quad (35)$$

则:

$$\begin{aligned} \frac{1}{A} \frac{\partial}{\partial t} [A(\alpha + K) \frac{\partial T}{\partial z}] &= \frac{1}{A} \frac{\partial}{\partial z} [AK_e \frac{\partial T}{\partial z}] \\ &= \frac{2}{A_{i+1}A_{i-1}} \cdot 0.5 \\ &\frac{(A_{i+1} + A_i)K_h(T_{i+1} - T_i) - (A_i + A_{i-1})K_{h-1}(T_i - T_{i-1})}{\Delta z^2} \end{aligned} \quad (36)$$

$$\frac{1}{A\rho c} \frac{\partial(A\Phi)}{\partial z} = (1 - \beta)Q_0 \cdot$$

$$\frac{\exp[-\eta(z_i + 0.5)] \frac{A_{i+1} + A_i}{2} - \exp[-\eta(z_i - 0.5)] \frac{A_i + A_{i-1}}{2}}{A_i \rho c_i \Delta z} \quad (37)$$

将方程 (34), (35), (36) 和 (37) 代入方程 (1) 中, 便可得到温度的数值解:

$$\begin{aligned} T_i^{j+1} &= T_i^j + \left\{ \frac{2}{A_{i+1}A_{i-1}} \cdot \right. \\ &\frac{(A_{i+1} + A_i)K_h(T_{i+1} - T_i) - (A_i + A_{i-1})K_{h-1}(T_i - T_{i-1})}{\Delta z^2} - \\ &(1 - \beta)Q_0 \cdot \left[\frac{\exp[-\eta(z_i + 0.5)] \frac{A_{i+1} + A_i}{2}}{A_i \rho c_i \Delta z} - \right. \\ &\left. \left. \frac{\exp[-\eta(z_i - 0.5)] \frac{A_i + A_{i-1}}{2}}{A_i \rho c_i \Delta z} \right] \right\} \end{aligned} \quad (38)$$

式中, $c_i (\text{J}/(\text{kg} \cdot \text{K}))$ 计算如下:

$$c_i^j = 4172 + 1.6648f_i^j \quad (39)$$

$$f_i^j = \exp(r_i^j/10.6) + \exp(-r_i^j/10.6) \quad (40)$$

$$r_i^j = 34.5 - T_i^j, \quad T_i^j < 34.5^{\circ}\text{C} \quad (41)$$

$$r_i^j = 2.08(T_i^j - 34.5)^{0.67}, \quad T_i^j \geq 34.5^{\circ}\text{C} \quad (42)$$

4 模型验证

将一个模型用来预测不同条件下的水质之前, 需对模型进行验证。验证本文建立的模型的数据取自 Texas 州的 Colorado city 湖 (北纬 33° ,

湖面积 8.2km²,最大深度 15.85m)的现场观测结果^[5]。每天的输入值由线性内插求得,用这些数据 $\eta=0.5, \beta=0.3$ 算得的水面温度及温度剖面与实测相比吻合很好,差值仅在 2℃ 以内,见图 1 和图 2。模拟计算的全年蒸发量为 2026mm,现场观测值 2125mm,相差仅 5%。

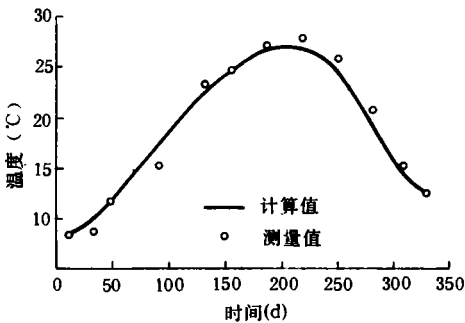


图 1 Colorado City 湖水面温度变化曲线

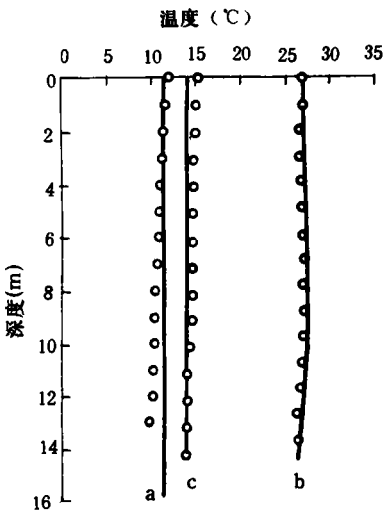


图 2 Colorado City 湖温度剖面计算值(实线)和测量值(黑点)的比较
a. 3月1日 b. 7月19日 c. 11月16日

5 模型应用

Calhoun 湖热结构的计算

Calhoun 湖位于 Minnesota 州的 Minneapolis 市,北纬 45°,水面面积为 1.71km²,最大深度 25m。

模拟计算时间为 183d,起始日期为 1974 年

4月15日,此时全湖均温,水温为 6.4℃,计算截止日期为 11月15日。

气象观测数据如空气温度、相对湿度、风速和云量取自观测值^[6]。计算中每天的输入值用气象参数每月的平均值进行线拟合得到。

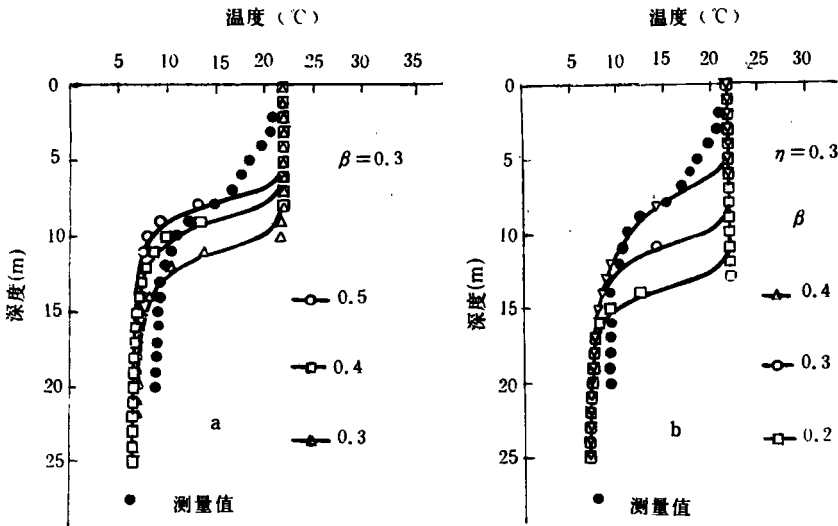


图 3 参数 β 和 η 灵敏度分析(1974 年 6 月 25 日)

参数 β 和 η 对热分层的影响如图 3 所示。由图 3 可知 η 增大,变温层将变薄,但对变温层和均温层中温度分布形态没有影响。 β 对热分层的影响与 η 相同。由图可知能得到最佳结果的 β 和 η 值分别是 0.4 和 0.3。

图 4 表明模拟计算的水面温度与实测值吻

合很好。只有在秋季计算值稍低于测量值,相差在 2°C 范围内。

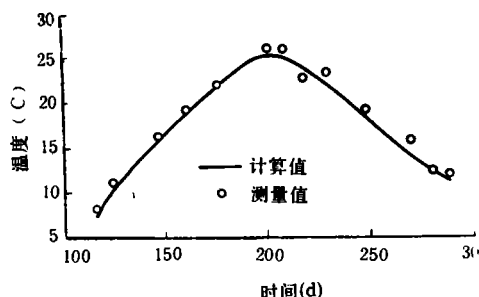


图 4 Calhoun 湖水温度变化曲线

图 5 是分层期的温度剖面分布。由图可见在变温层内温度分布的计算值与实测值吻合很好,差别出现在斜温层和均温层。在均温层中计算值稍低于测量值,相差在 2°C 范围内。

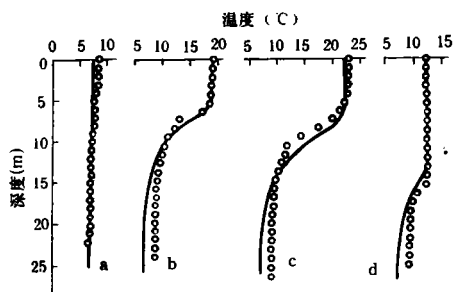


图 5 Calhoun 湖温度剖面计算值(实线)和测量值(圆圈)的比较
a. 4月26日 b. 6月9日 c. 8月16日 d. 10月7日

图 6 显示了模拟计算得到的涡扩散系数剖面。由图可见变温层内,春季紊流扩散系数大于夏季时的紊流扩散系数,这是由于春季有较强的风吹过湖面之故。较高的涡扩散系数之值所反映的强烈的混合抑制了大的温度梯度的形成,因此春季分层不明显(图 5 中 4 月 26 日)。夏季气温升高,风速又明显降低,使湖面紊流扩散减弱,达到一定深度后浮力梯度的作用使紊流扩散的影响为零,导致了斜温层的形成(图 5 中 6 月 9 日和 8 月 16 日)。随着时间的推移,斜温层逐渐下

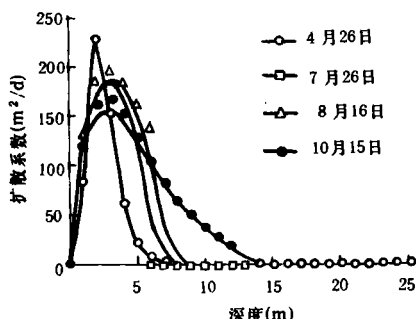


图 6 涡扩散系数的计算值

在整个模拟计算期间(4月15日—11月15日)蒸发量的计算值为 955mm。

6 结论

一维涡扩散模型能准确反映湖泊热量的平衡过程和输运机理。用实测值对模型进行验证和参数校正后,该模型便是计算湖泊热结构和蒸发量的有效工具。

模拟计算的温度剖面取决于湖面水体对太阳辐射的吸收系数 β 和光在水中的衰减系数 η 。湖泊不同, β 和 η 之值可能不同。

中纬度地区,较深湖泊的温度分层从春季开始,夏季十分明显,秋季后减弱,冬季某个时候起,全湖处于均温状态。

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Abstracts

Chinese Journal of Environmental Science

terms of the ratio of the concentration of dissolved ion (mg/L) to the activated sludge (kg VS/L).

Key words: anaerobic system, heavy metal, zinc.

Analysis and Prediction on the Traffic Noises from an Elevated Express Compound Highway. Chen Ziming, Lu Deming et al. (Dept. of Physics, Qingdao University of Oceanography, Qingdao 266003); *Chin. J. Environ. Sci.*, 15(2), 1994, pp. 24—28

Based on a section of the Jinan-Qingdao Express Highway in the suburbs of Qingdao City which is a typical compound highway of elevated structure, an analysis has been made on the relationship between its traffic noises and the factors involved. A mathematical model has been also established for predicting the traffic noises from an elevated express compound highway. It is found that the theoretical calculations based on this model are generally coincident with the actually simulated measurements.

Key words: elevated express compound highway, insonified zone, peak volume hour, running noise of a single vehicle.

Influence of Different Carbon Sources on Denitrification. Xu Yatong. (Dept. of Environmental Science, East China Normal University, Shanghai 200062); *Chin. J. Environ. Sci.*, 15(2), 1994, pp. 29—32

In a suspended sludge system, the rates of denitrification and carbon uptake can follow a zero-order reaction when the carbon supply was from an unlimited, single kind of carbon source at a suitable pH and temperature. It was found that a mixed volatile fatty acid (VFA) C-source resulted in a higher rate of denitrification than any of the individual VFA C-sources which formed the mixed one. Of the VFAs, acetate gave the highest rate of denitrification. The VFAs gave a higher rate of denitrification than their corresponding alcohols. The means of reducing the carbon consumption and the cost of treatment in a denitrification system had been also discussed.

Key words: biological denitrification, suspended sludge, carbon source, rate of denitrification, C-uptake rate.

Simulation of Thermal Structure and Evaporation for Lakes. Congzhi Zhou et al. (Teaching and Research Section for Water & Heating Logistic Engineering University, Chongqing 630041); *Chin. J. Environ. Sci.*, 15(2), 1994, pp. 33—37

An one-dimensional eddy diffusion model was used

for prediction of the annual vertical temperature profile and evaporation rate of a lake. The governing equation was a non-linear heat transfer equation assuming horizontal homogeneity. No lake-specific fitting of the parameters of the model was necessary. Eddy diffusivities were computed with Richardson Number. The heat exchange at the water surface was formulated by the energy balance approach, and the bottom of the lake was considered to be insulated. The governing equation was solved by finite difference. Computation was made for Lakes Colorado City and Calhoun. Quantitative and qualitative agreement between computed and measured temperature profiles was very good, with an error of less than 2°C. The computed annual evaporation capacity approached the one measured and the difference between them was only 5% of the measurement value.

Key words: lake, thermal structure, evaporation.

Promoting the Electrolysis Efficiency of a Bipolar Electrolyser by Adding a Coated Activated Carbon. Zhou Kanghan, Zhou Ding. (Harbin Institute of Technology, Harbin 150006); *Chin. J. Environ. Sci.*, 15(2), 1994, pp. 38—40

It has been found that improving the conditions of contact between particles in a bipolar electrolyser is a key to promoting its electrolysis efficiency. Such an improvement has been achieved by adding a coated activated carbon to the electrolyser and the satisfactory results were obtained. With a Proper formulation, the addition of a coated activated carbon at a ratio of 100 : 30 allowed the electrolysis efficiency to increase from 130% to 245% and a stable, uniform distribution could be kept during an operation for about 300h.

Key words: electrolysis efficiency, bipolar electrolyser, coated activated carbon.

Ozone Layer Depletion, Aerosol Pollution and Solar UV-B Radiation at the Ground. Wang Shaobin (Research Center for Eco-Environmental Sciences, Chinese Academy of Sciences, Beijing 100085); *Chin. J. Environ. Sci.*, 15(2), 1994, pp. 41—44

Ozone layer depletion and aerosol pollution have an influence on the solar ultraviolet radiation reaching the ground. A computation indicates that the radiation amplification factor (RAF) for a DNA damage varies from 2.3 to 4.2 when a reduction in column ozone goes from 1% to 30%. Since the industrial revolution, the daily UV-B radiation in urban and rural areas has decreased by about 45% and 10%, respectively, that can be attributed to aerosol pollution. The computed UV-B radiation due to the ozone reduction in Beijing shows an increasing